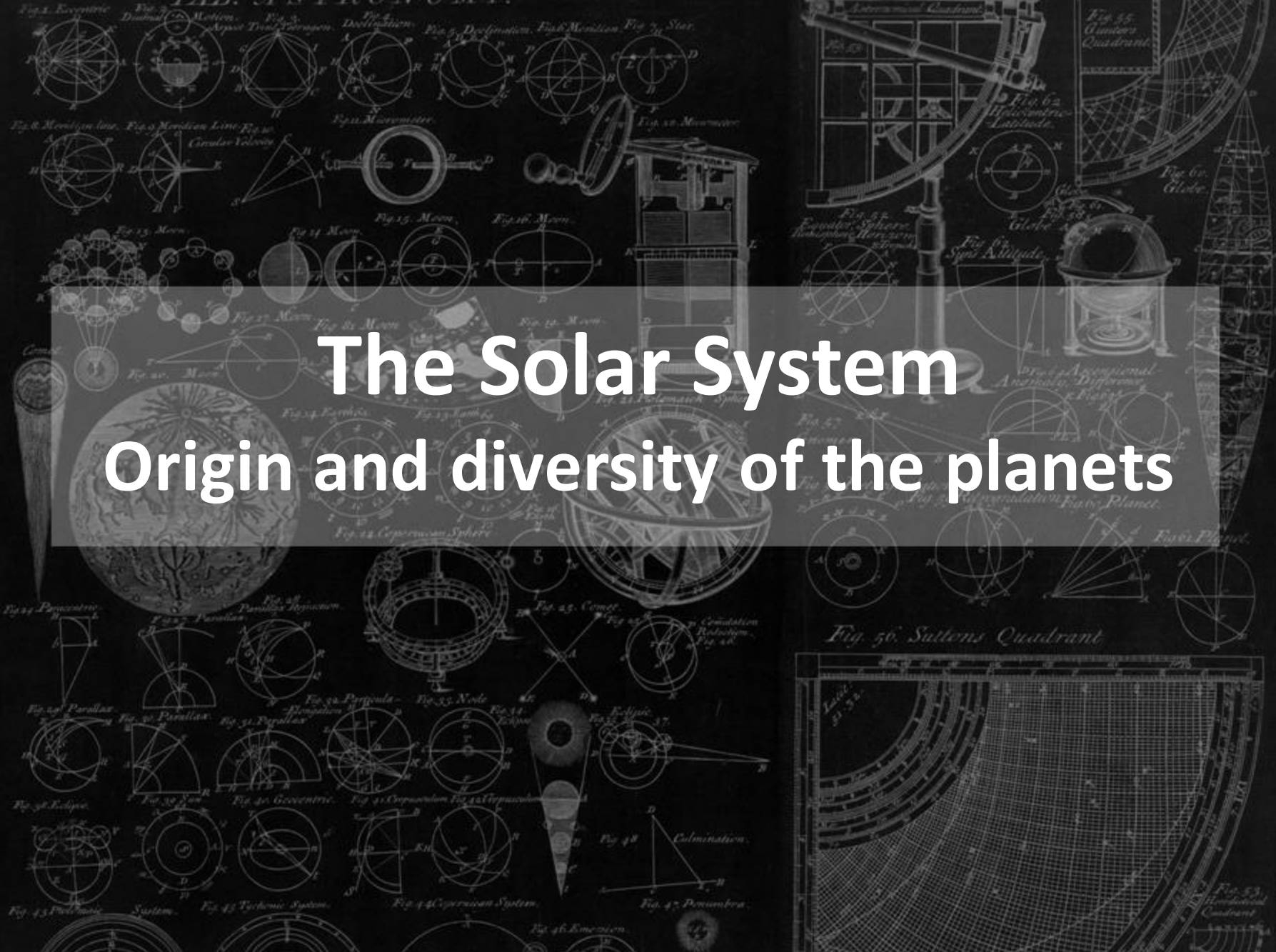


TAB. ASTRONOMY.



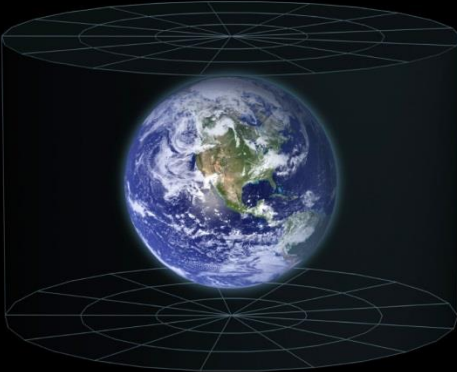
The Solar System

Origin and diversity of the planets

☀ Our place in the Universe

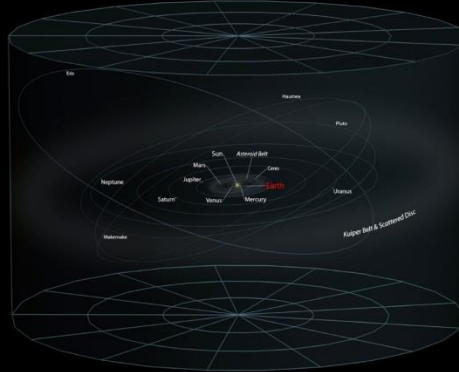
1 light-year = ~10 trillion km (10^{12} km)

EARTH



Earth diam.
 $13 \cdot 10^3$ km

OUR SOLAR SYSTEM



Dist. Earth-Sun
 $148 \cdot 10^6$ km
(17th century)

SOLAR INTERSTELLAR NEIGHBOURHOOD



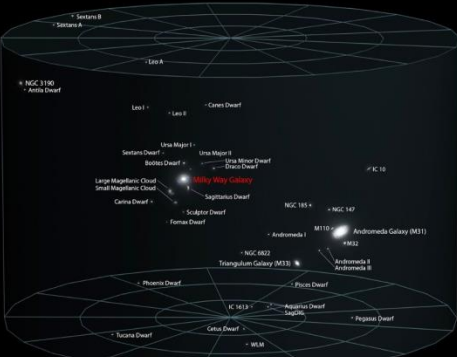
Dist. Earth-closest star
 $40 \cdot 10^{12}$ km
(4 light-years) (19th century)

MILKY WAY GALAXY (~100 billion stars)



Milky Way diam.
 10^{18} km
(10^5 light-years)

LOCAL GALACTIC GROUP

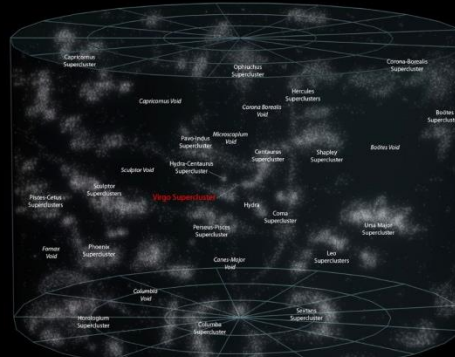


Dist. to nearest galaxy
 $2.5 \cdot 10^6$ l.y.

VIRGO SUPERCLUSTER

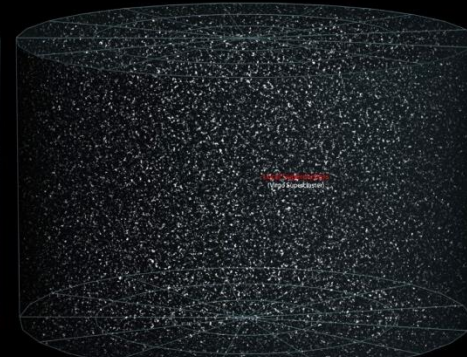


LOCAL SUPERCLUSTERS



Dist. to nearest supercluster
 $300 \cdot 10^6$ l.y.

OBSERVABLE UNIVERSE



Farthest galaxy observed
is $13.2 \cdot 10^9$ l.y. away

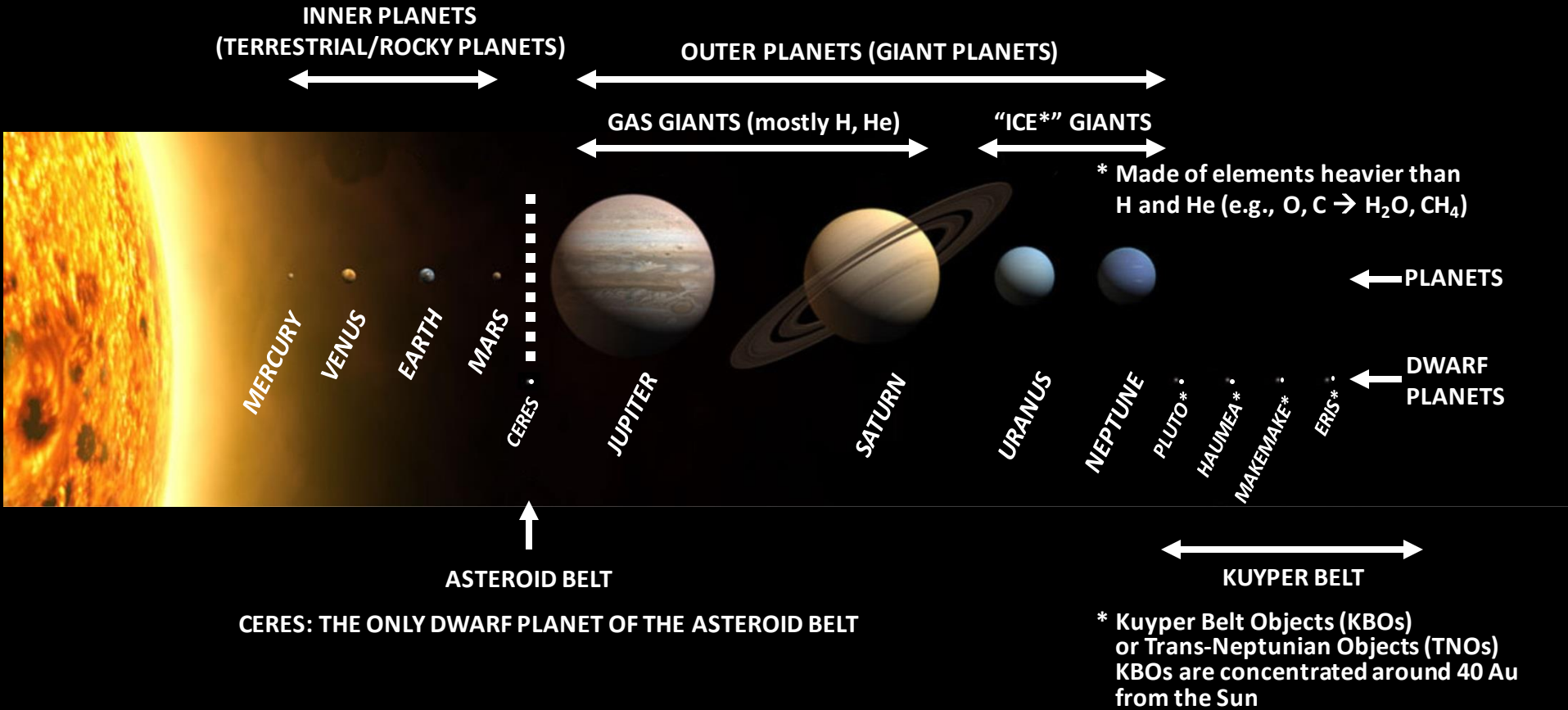
Galaxy Cluster Abell 68
Hubble Space Telescope

ACS/WFC F814W I
WFC3/IR F110W Y
WFC3/IR F160W H

500,000 light-years
153 kiloparsecs 49"



☀ Our Solar System



Age of the Universe:
13.7 billion years

Age of the Solar System
4.56 billion years

Objects of the Asteroid Belt are rich in silicate minerals and metals
Objects of the Kuyper Belt are rich in ices (source of comets)

☼ Planets and dwarf planets

MERCURY



2440 km
(radius)

VENUS



6052 km

EARTH (Moon: 1737 km)



6378 km

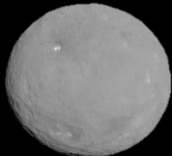
MARS



3388 km

Images by NASA

Ceres



475 km

Pluto



1160 km

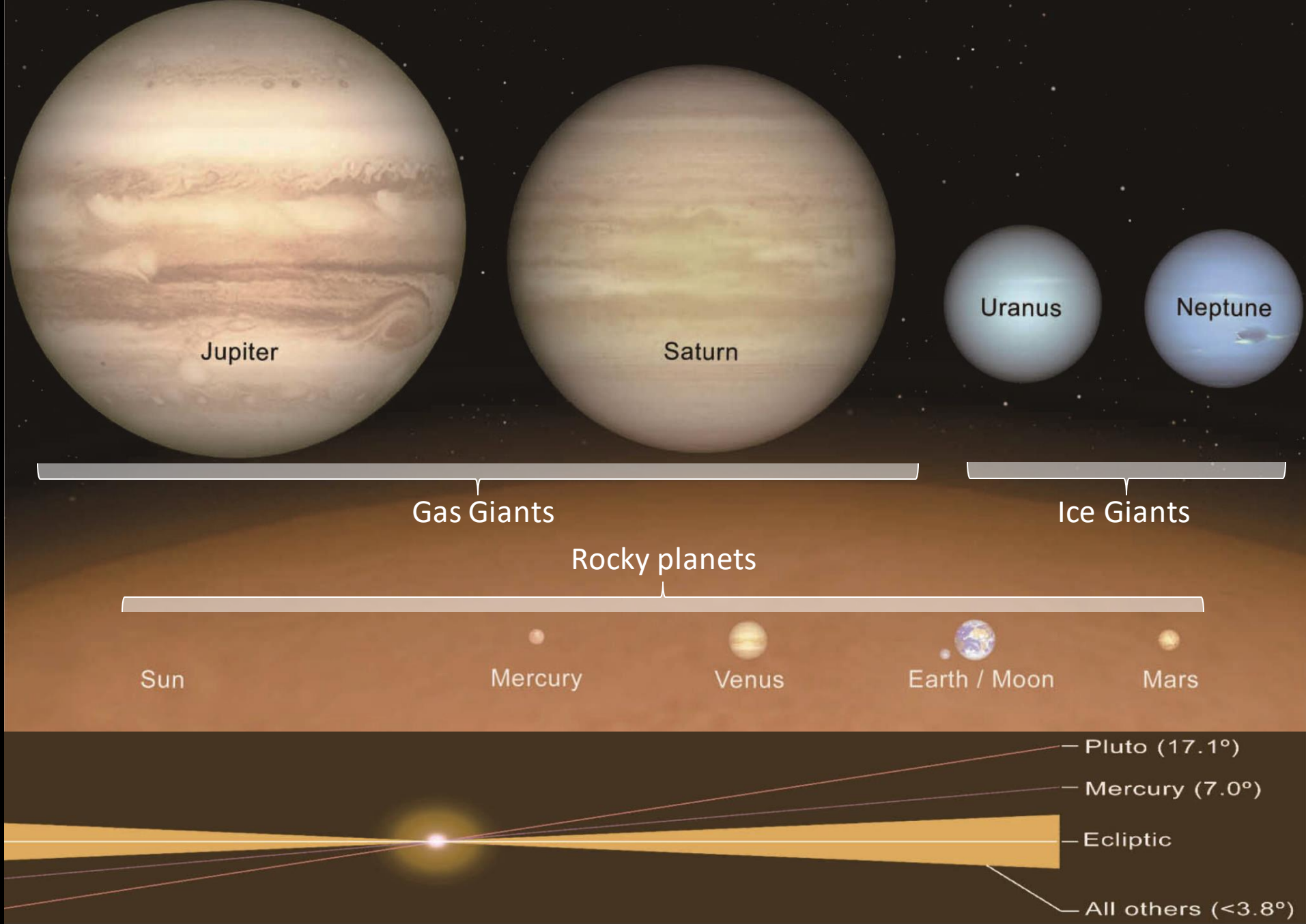
- A dwarf planet is an object:
 - in orbit around the Sun
 - with a rounded shape (gravity –mass– high enough to give it a rounded shape)
 - with a gravitational field too weak to have cleared its orbit; shares its orbit with objects of same size
 - that is not a satellite

