

## **Cambrian granulite to upper amphibolite facies metamorphism of post-797 Ma sediments in Madagascar**

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### **ABSTRACT**

The CHIME monazite and xenotime ages were determined for two paragneisses from Madagascar. The biotite-sillimanite-cordierite-quartz-microcline-plagioclase schist at Ihosy in southern Madagascar contains core-mantle monazite grains as well as chronologically unzoned ones. The unzoned monazite grains (8 of 11 analyzed grains) give a CHIME age of  $527 \pm 15$  Ma. Two zoned grains show core ages of  $1640 \pm 180$  and  $797 \pm 75$  Ma with mantles of  $550 \pm 50$  and  $535 \pm 10$  Ma, respectively. The kyanite- and corundum-bearing cordierite-sillimanite-muscovite-biotite-plagioclase-quartz gneiss at Maevatanana in northwestern Madagascar contains both monazite and xenotime. Monazite and xenotime grains are chronologically uniform, and give CHIME ages of  $534 \pm 10$  and  $530 \pm 32$  Ma, respectively. These chronological data suggest that the high-grade paragneisses in Madagascar, if not all, formed through a single thermal event at ca. 530 Ma from post-797 Ma sediments. This metamorphism can be linked to the continental collision that resulted in Gondwana supercontinent.

### **INTRODUCTION**

Gondwana supercontinent existed over 300 m.y. before its break-up in the Mesozoic. The evolution of this supercontinent, particularly of East Africa, Madagascar, Southern India, Sri Lanka and East Antarctica, was recently discussed on the basis of lithology, metamorphism and tectonics (e.g. Yoshida et al., 1992; Windley et al., 1994; Shiraishi et al., 1994; Kröner et al., 1996). The broad aspects of the picture for the Gondwana connection are generally agreed, but considerable debates still remain in detail, especially the location of the collision zone. Since Gondwana supercontinent formed through the late Cambrian (-latest Proterozoic) continental collision of East and West Gondwana, the distribution of Cambrian metamorphic rocks may play a key role in the analysis of the collision tectonics.

Recent geochronological studies have brought forth Cambrian metamorphic ages for high-grade paragneisses in the Kerala khondalite belt of southernmost India (Choudhary et al., 1992; Santosh et al., 1992; Bindu et al., 1996). Similar high-grade paragneisses are known to occur in Madagascar (Windley et al., 1994). Some of them were dated to be around 550 Ma (e.g. Kröner et al.,

1996, and their references). But little is known about the extension of the ca. 550 Ma metamorphic rocks in Madagascar. In order to shed more light on the collision tectonics of Gondwana, we have made the CHIME dating of monazite and xenotime from two samples of high-grade paragneisses from Madagascar.

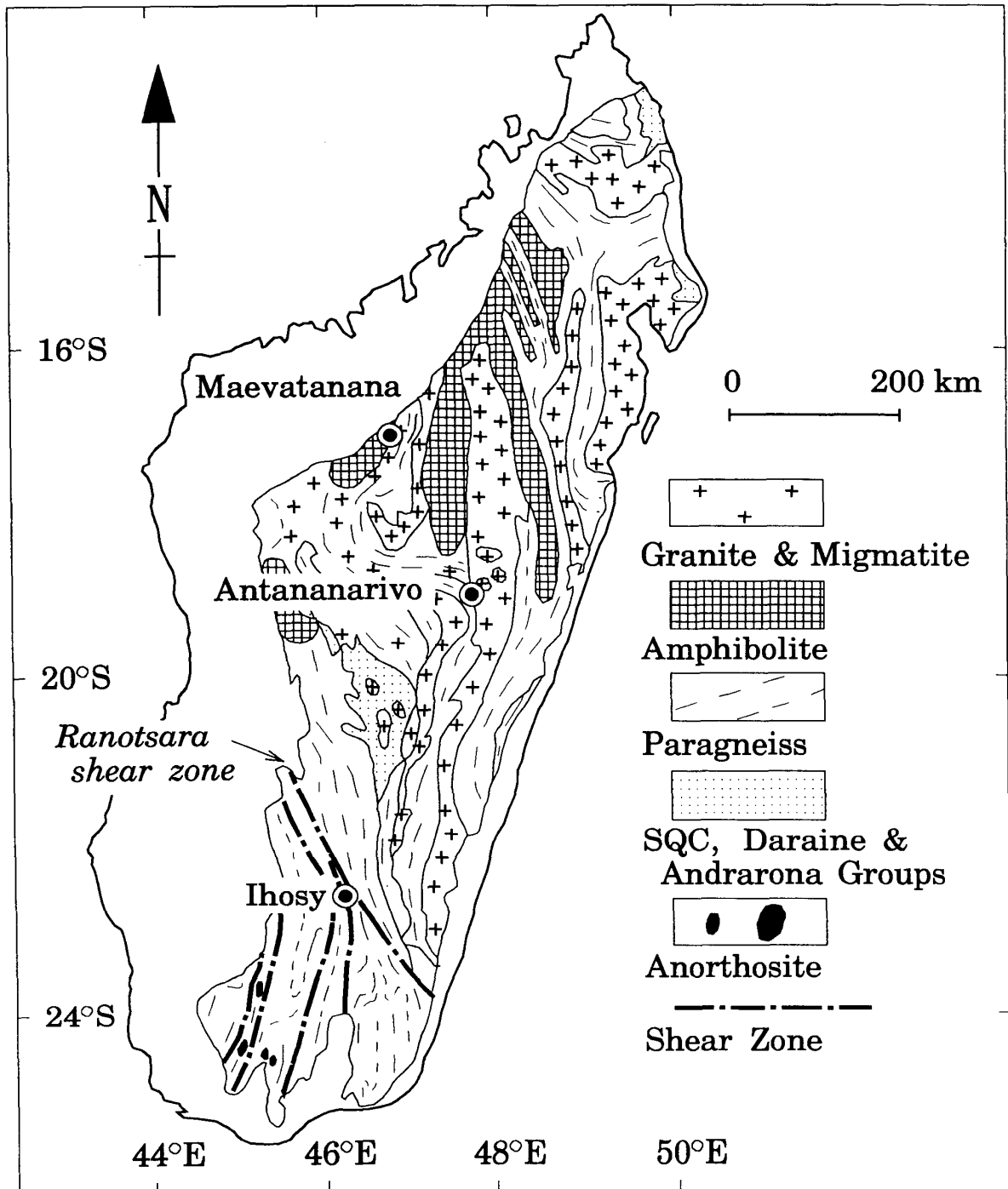


Fig. 1. Simplified and schematic geologic map showing the main basements of Madagascar (from Windley et al., 1994).