NB: The dwarf planet Eris has a radius of 1163 km

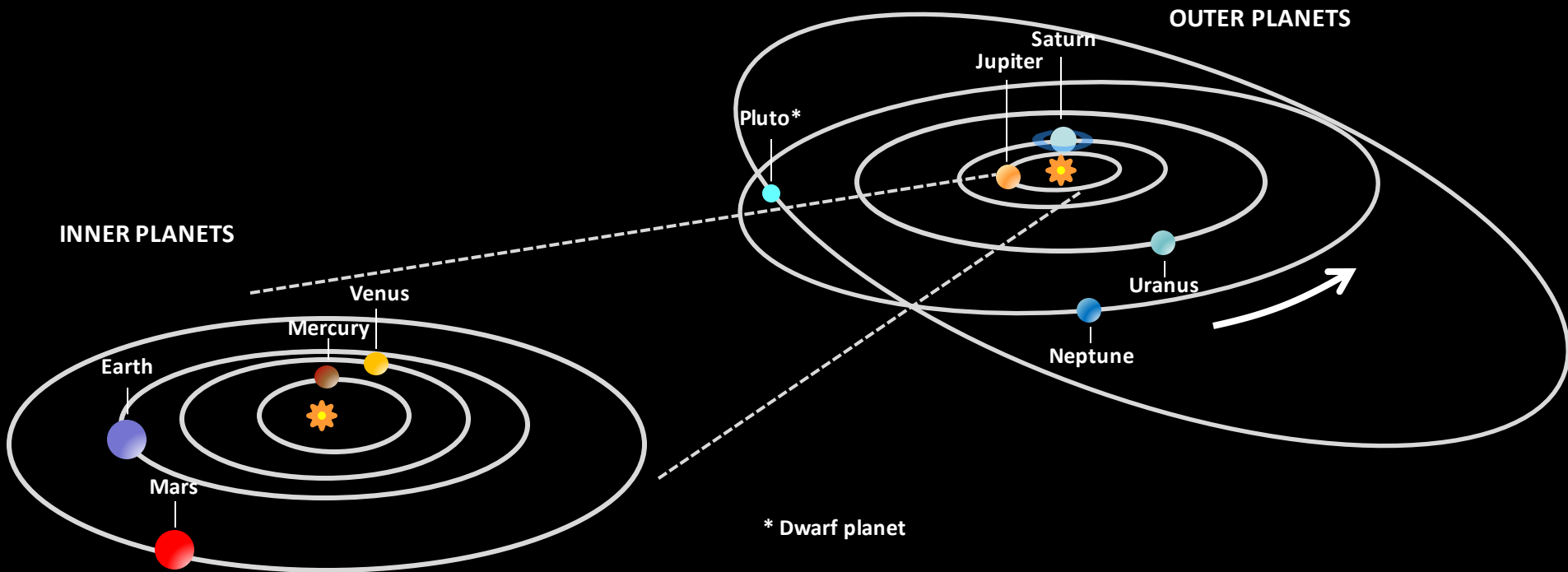
*Resolution B5 of the International
Astronomical Union (2006)*

Asteroids are relatively small
and have an irregular shape
(radius < 200-300 km)





☼ What the model should explain



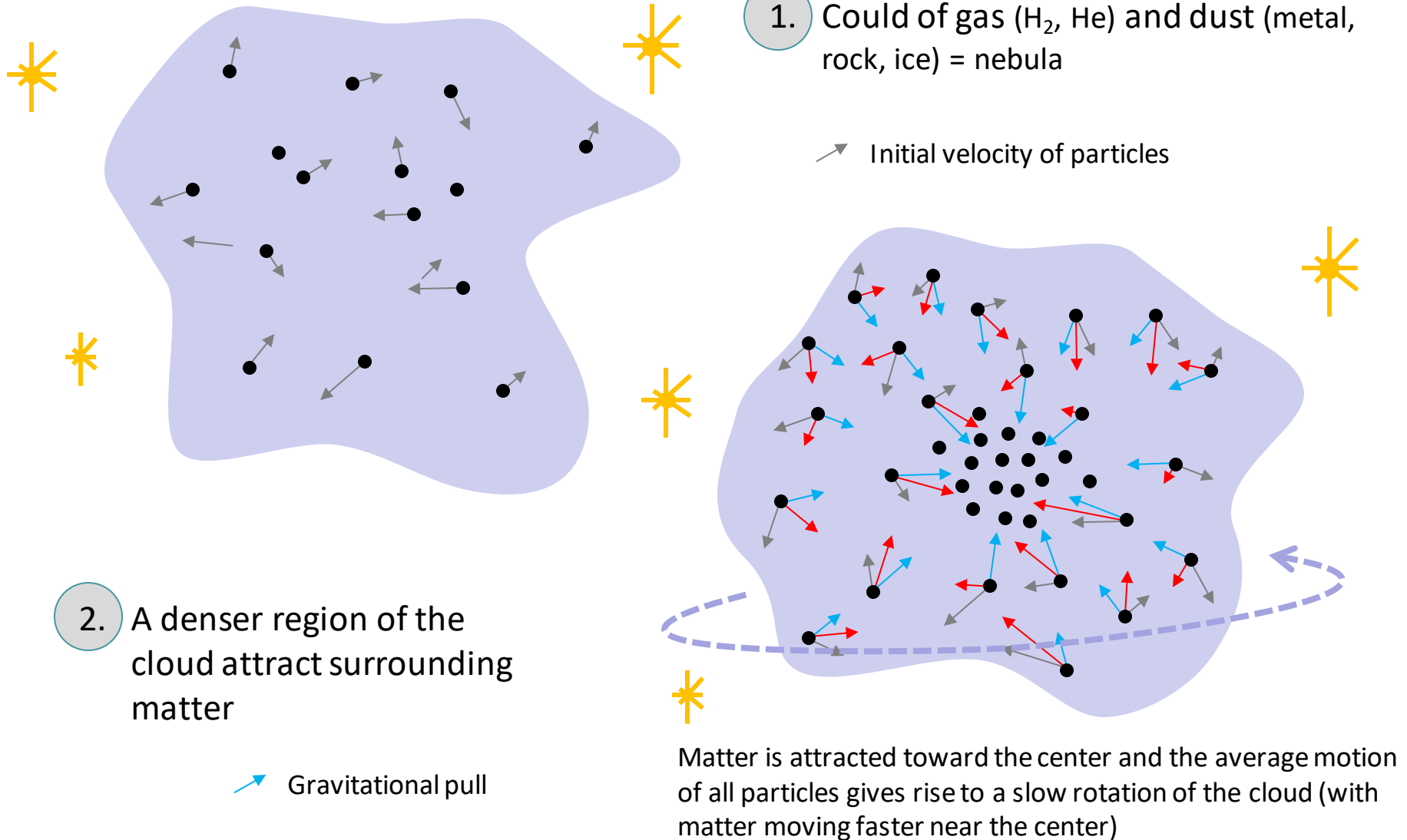
- Planets orbit the Sun in the same prograde direction (prograde = same direction as Sun's rotation)*
- Planetary orbits are nearly circular
- Planetary orbits lie in (or near) the equatorial plane of the Sun
- Inner planets (M, V, E, M) are small, dense, rocky (+ Fe)
- Outer planets (J, S, U, N) are large, made of gases (H, He) and ices
J, S: mostly H, He U, N: melted ices and rock

NB: prograde rotation except Venus (+ Uranus and Pluto with large tilts $> 90^\circ$)

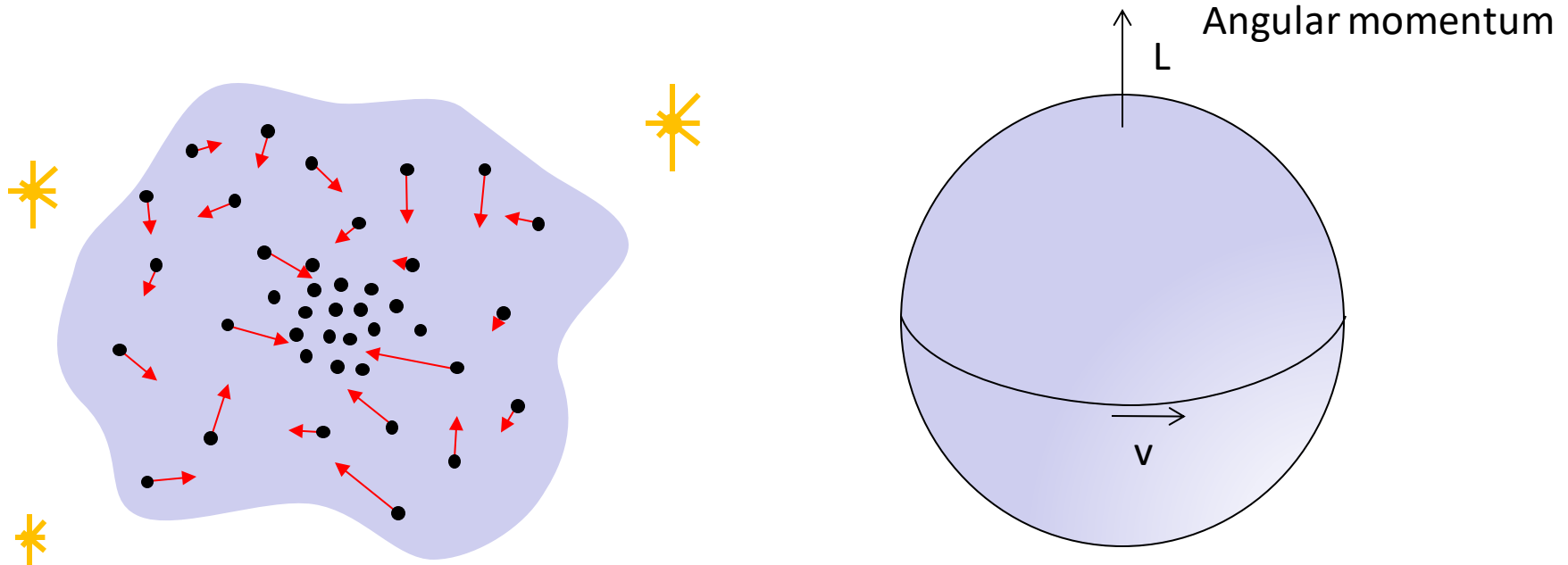
* Counterclockwise (as seen from "above" ecliptic plane)

☀ The nebular hypothesis Immanuel Kant (1724-1804) – Pierre Simon Laplace (1749-1827)

“The Sun and planets formed out of a single collapsing, rotating cloud of interstellar gas and dust, called a solar nebula.” from Lang (2011), Cambridge Guide to the Solar System



Due to the gravitational attraction of the denser region of the nebula, particles are slowly deviated toward that region. They move along a curved path toward the dense region of the cloud (path determined by their initial speed and the force of gravity). There is always more particles spiraling in one direction than in the other (unlikely to have exactly 50-50!). Due to collisions and friction, most particles will tend to move in the same direction after some time (either clockwise or anticlockwise). The resulting motion is therefore a slow rotation of the cloud that accelerates as the cloud collapse on itself due to gravity.



Eagle Nebula

(NASA, ESA, J. Hester, Hubble telescope)

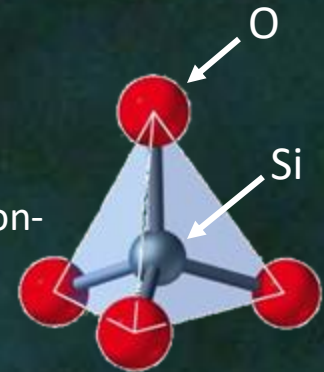
Events can produce variations in density within interstellar clouds:

Proximity of a massive star

Shockwaves of a supernova explosion

Dust particles rich in silicate minerals

Silicates are very common minerals on Earth. Their building blocks consist of silicon-oxygen tetrahedra



NB: One way to see dust clouds is to use telescopes which can detect the infrared radiation that dust particles emit.