## GENERAL GEOLOGY AND SAMPLE DESCRIPTION

The basement rocks in Madagascar are divided into the southern and central-northern sectors by the northwest trending Ranotsara shear zone (Windle et al., 1994; Fig. 1). The southern sector consists predominantly of high-grade metamorphic rocks formed under the granulite and upper amphibolite facies conditions. Particularly common are cordierite- and sillimanite-bearing paragneisses with intercalated marble, metaquartzite and amphibolite layers. They generally trend in NS direction. The central-northern sector is underlain mainly by granitoids and amphibolite (greenstone belt) with a subordinate amount of paragneiss. This sector is traversed by a north-trending 100–150 km wide dextral shear zone of probable Pan-African age; the western margin of the shear zone runs through around Antananarivo. Rocks in the dextral shear zone were metamorphosed under the granulite and upper amphibolite facies conditions, and include reworked older basement. Two paragneiss samples were collected for the present CHIME dating; biotite-sillimanite-cordierite gneiss from Ihosy in the southern sector and kyanite- and corundum-bearing cordierite-sillimanite-muscovite-biotite gneiss from Maevatanana in the western part of the central-northern sector.

### *Biotite-sillimanite-cordierite-quartz-microcline-plagioclase gneiss from Ihosy*

This rock is composed of plagioclase, microcline, quartz, biotite, cordierite, sillimanite and opaque mineral, together with small amounts of zircon, apatite and monazite. Plagioclase, sodic oligoclase in composition, is granoblastic and is slightly sericitized. Biotite, sillimanite and quartz together with fluid inclusions are found within plagioclase grains. Myrmekite is rarely developed. Microcline is commonly string and film perthite with well-developed microcline twinning. It includes sillimanite and plagioclase. Quartz shows often undulatory extinction and includes sillimanite, biotite, plagioclase, microcline, apatite and opaque mineral. Cordierite occurs as granoblastic grains with inclusions of sillimanite, biotite, quartz and opaque mineral. Some cordierite grains are twinned. It is slightly altered. Sillimanite is coarse-grained, sometimes euhedral or subhedral, and includes quartz, biotite, zircon and opaque mineral. It shows often undulatory extinction. The sillimanite inclusions in cordierite, quartz and biotite are of small and short prismatic form, and are frequently oriented. Biotite is oriented along the schistosity. It is pleochroic; X = nearly colourless, Y = brown and Z = light brown. Symplectic biotite is also found. Inclusions in biotite are sillimanite, apatite, zircon, monazite, and opaque mineral. Muscovite is very small in amount, and may be derived from the alteration of plagioclase. Monazite is anhedral, and ranges in size from 0.02 to 0.15 mm.

### *Kyanite- and corundum-bearing cordierite-sillimanite-muscovite-biotite-plagioclase-quartz gneiss from Maevatanana*

Biotite-rich layers are finely alternated with feldspathic layers which are somewhere coarse-grained. This rock is composed of plagioclase, quartz,

biotite, muscovite, cordierite and sillimanite together with accessory corundum, kyanite, zircon, apatite, monazite, xenotime and opaque mineral. Plagioclase (oligoclase) is granoblastic and is slightly sericitized. It includes biotite, sillimanite, muscovite and zircon. Quartz is also granoblastic with inclusions of biotite, sillimanite and muscovite. Biotite is commonly associated with muscovite and sillimanite. It is pleochroic; X = nearly colourless, Y = yellowish olive green and Z = olive green. Inclusions in biotite are apatite, zircon, monazite and opaque mineral. Sillimanite is fibrous to very long prismatic, intimately intergrown with muscovite and biotite. Cordierite is granoblastic with inclusions of biotite, sillimanite and opaque mineral. It is moderately, sometimes completely, altered. Muscovite flakes include plagioclase, quartz, biotite, sillimanite, corundum and kyanite. Corundum and kyanite are always coated with muscovite. Monazite and xenotime occur as tiny grains, but exceptionally large grains exceed 0.04 mm in size.

### CHIME AGES

Monazite and xenotime grains in polished thin sections, prepared for the conventional electron microprobe analyses, were analyzed on JEOL JXA-733 electron microprobe equipped with three wavelength dispersive-type spectrometers. Instrument operating conditions were 15 kv accelerating voltage, 0.02–0.15  $\mu\text{A}$  current and 5  $\mu\text{m}$  probe diameter. X-ray intensities were integrated 200s for the lines and 100s for backgrounds at two optimum positions on both sides of the lines. The detection limits of  $\text{ThO}_2$ ,  $\text{UO}_2$  and  $\text{PbO}$  at 2  $\sigma$  confidence level are 0.012, 0.007 and 0.004 wt.%, respectively, and possible errors in the determination of  $\text{ThO}_2$ ,  $\text{UO}_2$  and  $\text{PbO}$  are about 14% for 0.1 wt.% of the concentrations. The analytical results are listed in Table 1. CHIME ages were calculated through the method described by Suzuki and Adachi (1991a,b, 1994), Adachi and Suzuki (1992) and Suzuki et al. (1994).

#### *Biotite-sillimanite-cordierite-quartz-microcline-plagioclase gneiss from Ihosy*

Eleven monazite grains were analyzed. Nine of these grains are chronologically unzoned, and two grains show a core-mantle relation. Figure 2 shows the  $\text{PbO}$  vs.  $\text{ThO}_2^*$  plots of chronologically unzoned monazite grains. Data points, except those for grain margins (open circle), are arrayed linearly on the diagram. They are regressed with an isochron of  $527 \pm 15$  Ma (MSWD = 0.28) with an intercept value of  $-0.0010 \pm 0.0045$ .

Figure 3a shows the  $\text{PbO}$  vs.  $\text{ThO}_2^*$  plots of analytical data of a core-mantle monazite grain (M02). Data points for the core (solid square) and the mantle (solid circle) are separately arrayed, and define different isochrons of  $1640 \pm 180$  Ma (MSWD = 0.66) and  $550 \pm 50$  Ma, respectively. Data points between the 1640 and 550 Ma isochrons (open square, cross and triangle) are for the core-mantle boundary region, but do not define either a certain area in the cross section or any isochrons. The textural feature suggests an overlap of the 550 Ma mantle on the 1640 Ma core at analyzed points on the cross

Table 1. Microprobe analyses of ThO<sub>2</sub>, UO<sub>2</sub> and PbO of monazite (M) and xenotime (X) from metamorphic rocks at Ihosy and Maevatanana, Madagascar.

Spot No.	ThO <sub>2</sub> (wt.%)	UO <sub>2</sub> (wt.%)	PbO (wt.%)	Age (Ma)	ThO <sub>2</sub> * (wt.%)	Spot No.	ThO <sub>2</sub> (wt.%)	UO <sub>2</sub> (wt.%)	PbO (wt.%)	Age (Ma)	ThO <sub>2</sub> * (wt.%)		
<b>Biotite-sillimanite-cordierite gneiss (Ihosy)</b>						M02-35	3043	6.17	0.228	0.491	1604	7.00	
M01-01	6.38	0.084	0.149	525	6.66	M02-36	3049	5.58	0.166	0.276	1039	6.16	
M01-02	6.32	0.049	0.147	534	6.48	M02-37	3057	5.59	0.135	0.136	529	6.04	
M01-03	6.43	0.079	0.152	532	6.70	M02-38	3063	5.80	0.136	0.142	533	6.25	
M01-04	5.30	0.175	0.134	534	5.88	M02-39	3069	5.69	0.170	0.124	468	6.25	
M01-05	6.24	0.145	0.144	503	6.72	M02-40	3601	5.74	0.089	0.133	519	6.03	
M01-06	5.73	0.170	0.139	517	6.29	M02-41	3609	5.51	0.102	0.134	538	5.85	
M01-07 r	5.46	0.117	0.120	482	5.84	M02-42	3616	6.28	0.153	0.174	600	6.79	
M01-08 r	5.47	0.174	0.119	462	6.05	M02-43	3625	5.62	0.225	0.461	1634	6.44	
M01-09 r	5.67	0.170	0.129	488	6.23	M02-44	3631	5.53	0.302	0.471	1622	6.64	
M01-10 r	5.63	0.207	0.126	469	6.31	M02-45	3637	5.10	0.637	0.511	1577	7.42	
M01-11 r	5.75	0.191	0.130	479	6.38	M02-46	3644	6.03	0.237	0.449	1497	6.89	
M02-01	4030	6.98	0.348	0.574	1592	8.25	M02-47	3651	6.04	0.112	0.163	597	6.41
M02-02	4033	7.26	0.245	0.572	1605	8.15	M02-48	3657	5.65	0.135	0.138	531	6.10
M02-03	4038	7.29	0.246	0.587	1638	8.19	M02-49	3663	5.74	0.139	0.137	519	6.19
M02-04	4041	7.48	0.276	0.568	1534	8.48	M02-50	3669	5.52	0.167	0.112	434	6.07
M02-05	2034	7.09	0.335	0.510	1415	8.28	M02-51	4305	6.61	0.112	0.163	546	6.98
M02-06	4628	6.06	0.143	0.194	693	6.54	M02-52	4311	6.43	0.119	0.151	520	6.82
M02-07	5330	6.72	0.112	0.163	538	7.09	M02-53	4319	5.90	0.169	0.149	540	6.46
M02-08	2510	6.80	0.114	0.157	515	7.17	M02-54	4336	5.42	0.254	0.306	1125	6.31
M02-09	1349	6.72	0.138	0.164	536	7.18	M02-55	4345	5.54	0.107	0.130	519	5.90
M02-10	5343	6.82	0.154	0.169	541	7.34	M02-56	4351	5.61	0.105	0.137	540	5.96
M02-11	0236	6.31	0.174	0.135	460	6.88	M02-57	4357	5.54	0.132	0.134	527	5.98
M02-12	0826	6.21	0.143	0.130	458	6.68	M02-58	4366	5.48	0.177	0.116	448	6.06
M02-13	0836	6.15	0.150	0.150	530	6.65	M02-59	5014	5.46	0.124	0.121	485	5.87
M02-14	0641	6.00	0.123	0.133	488	6.41	M02-60	5023	4.15	0.093	0.103	541	4.46
M02-15	0852	5.67	0.170	0.139	524	6.23	M02-61	5037	5.28	0.108	0.126	527	5.64
M02-16	1417	5.57	0.141	0.131	510	6.04	M02-62	5049	5.62	0.137	0.137	530	6.07
M02-17	1425	5.26	0.133	0.131	540	5.70	M02-63	5056	5.48	0.251	0.143	532	6.31
M02-18	1433	5.53	0.097	0.127	511	5.85	M02-64	5809	5.73	0.150	0.118	446	6.23
M02-19	1453	5.52	0.148	0.130	509	6.01	M02-65	5823	5.27	0.128	0.122	505	5.69
M02-20	1459	5.25	0.127	0.119	495	5.67	M02-66	5840	5.81	0.153	0.128	476	6.31
M02-21	2206	5.50	0.133	0.129	509	5.94	M02-67	6415	5.63	0.142	0.140	538	6.10
M02-22	2212	5.56	0.124	0.137	539	5.97	M02-68	6425	5.71	0.136	0.126	480	6.15
M02-23	2219	5.69	0.100	0.129	503	6.02	M02-69	6429	5.95	0.152	0.121	442	6.46
M02-24	2224	4.10	0.040	0.119	658	4.24	M02-70	6436	3.05	0.765	0.102	432	5.56
M02-25	2238	5.40	0.247	0.438	1590	6.30	M02-71	6728	5.80	0.162	0.142	526	6.33
M02-26	2254	5.72	0.129	0.143	546	6.15	M02-72	4645	5.75	0.113	0.137	524	6.12
M02-27	2259	5.12	0.137	0.113	477	5.57	M02-73	4045	5.85	0.125	0.140	524	6.26
M02-28	3001	5.76	0.106	0.127	489	6.11	M02-74	3345	6.60	0.253	0.545	1655	7.53
M02-29	3007	5.76	0.082	0.127	496	6.03	M02-75	2445	5.59	0.148	0.136	526	6.08
M02-30	3013	5.72	0.101	0.139	538	6.05	M02-76	1945	5.77	0.133	0.134	508	6.21
M02-31	3020	5.47	0.098	0.131	529	5.80	M02-77	2152	5.98	0.157	0.149	538	6.50
M02-32	3026	5.84	0.382	0.528	1667	7.24	M02-78	2752	5.65	0.123	0.132	510	6.06
M02-33	3031	5.42	0.591	0.546	1647	7.58	M02-79	3252	5.89	0.173	0.233	840	6.48
M02-34	3038	6.10	0.271	0.503	1624	7.09	M02-80	3752	5.76	0.121	0.142	542	6.16

Table 1. (continued).