Image by NASA
















Mineral	Chemical formula	Cleavage planes and number of cleavage directions	Structure	Specimen
(a) Olivine	$(\text{Mg,Fe})_2\text{SiO}_4$	1 plane 	Isolated tetrahedra 	
(b) Pyroxene	$(\text{Mg,Fe})\text{SiO}_3$	2 planes at 90° 	Single chains 	
(c) Amphibole	$\text{Ca}_2(\text{Mg,Fe})_5\text{Si}_8\text{O}_{22}(\text{OH})_2$	2 planes at 60° and 120° 	Double chains 	
(d) Mica	Muscovite: $\text{KAl}_2(\text{AlSi}_3\text{O}_{10})(\text{OH})_2$	1 plane 	Sheets 	
(e) Feldspar	Orthoclase feldspar: KAlSi_3O_8 Plagioclase feldspar: $(\text{Ca,Na})\text{AlSi}_3\text{O}_8$	2 planes at 90° 	Three-dimensional frameworks 	

Image of the Eagle Nebula taken by Hubble (“pillars of creation”)

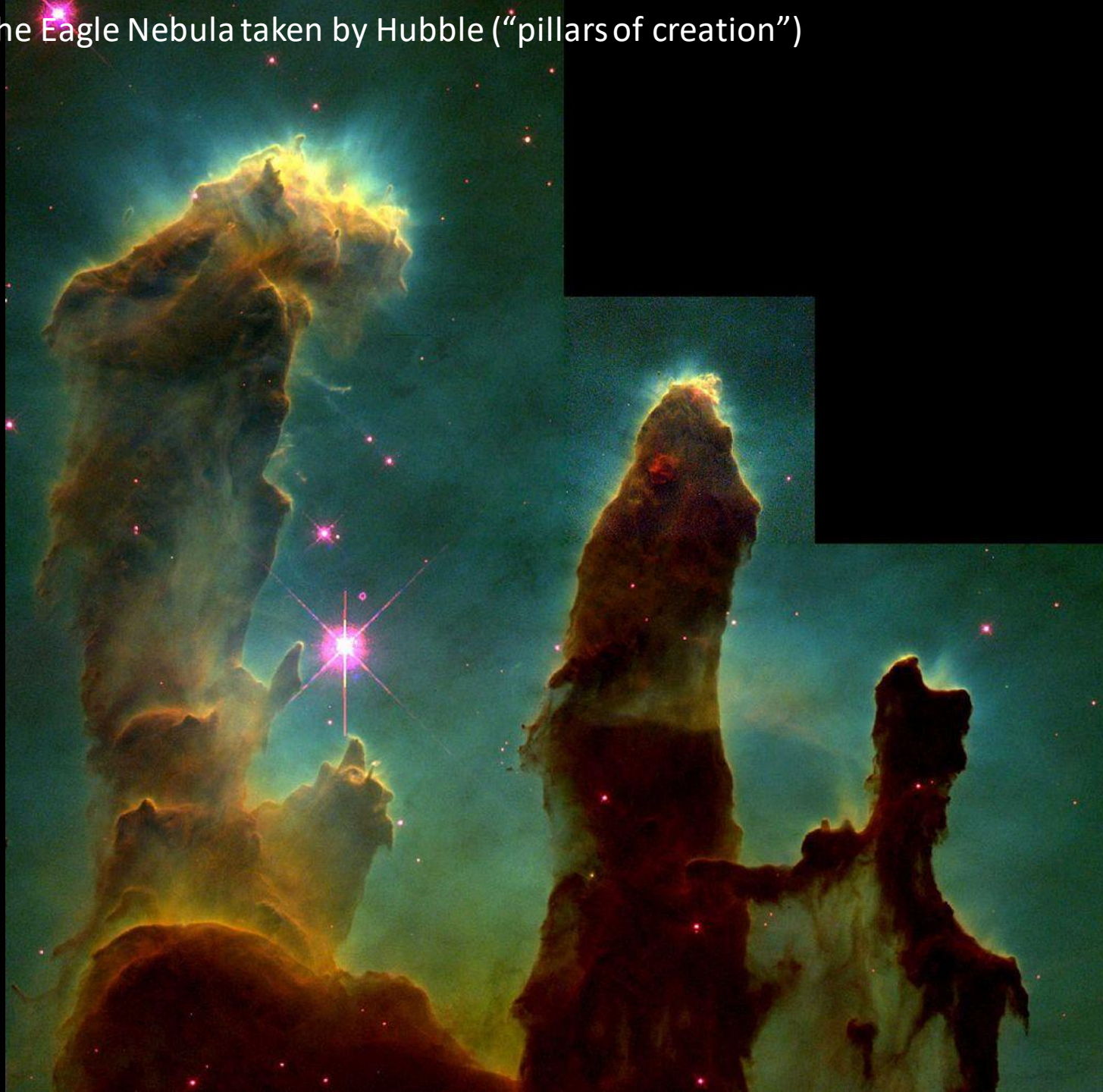
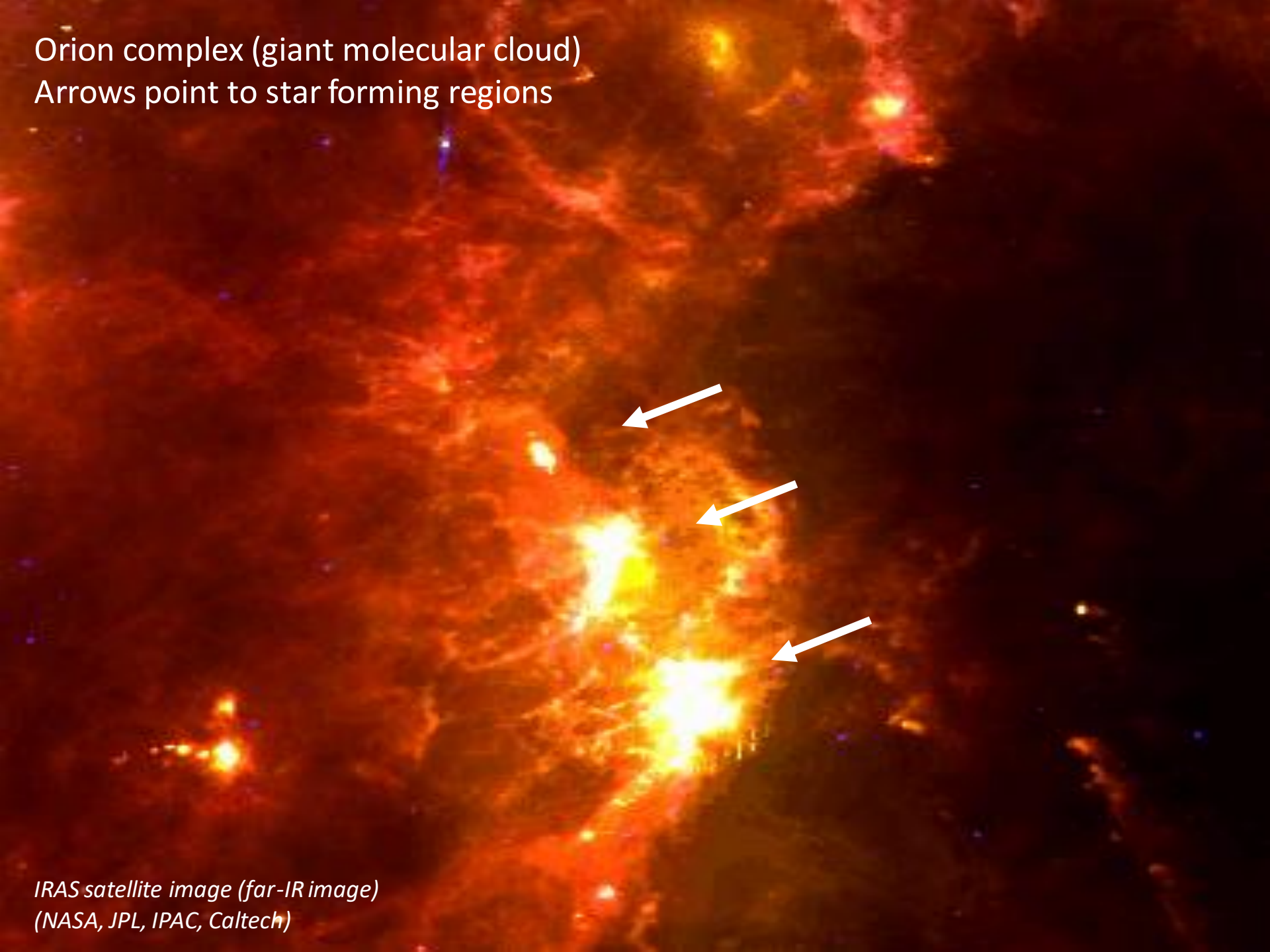


Image by NASA



Omega Nebula
(NASA, ESA, J. Hester, Hubble telescope)

Orion complex (giant molecular cloud)
Arrows point to star forming regions



IRAS satellite image (far-IR image)
(NASA, JPL, IPAC, Caltech)

3. Rotation velocity increases as the cloud contracts under the influence of gravity ("ice skater effect")

Conservation of angular momentum

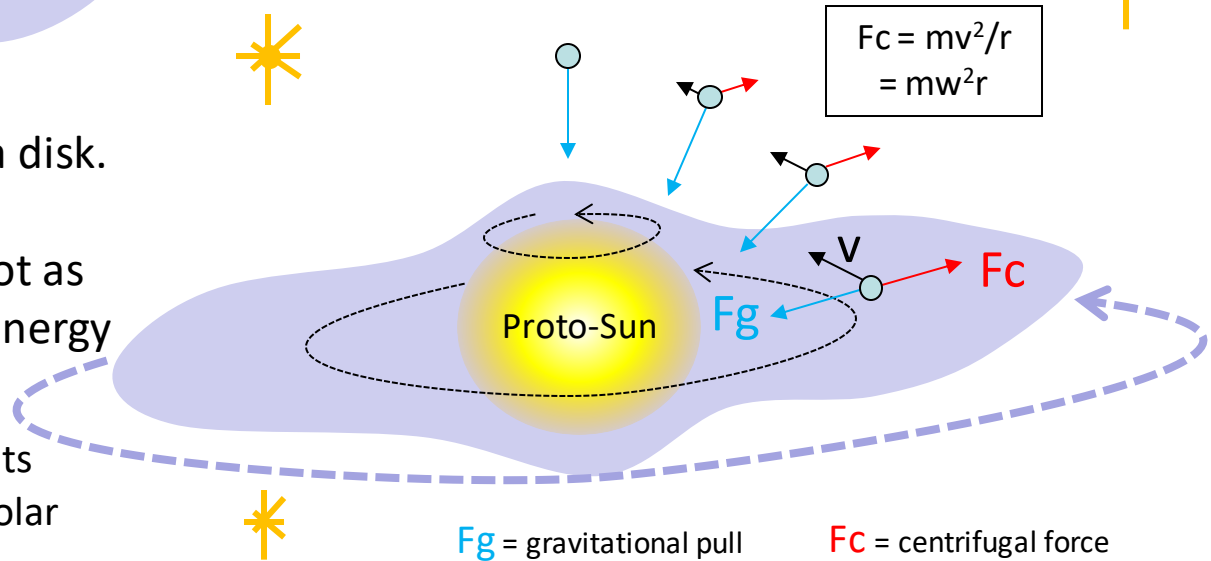
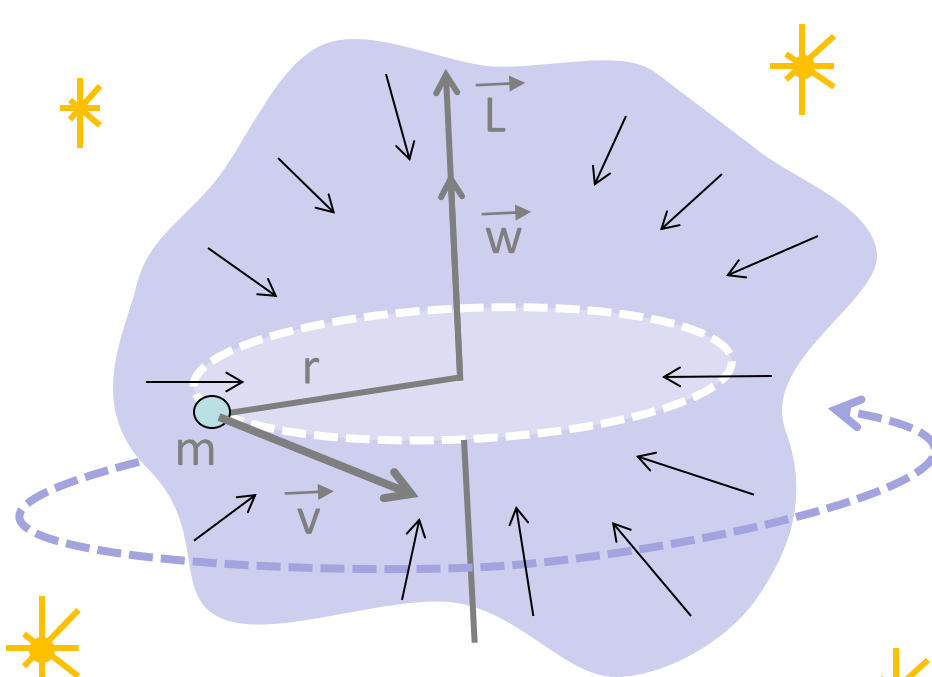
$$L = mvr = mwr^2$$

L = angular momentum
 m = mass of the particle
 r = distance from the center of rotation
 v = linear velocity
 w = angular velocity

4. The cloud flattens into a disk. The center of the disk becomes increasingly hot as gravitational potential energy is converted into heat

Note that the Sun represents 99.9% of the mass of the solar system!

The centrifugal force (inertia) is max in the equatorial plane (linear velocity max.) and opposes the gravitational pull. There is no centrifugal force at the poles (linear velocity = 0), hence "no" force opposes gravity (except gas pressure) and matter falls toward the center of the cloud ("pizza effect").



F_g = gravitational pull F_c = centrifugal force

$$F_c = \frac{mv^2}{r} = mw^2r$$

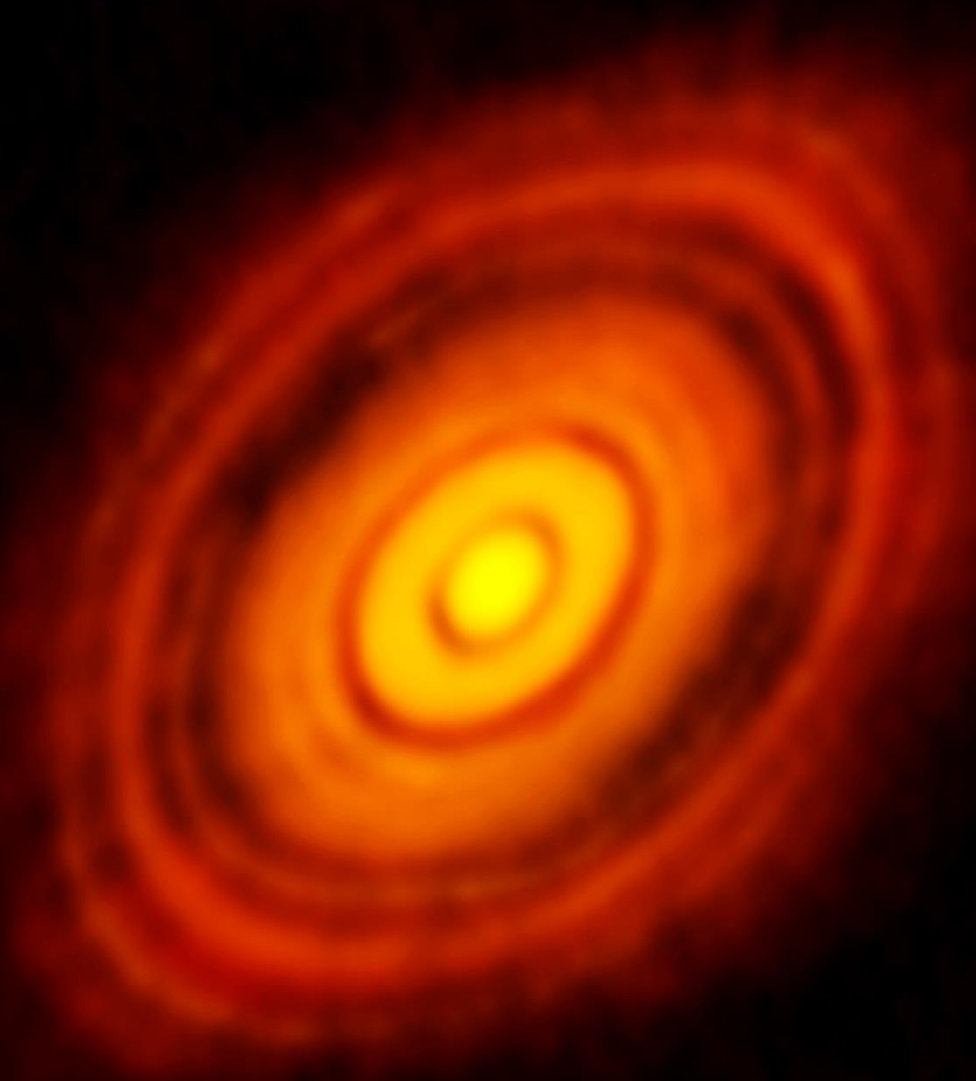
Young stars in the Orion nebula

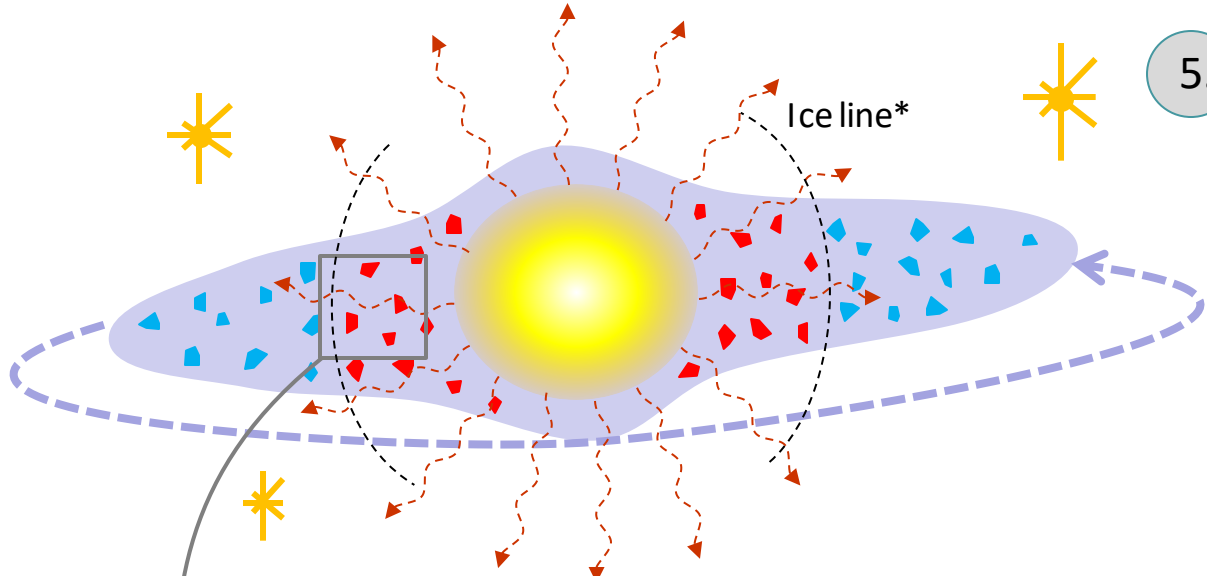


Protostellar disk

“ALMA image of the protoplanetary disc around HL Tauri - This is the sharpest image ever taken by ALMA — sharper than is routinely achieved in visible light with the NASA/ESA Hubble Space Telescope. It shows the protoplanetary disc surrounding the young star HL Tauri. These new ALMA observations reveal substructures within the disc that have never been seen before and even show the possible positions of planets forming in the dark patches within the system.”

Credit: ALMA (ESO/NAOJ/NRAO)
www.eso.org



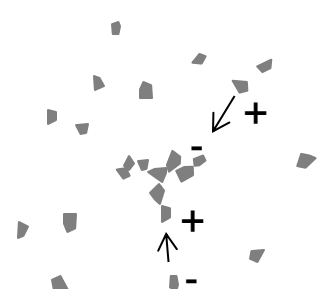


5. Volatile compounds (H_2 , He, H_2O ...) are blown away by the heat and solar wind (stream of particles, mainly protons and electrons). Compounds with a high melting point -refractory- (silicates, iron...) remains in the vicinity of the star.

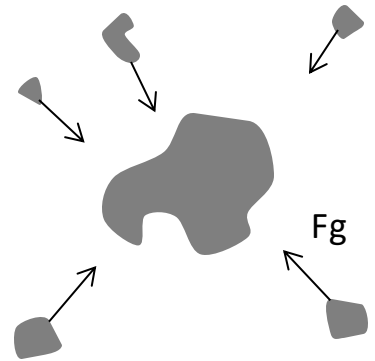
6. In the disk, matter coalesces into increasingly larger objects.

* Ice line \rightarrow beyond this line, the temperature is low enough to enable volatile compounds, such as water (H_2O), ammonia (NH_3), methane (CH_4), to condense into ice. The giant planets formed beyond this line.

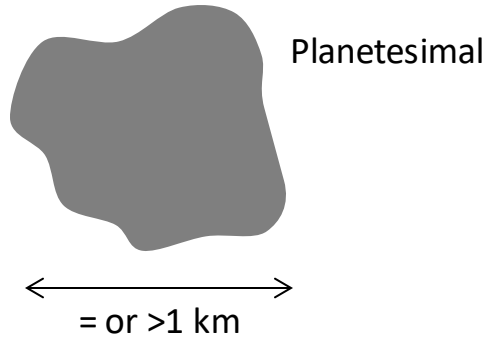
First, small grains are attracted by electrostatic force.



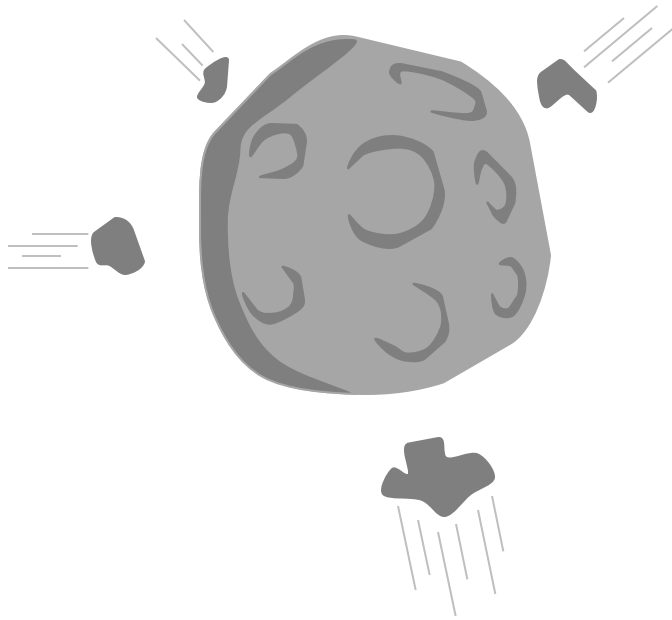
Then larger pieces begin to attract nearby objects by gravity



Gravity pulls together objects of increasing size.

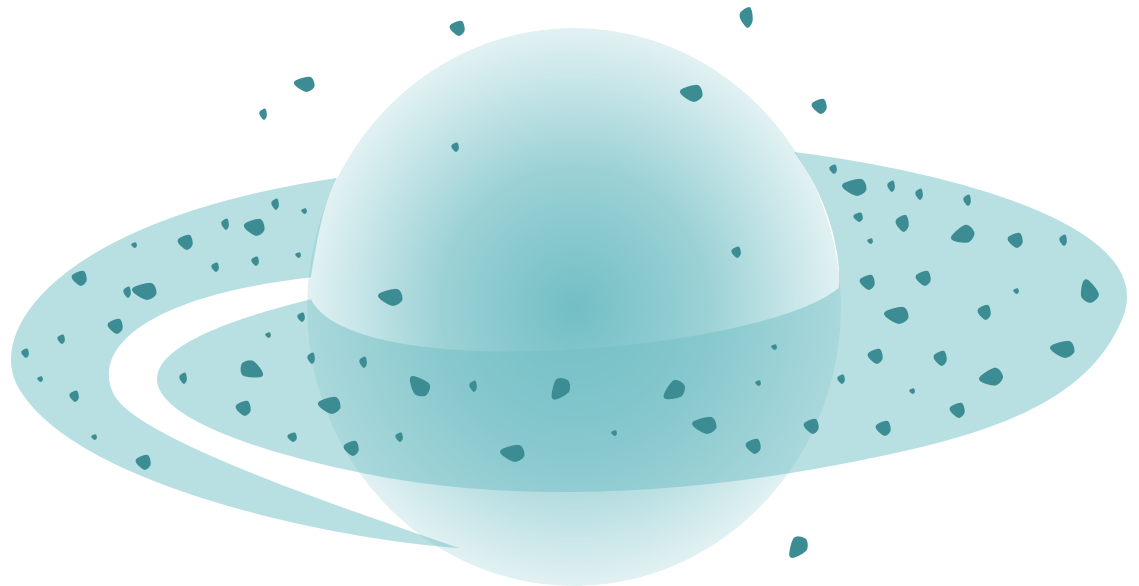


Closer to the proto-Sun:



Rocky material (e.g. silicates) and metals (e.g. Fe, Ni) coalesce to form the inner planets

Farther from the proto-Sun (beyond the ice line):



Gaseous (H_2 , He) and icy compounds (H_2O , NH_3 , CH_4)* coalesce to form the outer planets

* Together with some rocky material still present in the outer solar system

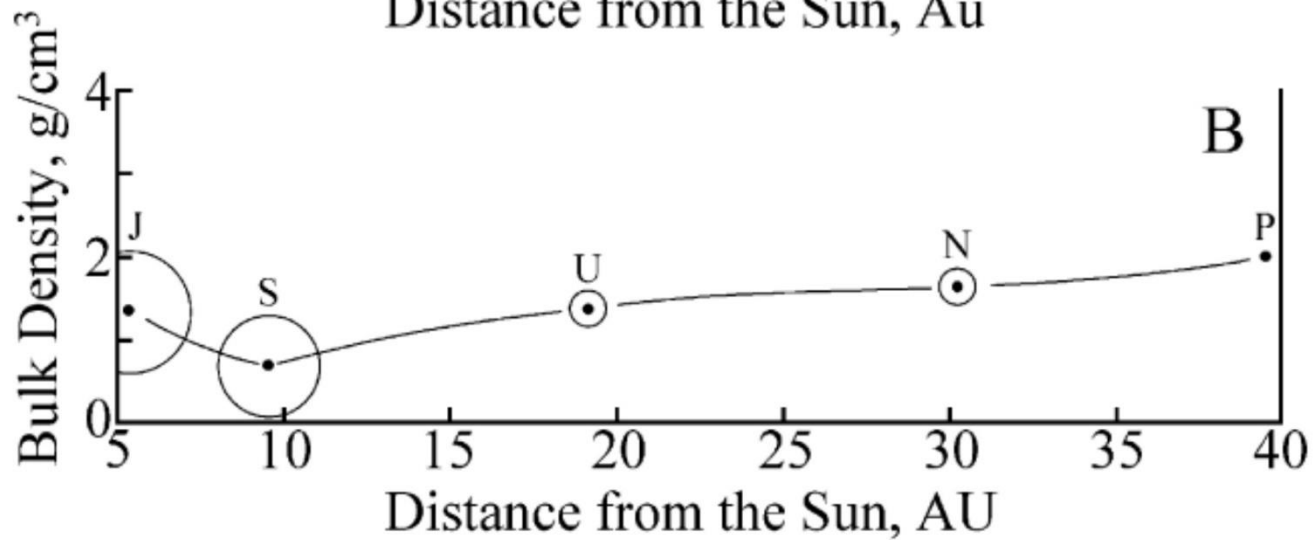
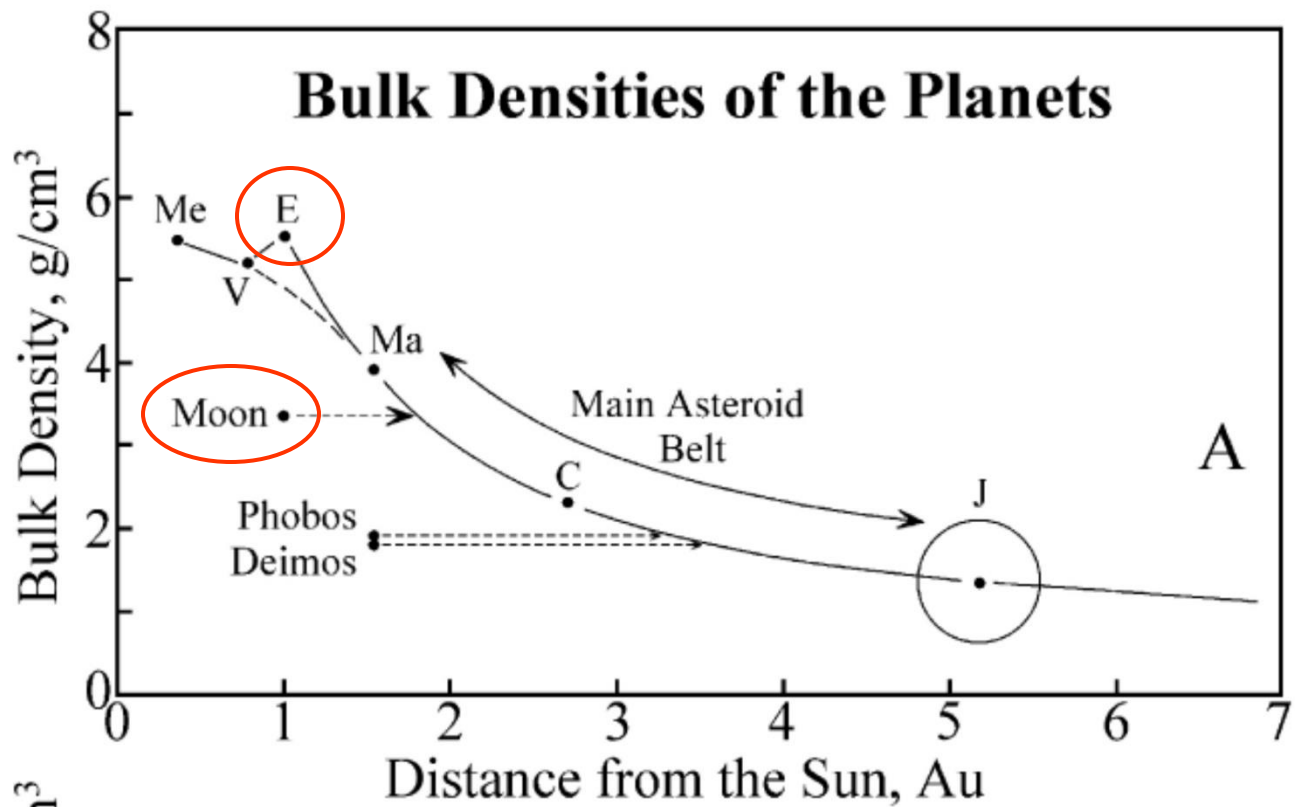
Ca-Al-rich inclusions found
in some meteorites →
Oldest age measured in the
solar system! → 4.567 Ga