Spot No.	ThO <sub>2</sub> (wt.%)	UO <sub>2</sub> (wt.%)	PbO (wt.%)	Age (Ma)	ThO <sub>2</sub> * (wt.%)	Spot No.	ThO <sub>2</sub> (wt.%)	UO <sub>2</sub> (wt.%)	PbO (wt.%)	Age (Ma)	ThO <sub>2</sub> * (wt.%)		
M02-81	4352	5.93	0.104	0.136	511	6.27	M08-13	6.74	0.131	0.161	526	7.17	
M02-82	4362	5.80	0.141	0.140	524	6.27	M08-14	6.88	0.141	0.156	498	7.35	
M02-83	3762	5.78	0.133	0.144	542	6.22	M08-15	6.69	0.138	0.155	510	7.15	
M02-84	3262	5.77	0.140	0.138	520	6.24	M08-16	6.67	0.136	0.159	523	7.12	
M02-85	2538	5.95	0.194	0.473	1623	6.66	M08-17	6.63	0.144	0.152	502	7.11	
M02-86	2530	6.15	0.308	0.509	1602	7.27	M08-18	6.81	0.126	0.164	534	7.23	
M02-87	2030	5.52	0.567	0.480	1457	7.56	M08-19	6.63	0.123	0.157	525	7.04	
M03-01		5.48	0.122	0.131	525	5.88	M08-20	6.77	0.128	0.157	513	7.20	
M03-02		5.72	0.168	0.136	509	6.28	M09-01	6.41	0.167	0.154	521	6.96	
M03-03		5.56	0.188	0.134	511	6.18	M09-02	6.45	0.161	0.150	503	6.98	
M03-04		5.35	0.177	0.140	554	5.94	M09-03	6.55	0.181	0.161	528	7.15	
M03-05		5.66	0.149	0.142	540	6.16	M09-04	6.53	0.158	0.163	542	7.06	
M04-01		5.68	0.172	0.138	519	6.25	M09-05	6.46	0.163	0.155	521	7.00	
M04-02		5.59	0.191	0.138	519	6.22	M09-06	6.44	0.142	0.151	513	6.91	
M04-03		5.58	0.177	0.143	544	6.17	M09-07	6.39	0.173	0.159	537	6.96	
M04-04		5.60	0.159	0.134	515	6.12	M09-08	6.27	0.141	0.153	532	6.74	
M04-05		5.65	0.171	0.131	495	6.21	M09-09	5.60	0.185	0.140	529	6.22	
M04-06		5.55	0.177	0.145	554	6.14	M09-10	6.56	0.175	0.160	527	7.14	
M04-07		5.77	0.199	0.134	491	6.42	M09-11	6.42	0.165	0.161	543	6.97	
M04-08		5.69	0.172	0.143	535	6.26	M09-12	6.61	0.146	0.156	518	7.10	
M04-09		5.73	0.206	0.144	526	6.41	M09-13	6.48	0.176	0.161	535	7.06	
M04-10		5.65	0.192	0.145	540	6.29	M09-14	5.85	0.191	0.148	537	6.48	
M05-01		2.90	0.328	0.083	492	3.98	M09-15	5.64	0.200	0.142	529	6.31	
M05-02		3.08	0.237	0.091	552	3.87	M10-01	7.25	0.136	0.170	518	7.70	
M05-03		2.97	0.252	0.081	502	3.80	M10-02	10.5	0.195	0.253	531	11.2	
M05-04		3.26	0.261	0.093	528	4.13	M10-03	13.0	0.233	0.313	534	13.8	
M05-05		3.31	0.274	0.094	521	4.22	M10-04	9.00	0.157	0.206	509	9.52	
M06-01		7.52	0.139	0.178	525	7.98	M10-05	9.05	0.150	0.216	531	9.55	
M06-02		7.56	0.124	0.187	550	7.97	M10-06	9.07	0.166	0.216	526	9.62	
M06-03		7.45	0.204	0.178	515	8.13	M10-07	6.45	0.176	0.160	534	7.03	
M06-04		7.53	0.154	0.181	529	8.04	M10-08	6.50	0.170	0.152	505	7.06	
M06-05		7.45	0.126	0.178	530	7.87	M10-09	6.55	0.120	0.152	515	6.95	
M07-01		6.85	0.163	0.163	518	7.40	M10-10	6.67	0.119	0.155	517	7.07	
M07-02		6.25	0.140	0.155	540	6.71	M10-11	6.69	0.165	0.162	526	7.24	
M08-01	r	6.66	0.120	0.140	468	7.06	M10-12	r	6.58	0.125	0.137	460	7.00
M08-02		7.03	0.141	0.158	496	7.49	M10-13	5.93	0.140	0.143	526	6.40	
M08-03		6.96	0.123	0.165	524	7.37	M10-14	5.50	0.150	0.135	531	5.99	
M08-04		6.67	0.125	0.153	509	7.08	M10-15	6.22	0.137	0.147	519	6.67	
M08-05		7.00	0.120	0.163	518	7.39	M10-16	5.93	0.153	0.147	536	6.43	
M08-06		6.84	0.124	0.155	503	7.25	M10-17	6.58	0.123	0.160	537	6.99	
M08-07		6.63	0.128	0.162	539	7.05	M10-18	11.6	0.181	0.273	527	12.2	
M08-08		6.75	0.130	0.160	525	7.18	M10-19	5.36	0.140	0.127	510	5.83	
M08-09		6.70	0.137	0.161	529	7.15	M10-20	5.71	0.142	0.140	533	6.18	
M08-10		6.45	0.144	0.157	532	6.92	M10-21	5.77	0.147	0.139	523	6.26	
M08-11		6.63	0.136	0.146	485	7.08	M10-22	4.94	0.148	0.124	537	5.43	
M08-12		6.74	0.126	0.163	533	7.16	M10-23	6.34	0.169	0.157	533	6.90	

Table 1. (continued).