■ Table 29.2 Condensation Sequence in the Solar Nebula

Temperature (K)	Condensations and Reactions Occurring
1600	Condensation of refractory oxides such as CaO, Al ₂ O ₃ , TiO ₂ , and rare-earth oxides
1300	Condensation of metallic nickel-iron alloy
1200	Condensation of enstatite (MgSiO ₃)
1200 – 490	Progressive oxidation of remaining metallic iron to FeO, which in turn reacts with enstatite to make olivine [(Fe, Mg) ₂ SiO ₄]
1000	Reaction of sodium with Al ₂ O ₃ and silicates to make feldspar and related minerals; condensation of potassium and other alkali metals
680	Reaction of H ₂ S with metallic iron to make troilite (FeS)
550	Combination of water vapor (H ₂ O) with calcium-bearing minerals to make tremolite
Ice line 425	Combination of water vapor with olivine to make serpentine
175	Condensation of water ice
150	Reaction of ammonia gas (NH ₃) with water ice to make the solid hydrate NH ₃ ·H ₂ O
120	Partial reaction of methane gas (CH ₄) with water ice to make the solid hydrate CH ₄ ·7H ₂ O
65	Condensation of argon and leftover methane gas into solid argon and methane
<25	Condensation of neon, hydrogen, and helium (temperature probably never fell this low, so this step did not occur in solar nebula)

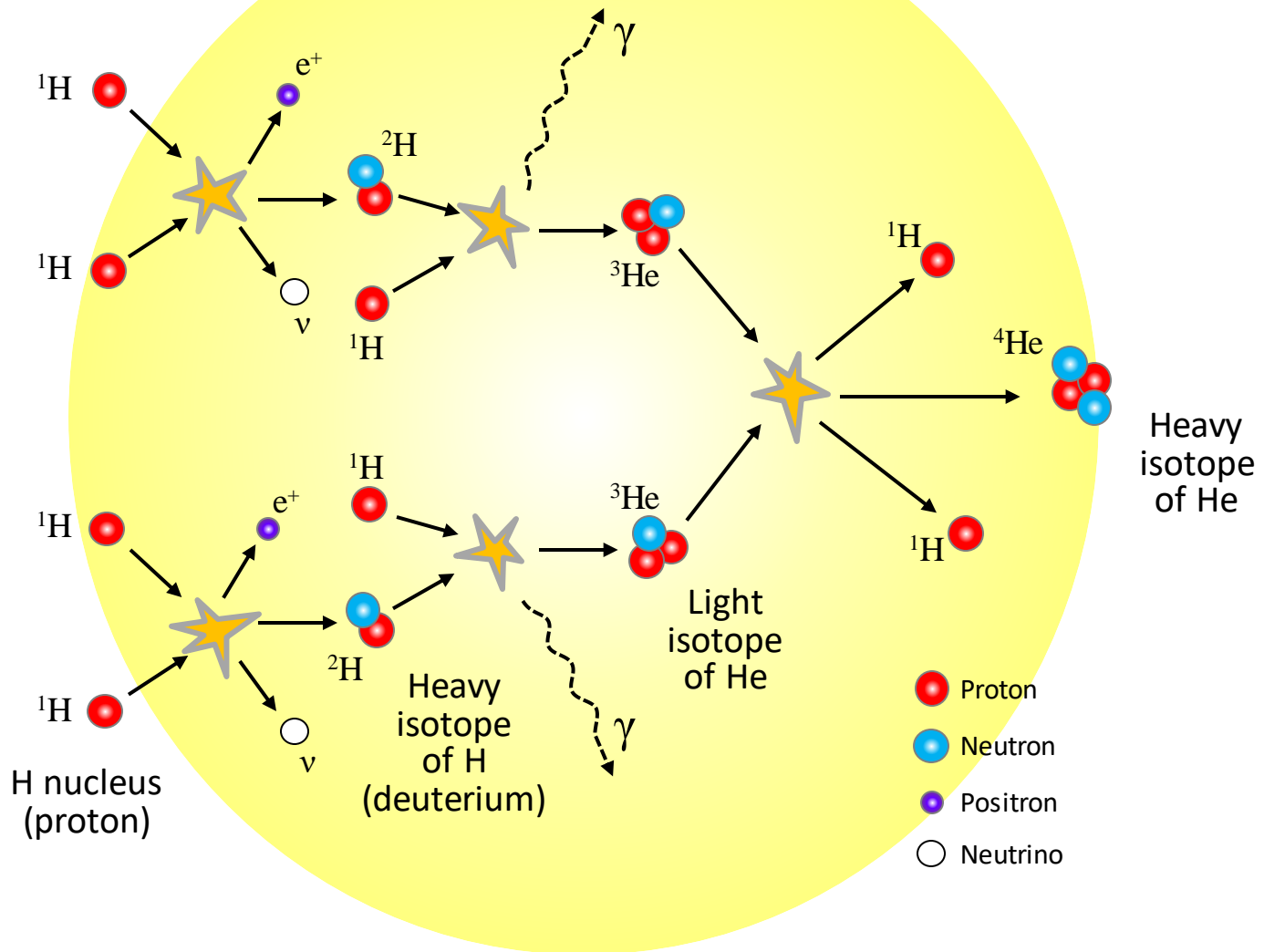
Modified from Lewis, John S. "The Chemistry of the Solar System." In *Scientific American*, vol. 230 (3), 1974.

$$0^{\circ}\text{K (kelvin)} = -273^{\circ}\text{C}$$

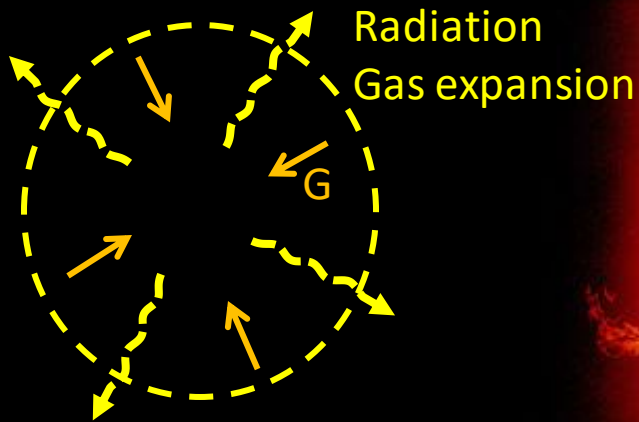


The center of the proto-Sun is so hot that hydrogen atoms are stripped of their electrons. The “soup” of electrons and protons is called a plasma. When temperature reaches 12×10^6 °C, nuclei fuse in a chain reaction producing a tremendous amount of energy (source of solar energy) = **NUCLEAR FUSION** → A star is born: the Sun

Proton-proton chain:

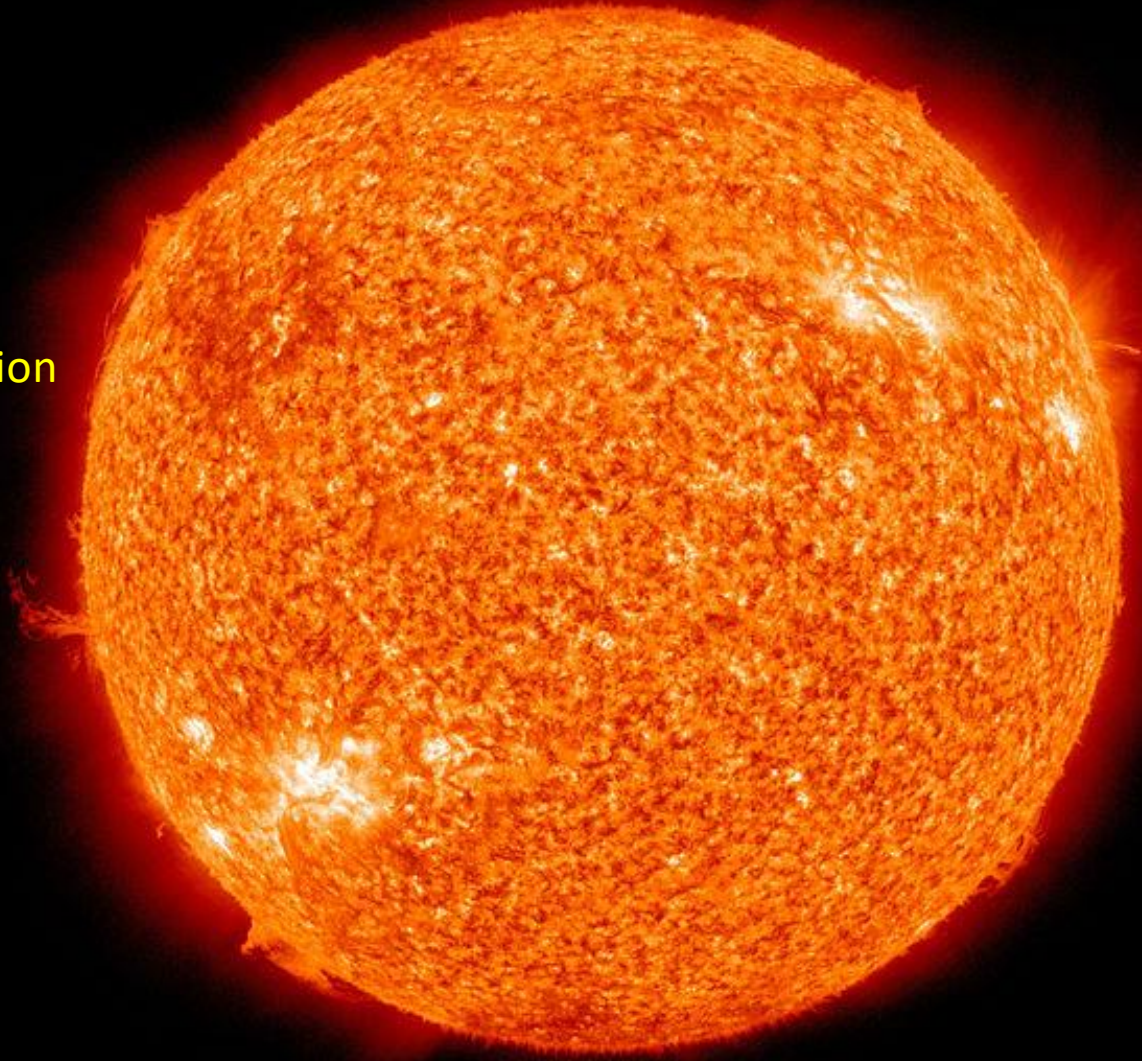


A star obtains the status of "main sequence star" (like our Sun) once the outward forces (radiation + gas expansion) balance the inward gravitational pull.



After all the hydrogen inside the star is used up, the star collapses and the hydrogen present in the outer layer of the star starts to fuse and the outer layer expands. The star becomes a red giant (100x Sun's current radius).

Our Sun has enough fuel for another 5 billion years.



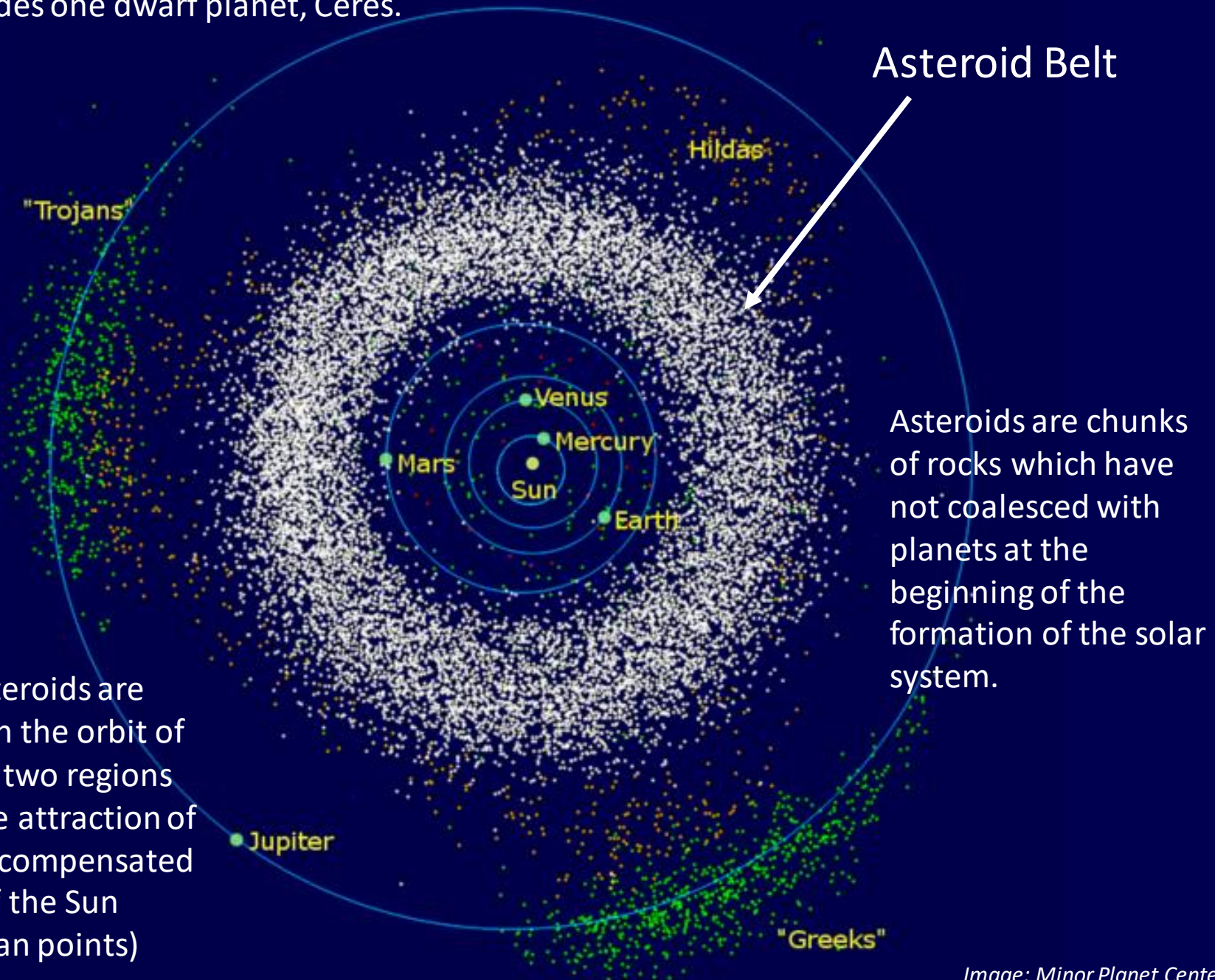
THE SUN

- Giant Molecular Cloud (nebula) • 0
- Protosun (central region warming up to 10,000 °C) • 2.03×10^6 yrs
- Solar nebula flattened into a disk • 2.13×10^6 yrs
- Main sequence star (Sun) • $30-50 \times 10^6$ yrs

THE PLANETS

- Planetesimals/protoplanets • 2.2×10^6 yrs
- Gas Giants, asteroids, comets • $2-10 \times 10^6$ yrs
- Terrestrial planets • $10-100 \times 10^6$ yrs
- Heavy Bombardment • $100-1300 \times 10^6$ yrs

The Asteroid Belt, between the orbit of Mars and Jupiter, is composed of billions of asteroids and includes one dwarf planet, Ceres.



Asteroid Belt

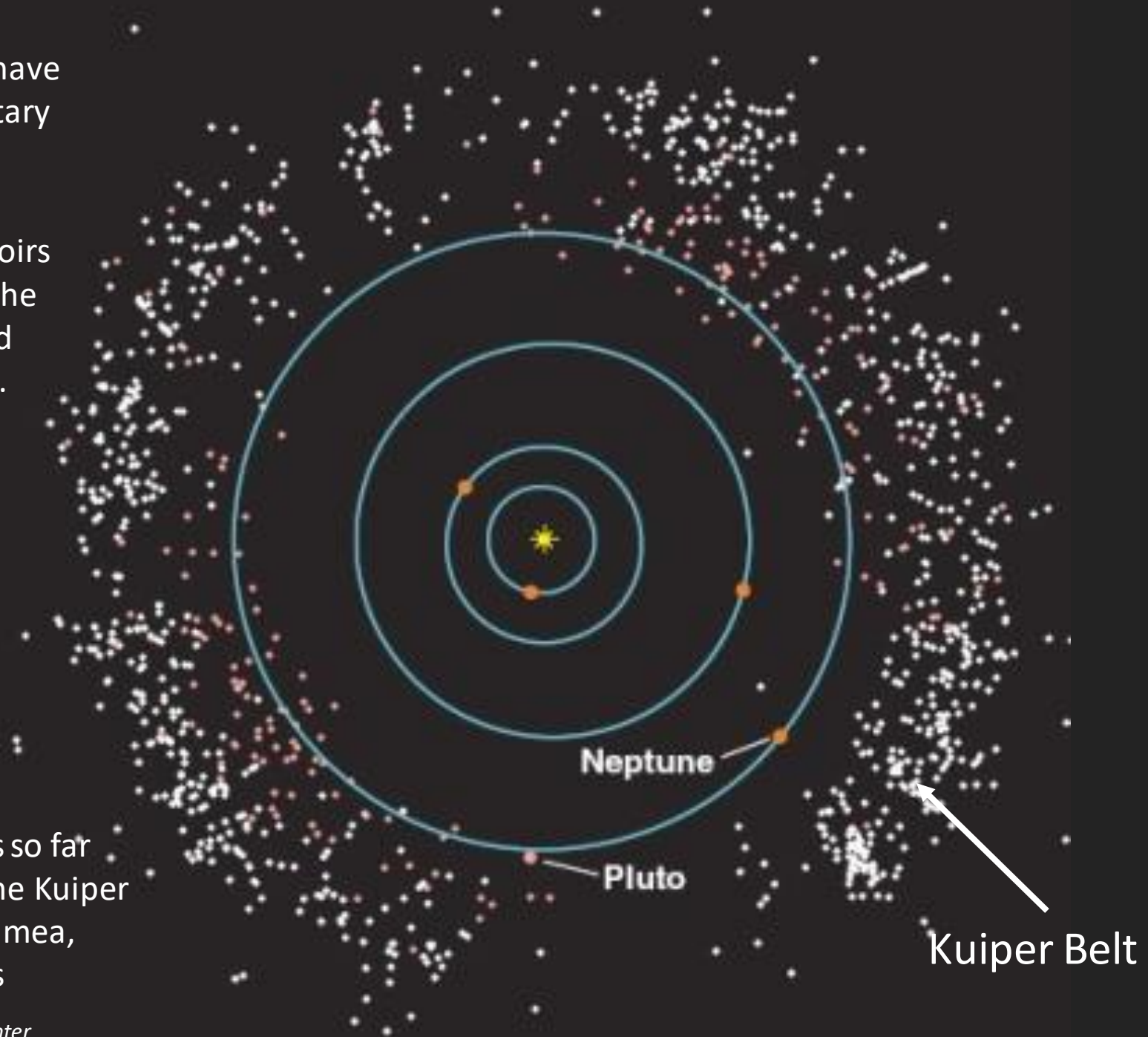
Asteroids are chunks of rocks which have not coalesced with planets at the beginning of the formation of the solar system.

Trojan asteroids are located on the orbit of Jupiter in two regions where the attraction of Jupiter is compensated by that of the Sun (Lagrangian points)

Comets are icy objects which have escaped planetary accretion (like asteroids).

The two reservoirs of comets are the Kuiper Belt and the Oort Cloud.

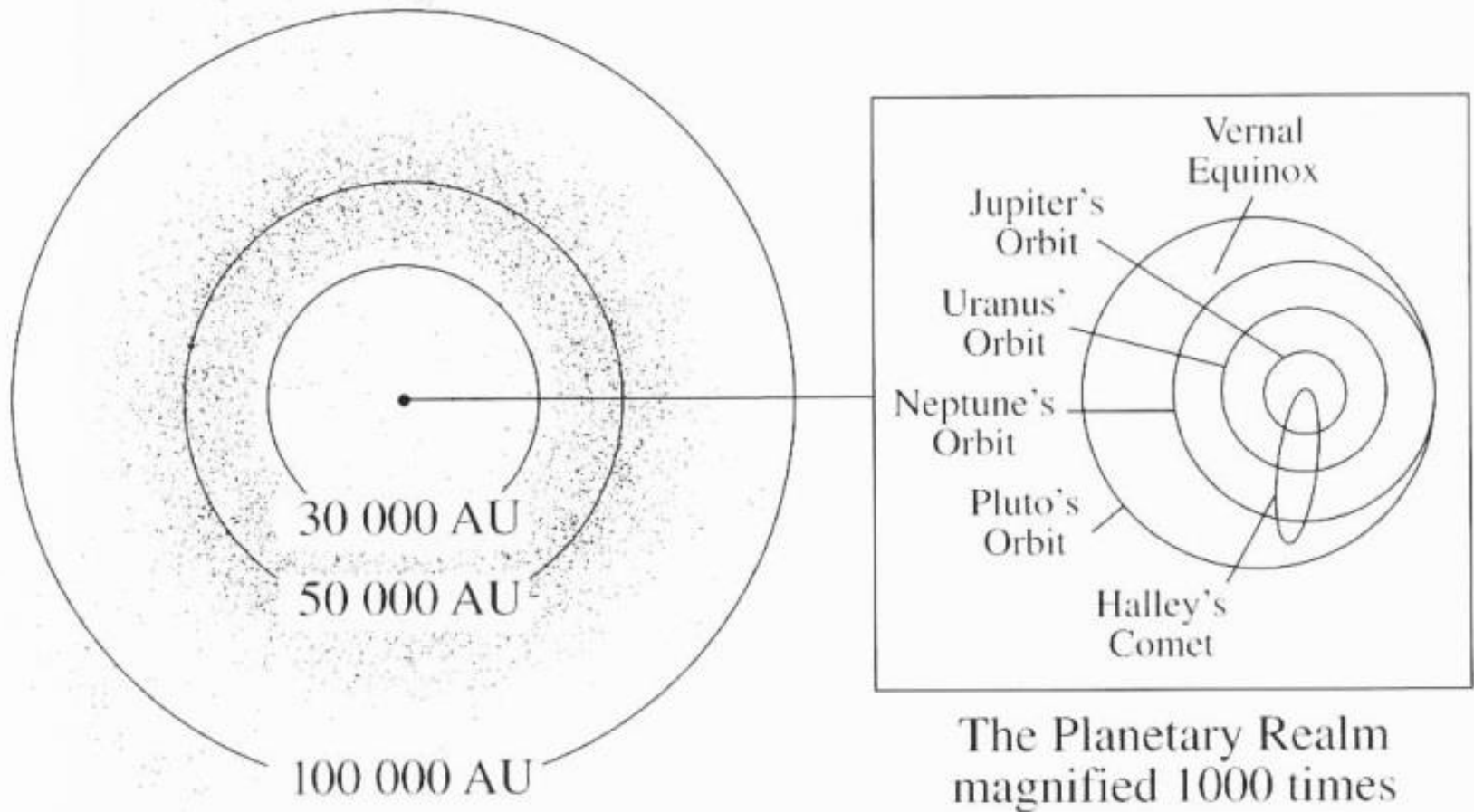
4 dwarf planets so far discovered in the Kuiper belt: Pluto, Haumea, Makemake, Eris



Kuiper Belt

Oort Cloud

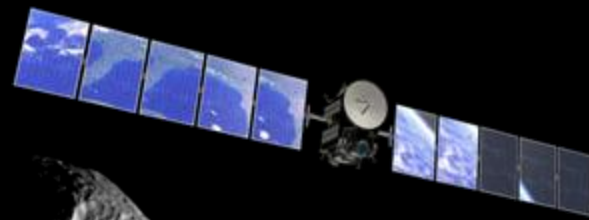
By studying the size and trajectory of comets, the Dutch astronomer Jan H. Oort (1900-1992) has postulated the existence of a "comet reservoir" surrounding the solar system (way beyond the Kuyper Belt).



Asteroids and comets give us information about the raw material that accreted to form the planets.

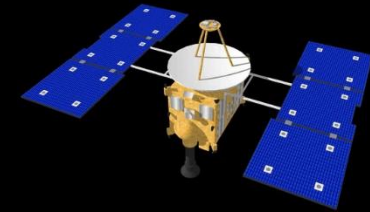
Most meteorites falling on Earth are pieces ejected during collisions between asteroids. Hence, their age can give us an idea of the age of the Solar System. The currently accepted age of the Earth is 4.56 Ga based on the age of meteorites.

Computer models suggest that it may have taken only 10 million years for the planets to form.



NASA **DAWN** MISSION

Exploration of Vesta and Ceres (Asteroid Belt, launched 2007, explored Vesta & Ceres)



JAXA **HAYABUSA I* & II**** MISSION

*Mapping and sampling of asteroid Itokawa & Ryugu (*launched 2003, returned 2010; **launched 2014, reached target 2018)*

ESA **ROSETTA** MISSION

Mapping and landing probe on comet Chury (launched 2004, reached target 2014)



NASA **OSIRIS-Rex** MISSION

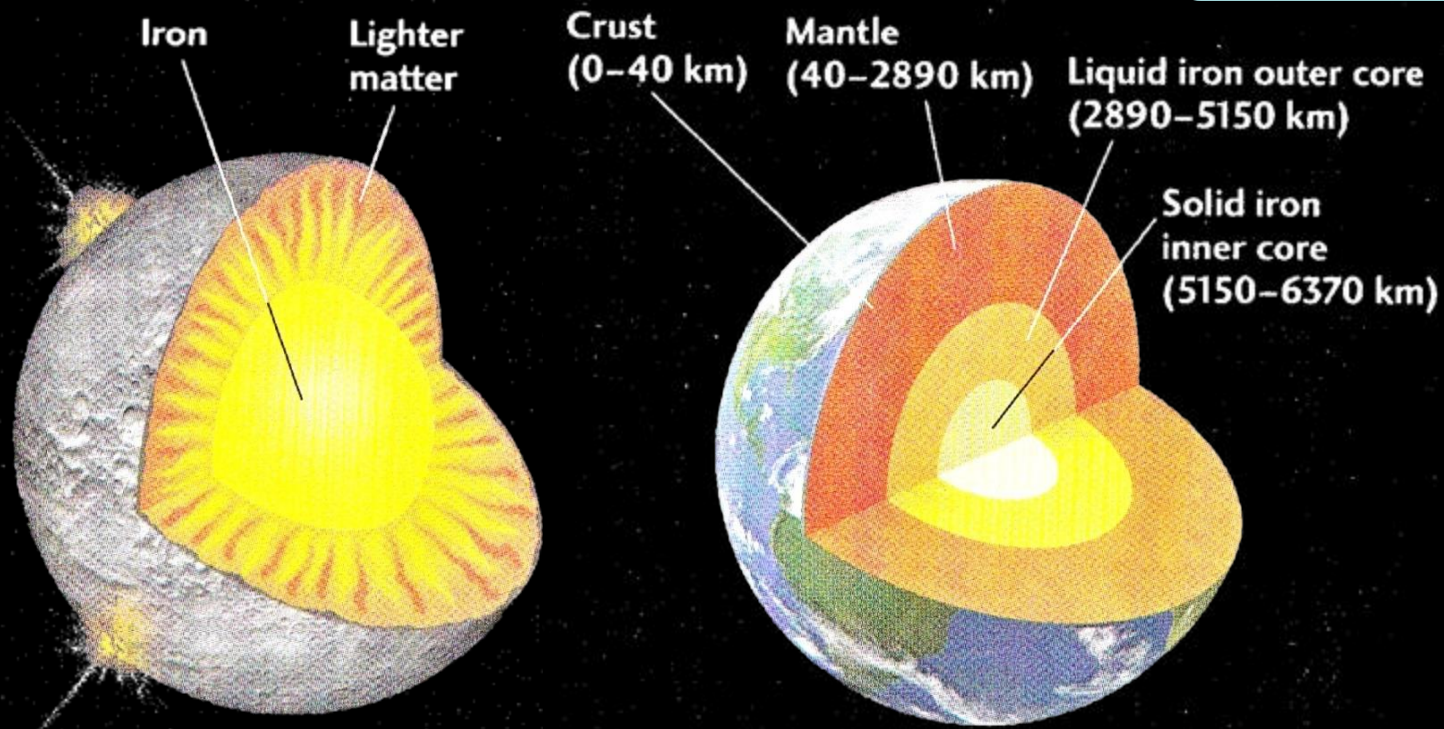
Mapping and sampling of carbon-rich asteroid Bennu (launched in 2016, sampling-return 2020-2021)



☀ Formation of Earth's layers

- Earth was in a molten state at the beginning of its history.
 - Impacts with meteorites and planetesimals (conversion of kinetic energy)
 - Radioactive decay
 - Gravitational contraction (conversion of gravitational potential energy)
- Matter redistributed according to differences in density:

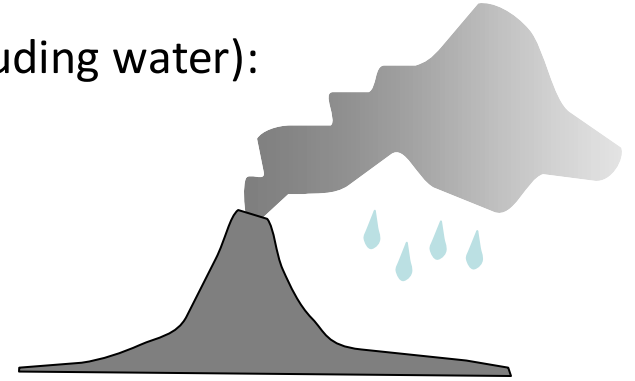
Gravitational differentiation



☀ Origin of Earth's atmosphere and water

■ Two possible sources for volatile components (including water):

- Volatiles trapped in the **Earth's interior** and released during volcanic activity



- Volatiles from **meteorites** and (especially) **comets**.