Spot No.	ThO <sub>2</sub> (wt.%)	UO <sub>2</sub> (wt.%)	PbO (wt.%)	Age (Ma)	ThO <sub>2</sub> * (wt.%)	Spot No.	ThO <sub>2</sub> (wt.%)	UO <sub>2</sub> (wt.%)	PbO (wt.%)	Age (Ma)	ThO <sub>2</sub> * (wt.%)
M10-24	4.49	0.321	0.126	531	5.56	M11-40 4364	3.75	0.079	0.0917	537	4.01
M10-25	6.31	0.132	0.149	518	6.75	M11-41 5103	7.25	0.131	0.177	541	7.68
M10-26	7.21	0.163	0.173	526	7.74	M11-42 5107	7.58	0.138	0.184	536	8.04
M10-27 r	6.44	0.387	0.155	471	7.71	M11-43 5121	8.46	0.117	0.203	538	8.85
M10-28	6.64	0.115	0.160	536	7.02	M11-44 5135	10.5	0.384	0.411	813	11.8
M10-29	6.63	0.139	0.159	525	7.09	M11-45 5143	8.86	0.500	0.243	542	10.5
M10-30	5.19	0.687	0.169	532	7.47	M11-46 5151	8.16	0.275	0.205	532	9.07
M11-01 0632	6.64	0.181	0.163	530	7.24	M11-47 5158	8.03	0.146	0.196	541	8.51
M11-02 0638	6.66	0.185	0.161	520	7.27	M11-48 5164	6.79	0.157	0.169	544	7.31
M11-03 0646	6.29	0.166	0.152	523	6.83	M11-49 6002	7.19	0.126	0.173	533	7.61
M11-04 0654	5.97	0.192	0.149	531	6.60	M11-50 6013	8.31	0.132	0.201	539	8.75
M11-05 0051	5.27	0.429	0.144	507	6.69	M11-51 6026	10.7	0.371	0.408	795	12.0
M11-06 1030	6.64	0.159	0.160	523	7.17	M11-52 6033	10.4	0.346	0.393	792	11.6
M11-07 1034	6.64	0.145	0.158	522	7.12	M11-53 6044	8.29	0.141	0.199	535	8.76
M11-08 1043	6.67	0.164	0.164	533	7.21	M11-54 6053	8.17	0.137	0.198	539	8.63
M11-09 1051	6.36	0.169	0.158	535	6.92	M11-55 6060	7.01	0.151	0.171	535	7.51
M11-10 1056	6.06	0.168	0.157	557	6.62	M11-56 6064	5.03	0.150	0.128	545	5.52
M11-11 1814	6.69	0.097	0.140	470	7.01	M11-57 6903	8.46	0.141	0.204	536	8.93
M11-12 1819	6.76	0.118	0.164	539	7.15	M11-58 6911	7.79	0.108	0.184	531	8.15
M11-13 1829	7.07	0.134	0.167	522	7.51	M11-59 6917	11.0	0.358	0.420	801	12.2
M11-14 1837	7.21	0.131	0.174	534	7.64	M11-60 6926	9.70	0.320	0.371	804	10.8
M11-15 1844	7.21	0.116	0.179	553	7.59	M11-61 6935	8.34	0.129	0.201	539	8.76
M11-16 1855	6.48	0.161	0.163	544	7.02	M11-62 6942	8.09	0.132	0.191	525	8.52
M11-17 1861	6.67	0.158	0.167	546	7.19	M11-63 6949	8.20	0.132	0.195	531	8.63
M11-18 1869	4.63	0.143	0.106	488	5.10	M11-64 6955	7.64	0.139	0.187	542	8.10
M11-19 2608	6.80	0.156	0.160	512	7.31	M11-65 6961	2.22	0.245	0.0709	549	3.03
M11-20 2615	7.25	0.124	0.173	529	7.66	M11-66 7809	7.74	0.126	0.183	528	8.15
M11-21 2621	7.07	0.125	0.173	543	7.49	M11-67 7817	9.33	0.358	0.360	798	10.5
M11-22 2632	7.06	0.131	0.173	541	7.49	M11-68 7825	9.35	0.366	0.365	805	10.6
M11-23 2641	7.68	0.116	0.183	534	8.07	M11-39 7833	8.84	0.326	0.343	803	9.95
M11-24 2648	7.06	0.104	0.171	541	7.41	M11-70 7841	8.67	0.151	0.209	534	9.17
M11-25 2655	6.61	0.119	0.160	536	7.00	M11-71 7849	7.97	0.130	0.188	527	8.40
M11-26 2663	6.58	0.170	0.158	518	7.15	M11-72 7856	3.85	0.309	0.114	549	4.88
M11-27 2671	4.96	0.343	0.140	539	6.10	M11-73 8815	11.0	0.347	0.277	534	12.2
M11-28 3506	7.05	0.134	0.174	545	7.50	M11-74 8820	9.10	0.387	0.237	536	10.4
M11-29 3513	7.23	0.121	0.176	541	7.63	M11-75 8833	11.0	0.422	0.284	539	12.4
M11-30 3521	7.36	0.077	0.173	532	7.62	M11-76 8839	8.98	0.218	0.221	535	9.70
M11-31 3529	7.56	0.126	0.183	538	7.98	M11-77 8844	6.39	0.172	0.161	542	6.96
M11-32 3536	7.62	0.140	0.184	534	8.08	M11-78 8850	2.32	0.261	0.0729	537	3.19
M11-33 3543	7.74	0.327	0.203	541	8.83	M11-79 9628	6.56	0.254	0.170	539	7.40
M11-34 3553	7.32	0.103	0.173	529	7.66	M11-80 9633	5.61	0.247	0.146	535	6.42
M11-35 4300	7.30	0.121	0.157	481	7.69	M11-81 9638	5.82	0.425	0.164	534	7.23
M11-36 4322	7.50	0.128	0.182	541	7.92	M11-82 9641	4.28	0.465	0.138	556	5.82
M11-37 4340	8.23	0.453	0.230	556	9.73						
M11-38 4349	7.03	0.306	0.179	521	8.05						
M11-39 4357	7.69	0.124	0.185	537	8.10						

Table 1. (continued).

Spot No.	ThO <sub>2</sub> (wt.%)	UO <sub>2</sub> (wt.%)	PbO (wt.%)	Age (Ma)	ThO <sub>2</sub> * (wt.%)	Spot No.	ThO <sub>2</sub> (wt.%)	UO <sub>2</sub> (wt.%)	PbO (wt.%)	Age (Ma)	UO <sub>2</sub> * (wt.%)
<b>Kyanite- and corundum-bearing cordierite- sillimanite-muscovite gneiss (Maevatanana)</b>						X01-01	0.411	1.49	0.123	545	1.61
M01-01	10.7	1.34	0.350	544	15.1	X01-02	0.477	1.64	0.136	542	1.79
M01-02	10.8	1.31	0.342	531	15.1	X01-03	0.482	2.06	0.166	537	2.21
M01-03	10.4	1.19	0.323	529	14.3	X01-04	0.397	1.41	0.112	521	1.53
M01-04	10.6	0.890	0.308	534	13.5	X01-05	0.285	1.17	0.093	532	1.26
M01-05	11.1	1.13	0.343	542	14.9	X01-06	0.132	1.25	0.097	538	1.29
M01-06	13.9	0.787	0.378	537	16.5	X01-07	0.142	1.19	0.093	541	1.23
M01-07	12.9	0.787	0.361	545	15.5	X01-08	0.148	1.13	0.088	532	1.18
M01-08	11.7	0.769	0.327	540	14.2	X01-09	0.239	1.26	0.103	548	1.33
M01-09	10.2	0.974	0.306	537	13.4	X01-10	0.439	1.82	0.148	540	1.95
M01-10	10.8	1.23	0.339	536	14.9	X01-11	0.390	1.46	0.117	533	1.57
M01-11	12.9	0.941	0.364	532	16.1	X01-12	0.373	1.71	0.136	534	1.82
M01-12	16.0	0.655	0.417	538	18.2	X01-13	0.364	1.37	0.113	545	1.48
M01-13	8.88	0.599	0.251	542	10.9	X02-01	0.216	1.41	0.115	556	1.47
M01-14	13.6	1.06	0.396	543	17.1	X02-02	0.203	1.27	0.102	546	1.33
M01-15	11.4	0.781	0.318	535	14.0	X02-03	0.153	1.58	0.118	522	1.62
M01-16	13.0	1.34	0.394	533	17.4	X02-04	0.128	1.49	0.112	525	1.53
M01-17	10.8	1.37	0.352	539	15.3	X02-05	0.278	1.66	0.132	541	1.74
M02-01	11.6	0.938	0.332	532	14.7	X02-06	0.342	1.32	0.107	535	1.43
M02-02	11.8	0.941	0.339	533	14.9						
M02-03	12.0	0.932	0.338	529	15.0						
M02-04	7.09	1.28	0.259	536	11.3						
M02-05	10.9	1.05	0.324	532	14.3						
M02-06	10.3	1.12	0.320	536	14.0						
M02-07	9.58	1.50	0.331	534	14.5						
M02-08	9.81	1.59	0.343	535	15.1						
M02-09	8.97	1.36	0.305	530	13.5						
M02-10	7.31	1.52	0.281	535	12.4						
M02-11	8.05	1.46	0.290	528	12.9						
M02-12	7.85	0.776	0.236	531	10.4						
M02-13	6.85	0.560	0.202	544	8.71						
M02-14	9.36	1.07	0.295	537	12.9						
M02-15	8.11	1.42	0.291	532	12.8						
M02-16	9.02	1.45	0.313	531	13.8						
M02-17	8.90	1.44	0.310	532	13.7						
M02-18	11.9	0.904	0.333	527	14.8						
M02-19	12.2	0.758	0.334	532	14.8						
M02-20	12.0	0.869	0.343	541	14.9						
M02-21	11.6	1.02	0.339	531	15.0						
M02-22	8.84	1.33	0.300	533	13.2						
M02-23	11.4	1.03	0.337	535	14.8						
M02-24	10.6	1.24	0.341	545	14.7						
M02-25	11.3	1.06	0.335	530	14.9						

**Note**

ThO<sub>2</sub>\*: sum of the measured ThO<sub>2</sub> and ThO<sub>2</sub> equivalent of the measured UO<sub>2</sub>.