■ The composition of the primitive atmosphere was different from today:

→ H_2 , N_2 , CO_2 , and H_2O . **NO OXYGEN!**

- Oxygen was produced later by biological activity (**photosynthesis**)!
- Once temperature began to fall, water vapor condensed in a liquid state and primitive oceans began to form (probably less than 200 Ma after the beginning of Earth's formation).

- The Moon



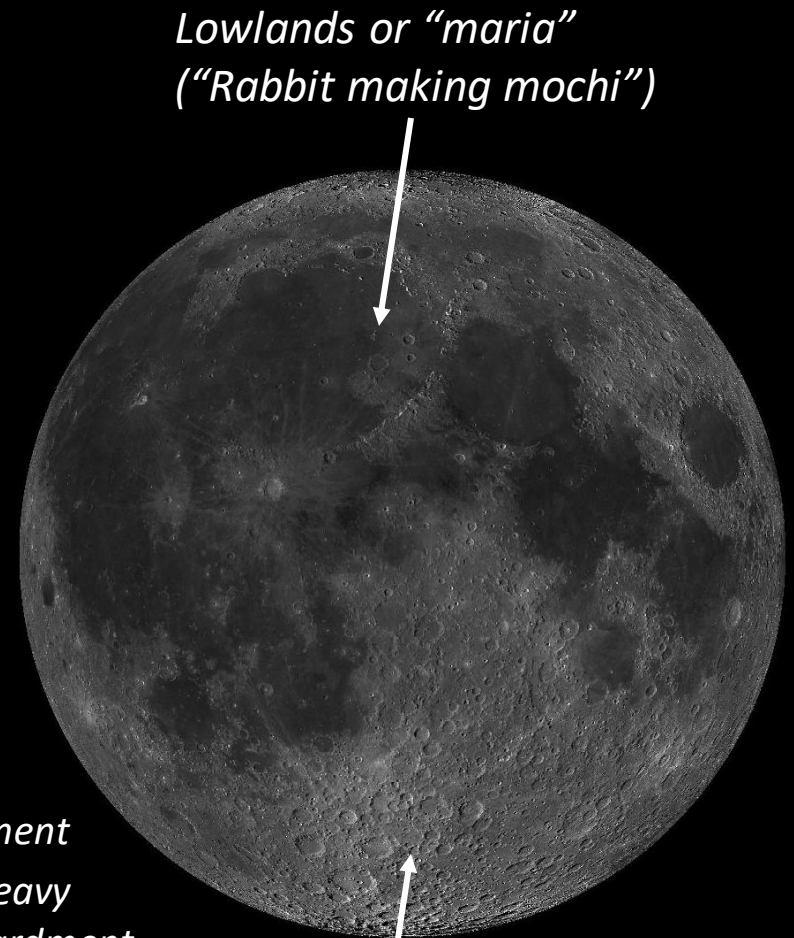
In the case of the Moon, we have access to the absolute age of surfaces with different crater densities (lunar rock samples brought back to Earth have been dated!).

Assuming craters form at a constant (known) rate, it is possible to calculate the absolute age of planetary surfaces.

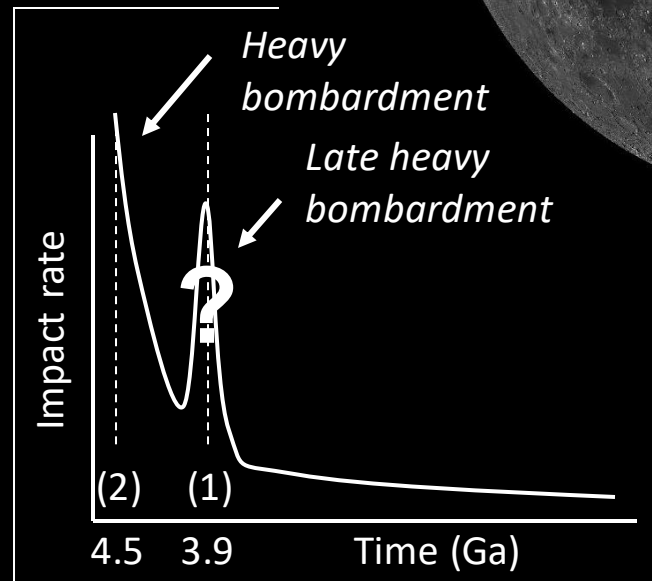
An old surface has more craters than a young surface.

It is possible to know the relative age of two surfaces by comparing their crater density.

- Main features of the Moon surface:
 - **Lowlands** (mare*, pl. maria)
Surface with few craters (younger), consisting of basalt that has filled large craters
 - **Highlands**
Surface with many craters (older)

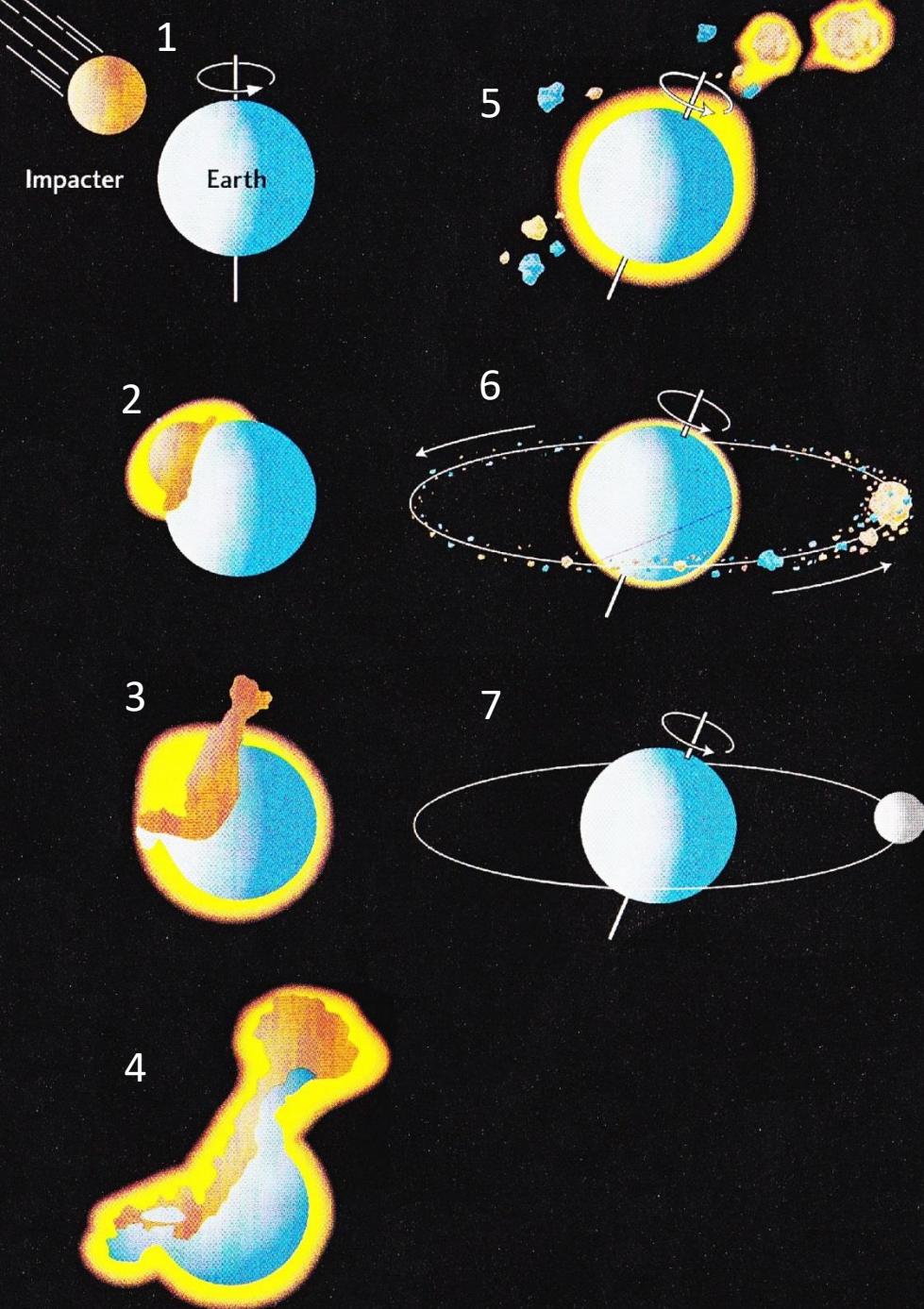


- Radiometric dating:
 - Lowlands (maria):
4 to 3.2 Ga **(1)**
 - Highlands:
4.5 to 4 Ga **(2)**



*Mare = latin word for "sea"

NB: Hadean Eon: 4.6-4 billion years

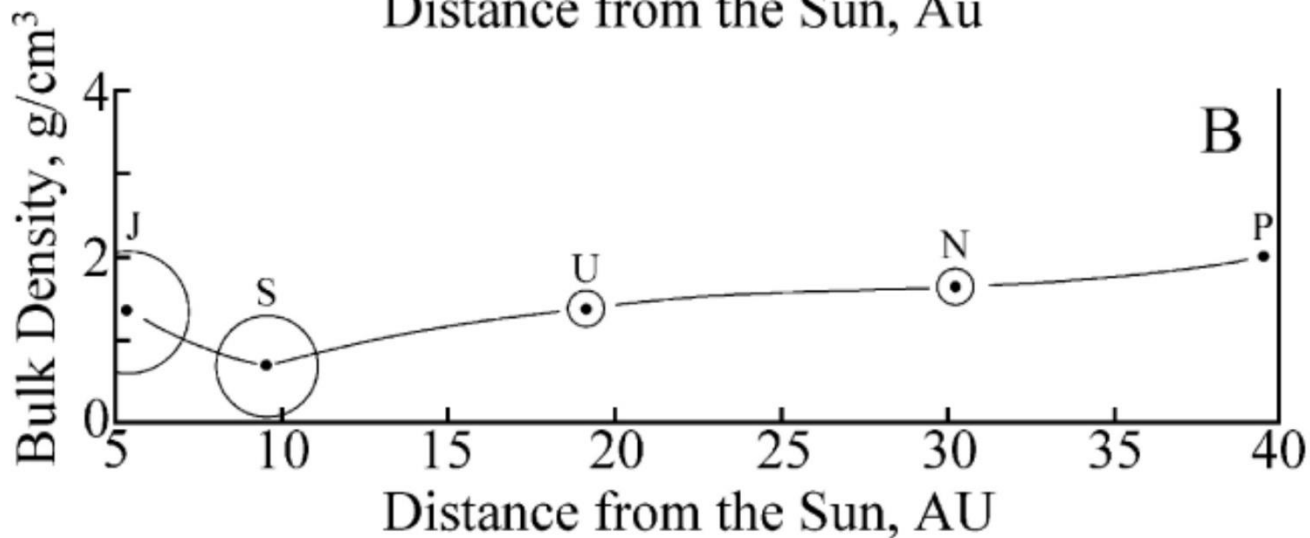
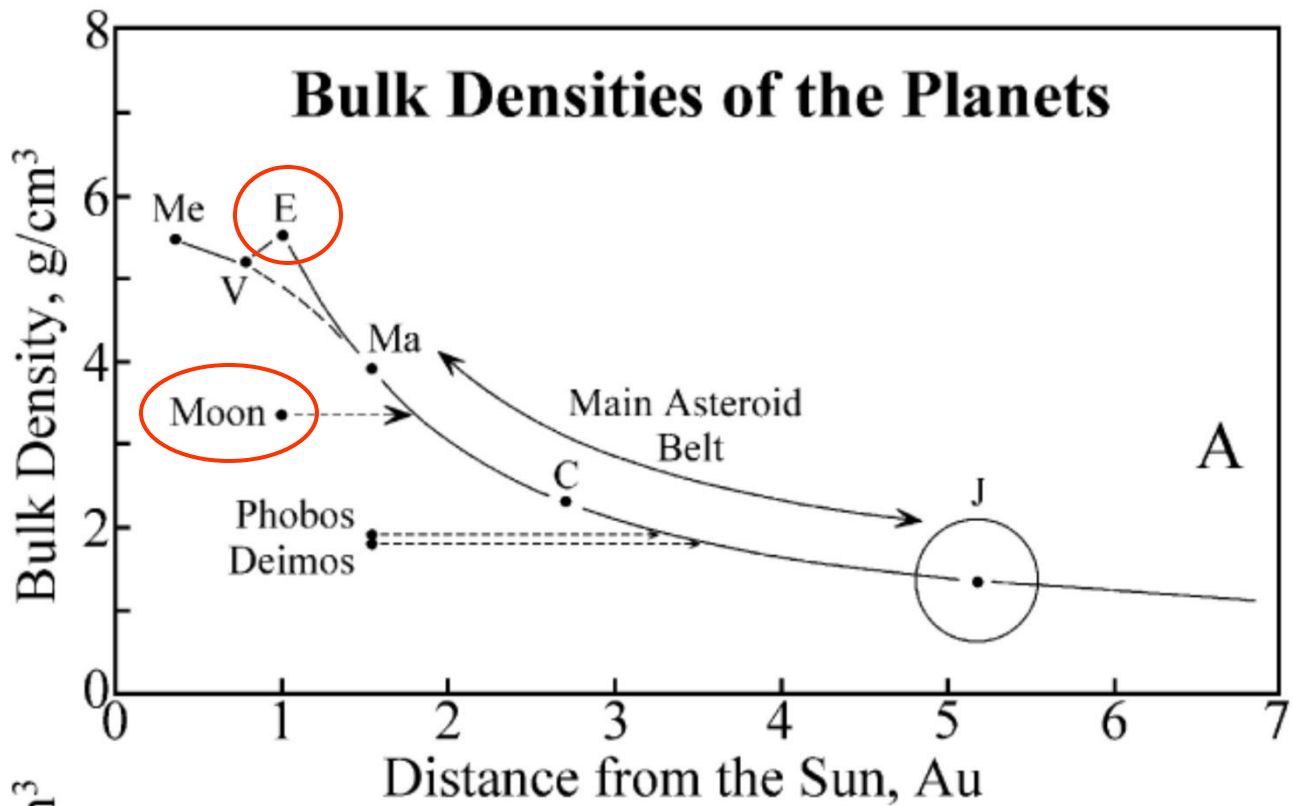


The Giant Impact Hypothesis suggests a large planetesimal impacted the Earth 4.6 billion years ago and the Moon formed out of the ejected material.