UO<sub>2</sub>\*: sum of the measured UO<sub>2</sub> and UO<sub>2</sub> equivalent of the measured ThO<sub>2</sub>.

Age: apparent age calculated through equation 1 of Suzuki and Adachi (1991a,b, 1994), Suzuki et al. (1994) and Adachi and Suzuki (1992).

r: rim

Analyzed position is shown for M02 and M11 monazite grains: for example, M02-01 4030 means 40 μm from right and 30 μm from top.



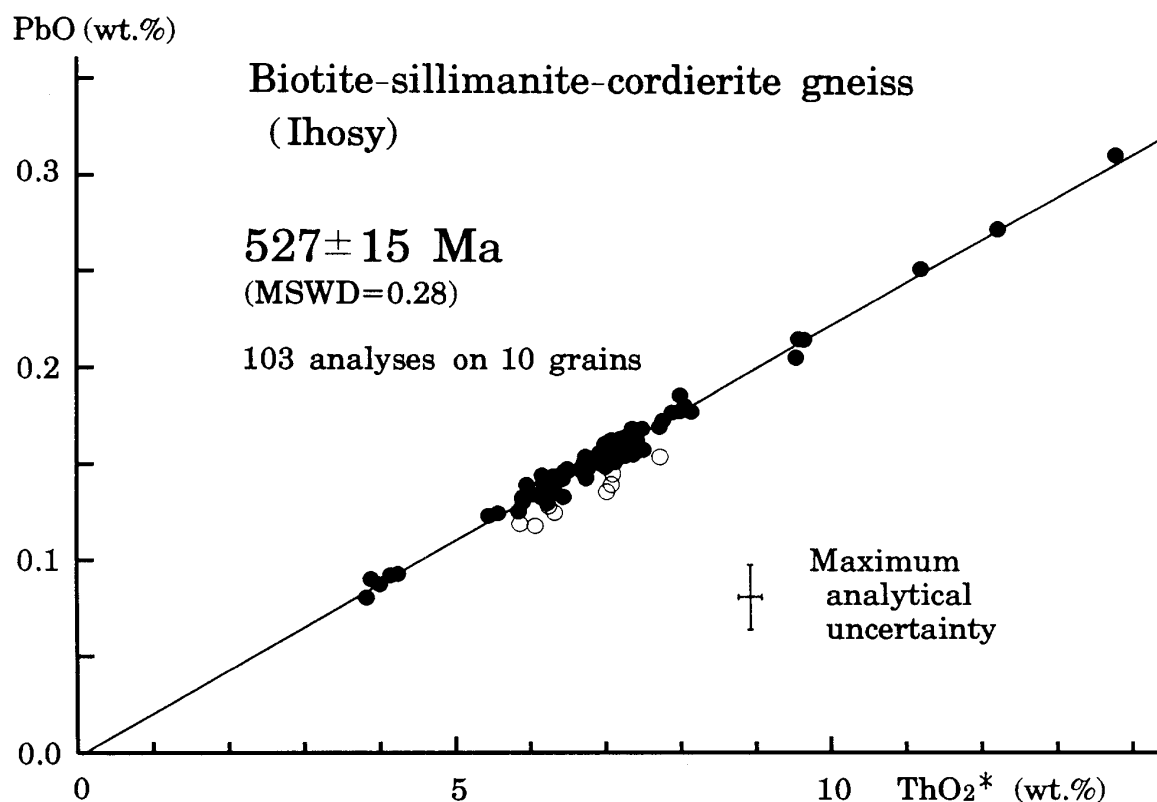


Fig. 2. Plot of PbO vs.  $\text{ThO}_2^*$  for unzoned monazite grains from biotite-sillimanite-cordierite gneiss from Ihosy. Open circle represents data points for grain margins. Error bars in the figure represent  $2\sigma$  analytical uncertainty, and error quoted in age is of  $2\sigma$ .

section. Data points (open circle) below the 550 Ma isochron may result from Pb loss.

Figure 3b shows the PbO vs.  $\text{ThO}_2^*$  plots of analytical data of a core-mantle monazite grain (M11). Within this grain  $\text{UO}_2$  is heterogeneously distributed with respect to the shape of the cross section, showing a homogeneous concentration (0.320–0.384 wt.%) in the core and variable concentrations (0.077–0.500 wt.%) in the mantle. The  $\text{ThO}_2$  content of the core (8.843–11.04 wt.%) is higher than that of the mantle (2.22–11.03 wt.% and mostly in the range between 6.5 and 8.5 wt.%). This compositional discontinuity suggests overgrowth of the mantle under a different chemical environment. The CHIME ages are  $797 \pm 75$  Ma for the core and  $535 \pm 10$  Ma for the mantle.

*Kyanite- and corundum-bearing cordierite-sillimanite-muscovite-biotite-plagioclase-quartz gneiss from Maevatanana*

Most monazite grains from this sample are very small in size, but two grains have a size sufficient for microprobe analyses. Seventeen spots were measured on M01-grain and 25 spots on M02-grain. The  $\text{ThO}_2$  and  $\text{UO}_2$  distributions are highly variable within individual grains (Table 1), but data points are arrayed linearly on the PbO vs.  $\text{ThO}_2^*$  diagram (Fig. 4a). The best

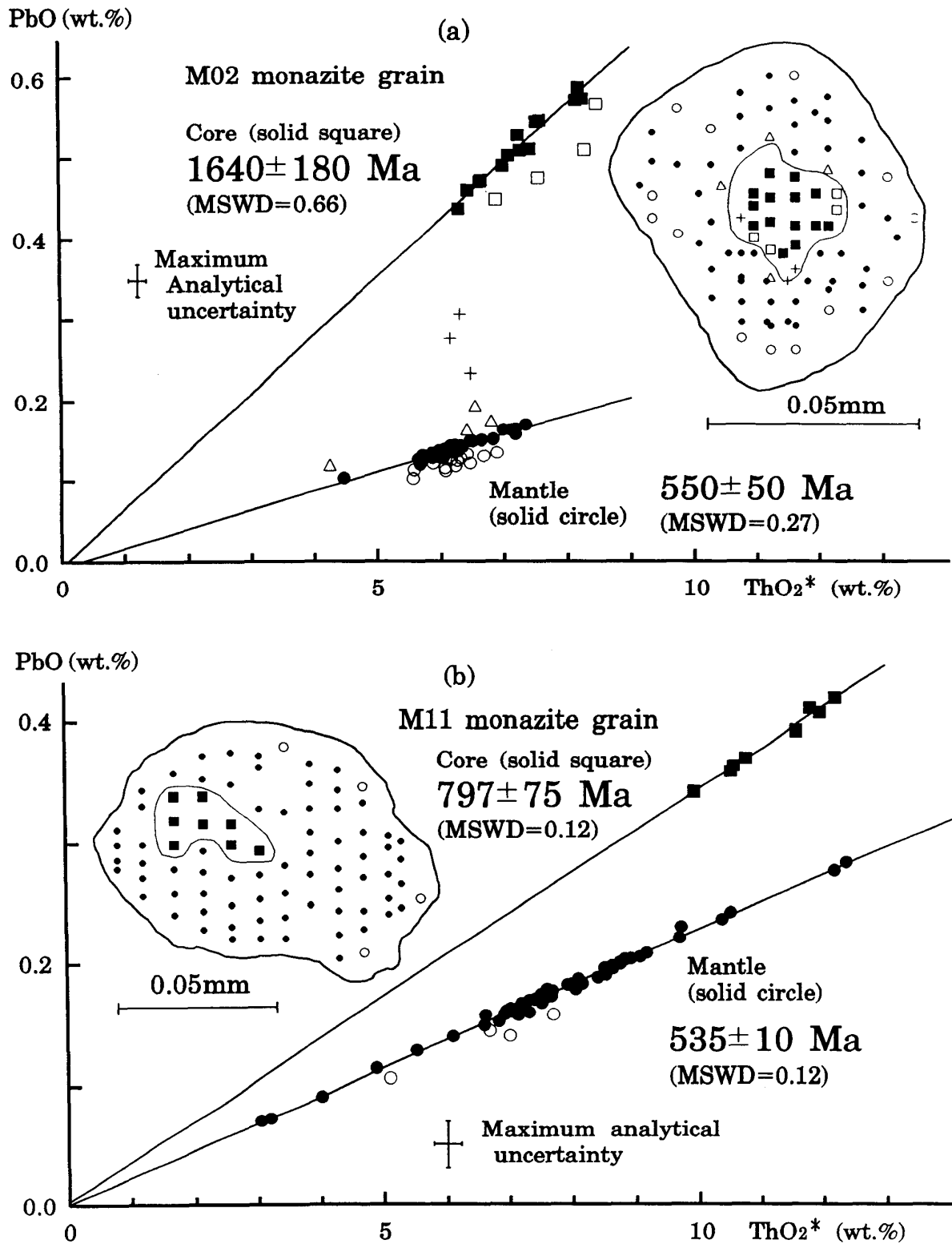


Fig. 3. Plot of PbO vs. ThO<sub>2</sub>\* for core-mantle monazite grains M02 (a) and M11(b) from biotite-sillimanite-cordierite gneiss from Ihosy. Solid square represents data points for the core and solid circle represents those for the mantle. Data points for grain margin are shown with open circle, and those for the core-mantle boundary with open square and cross. Explanation for errors is the same as Fig. 2.

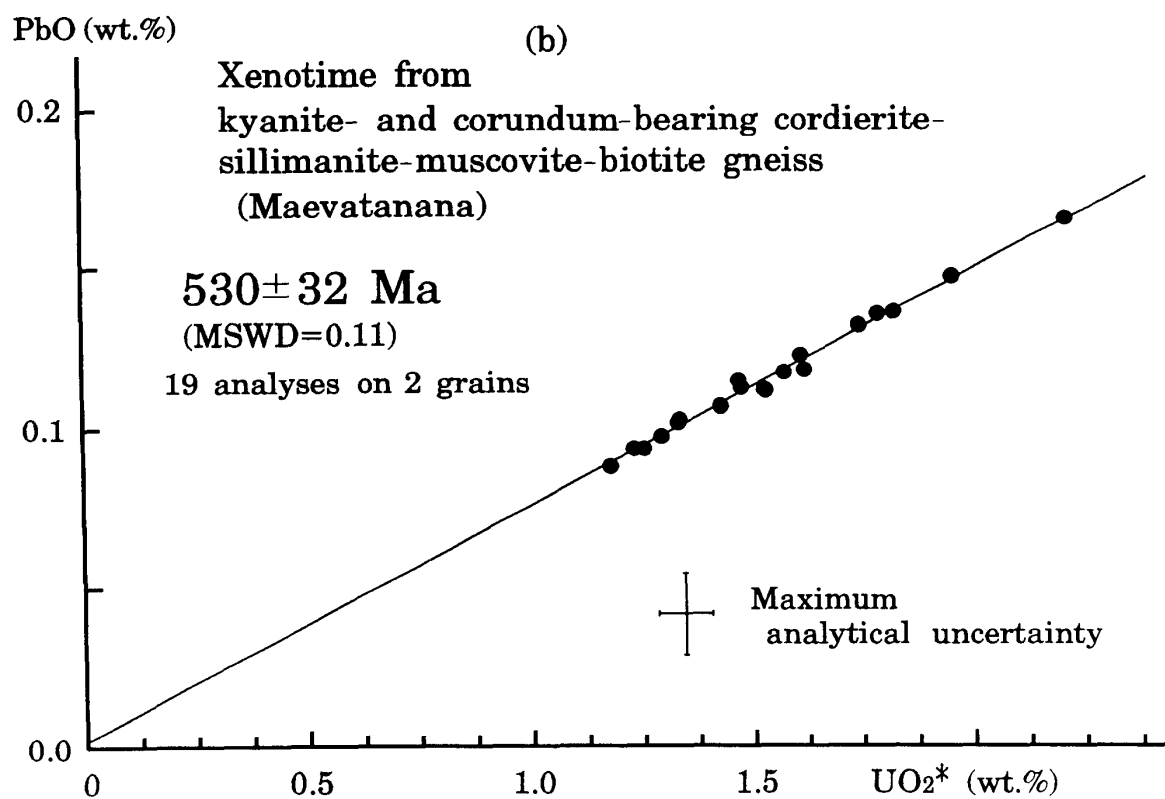
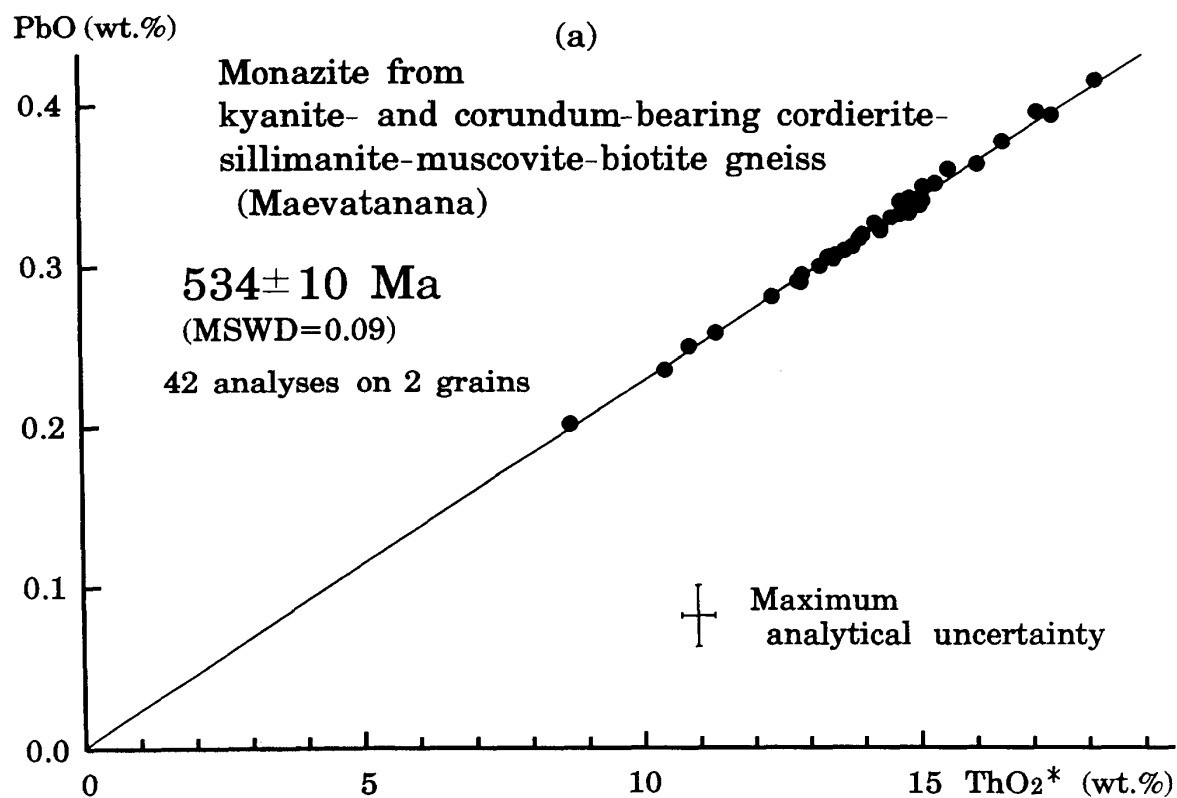


Fig. 4. Plot of PbO vs. ThO<sub>2</sub>\* for monazite (a) and that of PbO vs. UO<sub>2</sub>\* for xenotime (b) from kyanite- and corundum-bearing cordierite-sillimanite-muscovite-biotite gneiss from Maevatanana. Explanation for errors is the same as Fig. 2.

fit regression line yields an age of  $534 \pm 10$  Ma (MSWD = 0.09) and an intercept value of  $0.0010 \pm 0.0062$ . Figure 4b shows the PbO vs.  $UO_2$  plots of 19 analytical data for two xenotime grains. All data points are arrayed linearly, and define an isochron of  $530 \pm 32$  Ma (MSWD = 0.11) with an intercept value of  $0.0019 \pm 0.0072$ . Despite the detailed microprobe analysis, no age signature older than ca. 530 Ma can be obtained on both the monazite and xenotime grains.

## DISCUSSION AND CONCLUSION

The CHIME dating disclosed that the biotite-sillimanite-cordierite gneiss from Ihosy contains chronologically zoned monazite grains as well as unzoned ones. The core ages are  $1640 \pm 180$  and  $797 \pm 75$  Ma, although the mantles and unzoned grains give unequivocal ages of 527–550 Ma. The zoning pattern coupled with the considerable difference in core ages suggests that the older core are of detrital origin, retaining pre-metamorphic information. Inheritance of monazite grains in the uppermost amphibolite to granulite facies paragneisses is not uncommon (Parrish, 1990; Suzuki and Adachi, 1994; Suzuki et al., 1994). Thus, we conclude that the biotite-sillimanite-cordierite gneiss at Ihosy formed through the ca. 530 (527–550) Ma single metamorphic event from a post- $795 \pm 75$  Ma sediment. This reinforces the previous age assignment of late Proterozoic (< ca. 720 Ma) sedimentation and ca. 550 Ma metamorphism for the gneisses exposed at Ihosy (Kröner et al., 1996).

The sample of kyanite- and corundum-bearing cordierite-sillimanite-muscovite-biotite gneiss was collected from the metamorphic sequence peripheral to the Maevatanana greenstone belt of possible Archean age. The nature of the boundary between the peripheral rocks and those of the greenstone proper is still unclear because of poor exposure. Thus, some may regard the metamorphic sequence as the basement unit, and others may consider that the metamorphic sequence overlain unconformably the greenstone belt. This rock gives CHIME monazite age of  $534 \pm 10$  Ma and xenotime age of  $530 \pm 32$  Ma. The well-defined isochrons with no indication of older and younger ages (Fig 4a,b) suggest a single-stage crystallization of monazite and xenotime during the kyanite-sillimanite type metamorphism ca. 530 Ma ago. We, therefore, consider that the paragneisses at Maevatanana represent the northern extension of the ca. 530 Ma metamorphic rocks at Ihosy in the southern sector. Presumably the central NS trending ductile shear zone of Pan-African age is much wider than has been thought (Windley et al., 1994).

The high-grade metamorphism at ca. 530 Ma in Madagascar appears to be essentially synchronous with the granulite facies metamorphism in the south of the so-called Achankovil shear zone in southern India ( $539 \pm 20$  Ma, Santosh et al., 1992;  $558 \pm 11$  Ma, Choudhary et al., 1992;  $527 \pm 10$  Ma, Bindu et al., 1996). Bindu et al. (1996) also reported a  $640 \pm 30$  Ma CHIME age for detrital zircon from the cordierite charnockite sample that gives  $527 \pm 10$  Ma CHIME monazite age. This clearly shows that gneisses in southernmost India are not the

polymetamorphosed Proterozoic complex; they formed newly in the Cambrian from post-640 Ma protolith. Similar Cambrian high-grade paragneisses derived from post-620–760 Ma sediments occur in the Yamato Mountains and the Sør Rondane Mountains in East Antarctica (Shiraishi et al., 1994; Asami et al., 1997).

It has been known that metamorphic rocks in Sri Lanka bear many similarities with those in southernmost India and East Antarctica (Yoshida et al., 1992; Shiraishi et al., 1994). We, therefore, consider that the Cambrian (ca. 530 Ma) metamorphic rocks in Madagascar, southernmost India, Sri Lanka and East Antarctica constituted a single metamorphic terrain that can be linked directly to the collision of East and West Gondwana.

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