Explains:

- Earth's tilt
- Lack of volatile material on the Moon (vaporized during its formation)
- Tiny iron-rich core of the Moon (source material: only crust and mantle)

The Moon may have derived primarily from the crust and mantle of the Earth and impactor, explaining the low density of the Moon.



☀ The terrestrial planets and their characteristics

- Similarity

 - Layered structure (by gravitational differentiation)

 - A Fe-Ni core, a silicate mantle, and an outer crust

- Differences

 - Size

 - Mass

 - Dynamics (presence/absence of geological activity)

 - Atmospheric composition and pressure

 - Surface temperature

 - Density of craters (related to age of planetary surfaces)

■ Basic characteristics of the terrestrial planets

	Mercury	Venus	Earth	Mars	Moon
Radius (km)	2440	6052	6378	3388	1737
Mass (Earth=1)	0.06	0.81	1	0.11	0.01
	(large Fe-Ni core)		(6×10^{24} kg)		
Mean density (g/cm ³)	5.43	5.24	5.52	3.94	3.34
Orbit period (Earth days)	88	224	365	687	27
Distance from the Sun (millions of km)	57	108	148	228	
Av. atmospheric pressure (bar)	10^{-12}	92	1	7×10^{-3}	
Composition of atmosphere	(H ₂ , He)	CO ₂ , H ₂ SO ₄	N ₂ , O ₂ , CO ₂	CO ₂	
Temperature (°C)	-170 to 470	475	-60 to 50	-155 to 20	
DENSITY OF CRATERS	HIGH	LOW	LOW	HIGH	HIGH

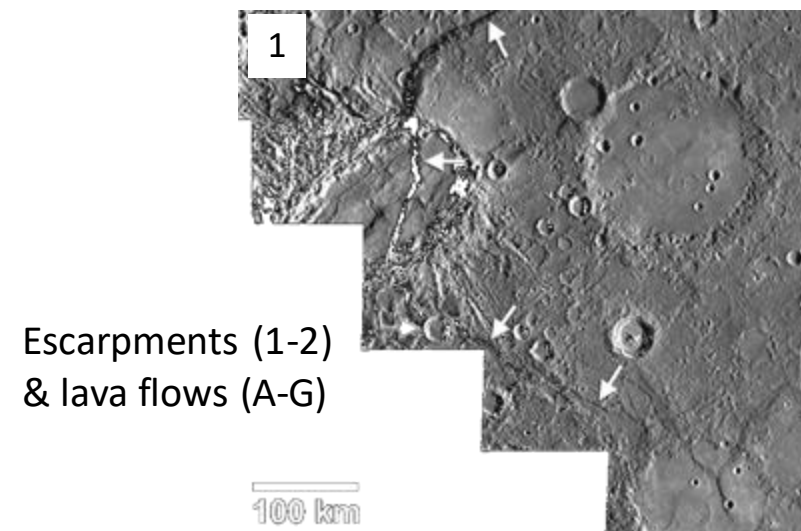
} Control on T

- Mercury

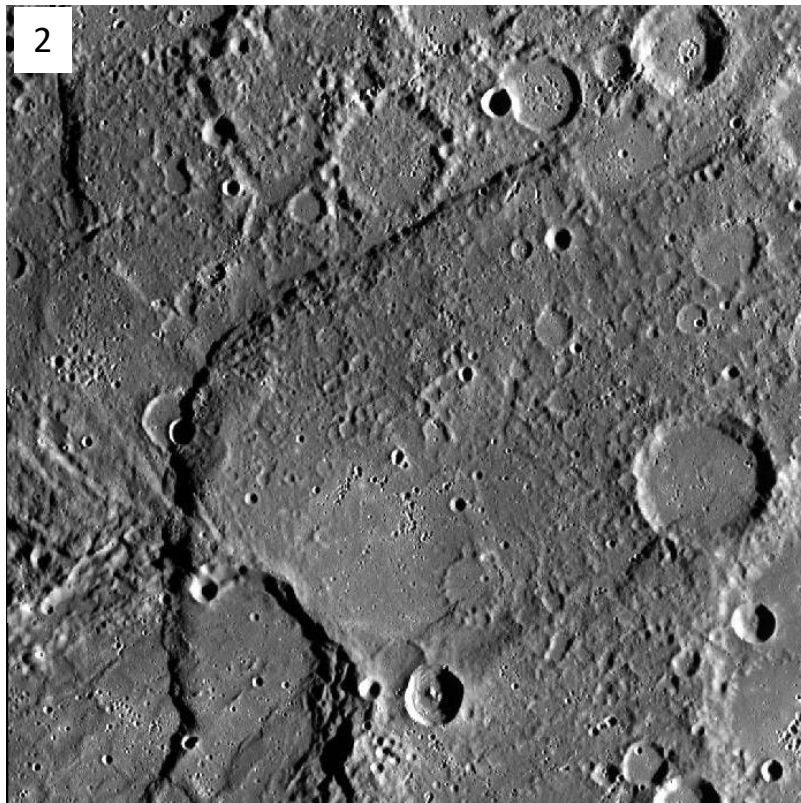
- Images of the surface of Mercury taken by Mariner 10 in 1974 and by Messenger (2004-2015).
- Very old surface: many craters
- Numerous escarpments (1-2 km high, 100s of km long), probably faults resulting from contraction during cooling of the crust.



Moon-like surface



Escarpments (1-2)
& lava flows (A-G)



Solomon et al. (2008)

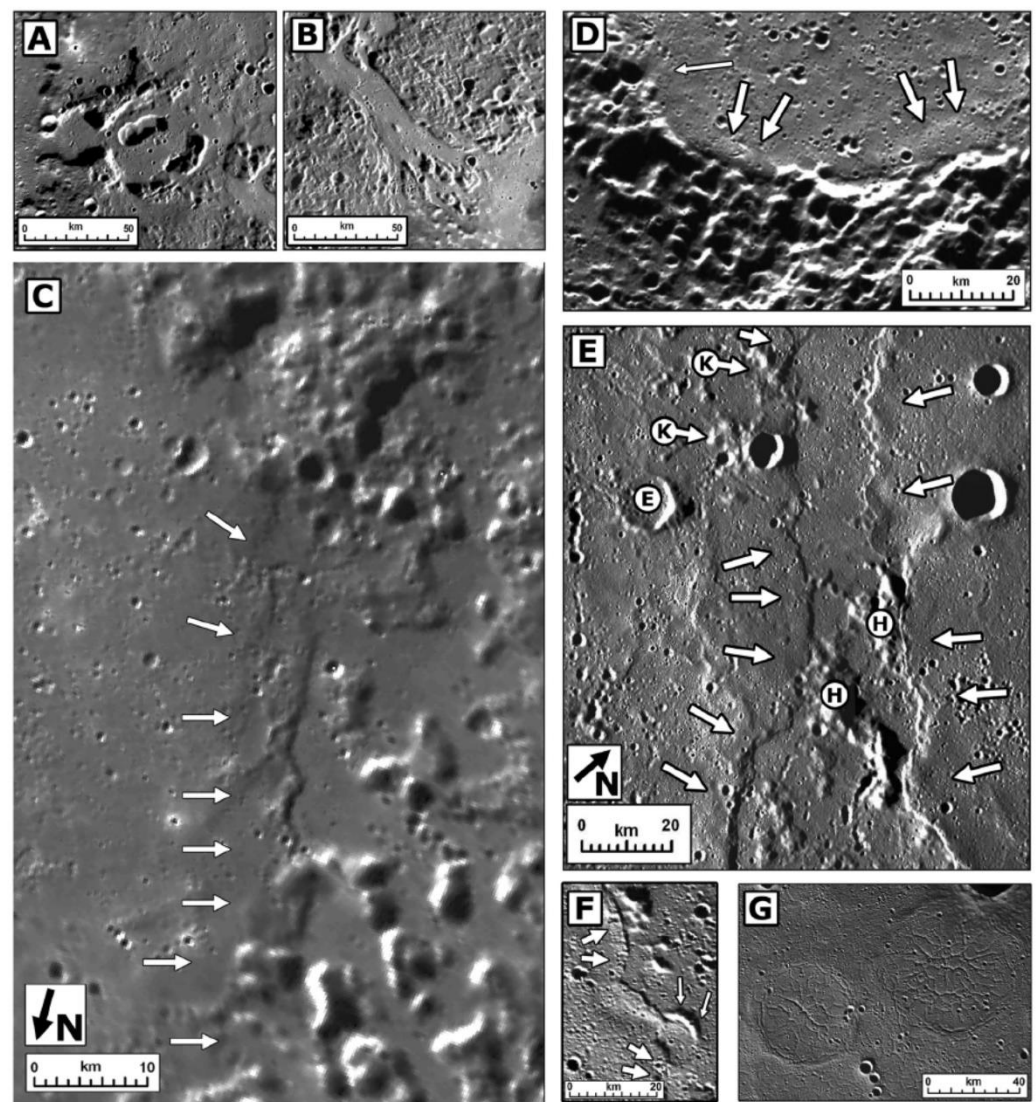


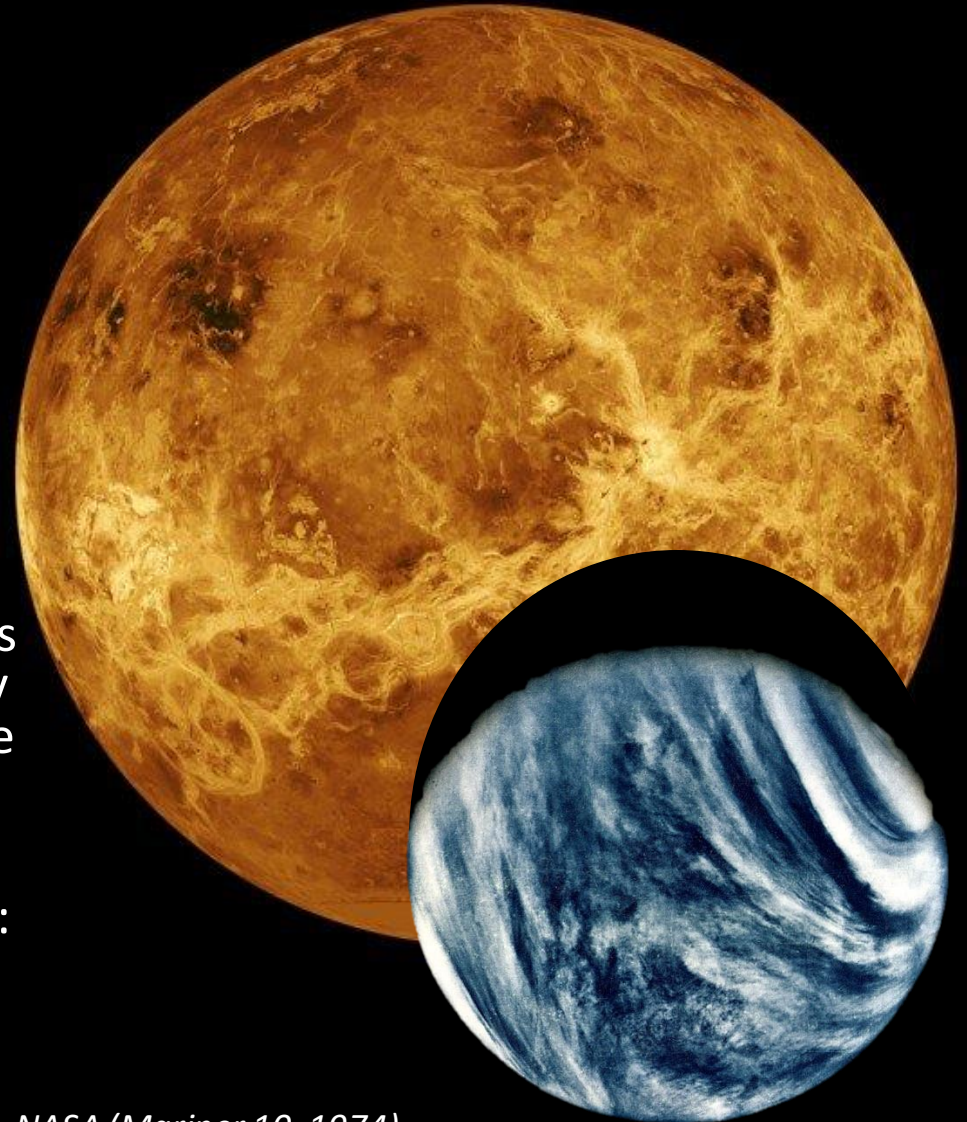
Fig. 3. Features interpreted to be related to lava flows in the northern smooth plains and vicinity. (A) Pits interpreted as source vents. (B) Teardrop-shaped hills and channel. (C) Distal smooth plains (left) embaying rough plains (right) (see Fig. 2) along a flow front-like contact (arrows). (D) Steep flow margin (broad arrows) within a flooded impact crater; narrow arrow points to possible flooding of a smaller crater. (E) Candidate lava flow fronts (arrows) embaying crater (E) and flooding the hills (H) between flow fronts to form kipukas (K). (F) Evidence for a flow front (broad arrows) descending into a preexisting crater (narrow arrows; note adjacent depressed part of flow). (G) Fissures inside flooded craters in the Goethe basin (G in Fig. 1). The location of these images and MESSENGER image numbers are given in the SOM.

Head et al. (2011)

- Venus

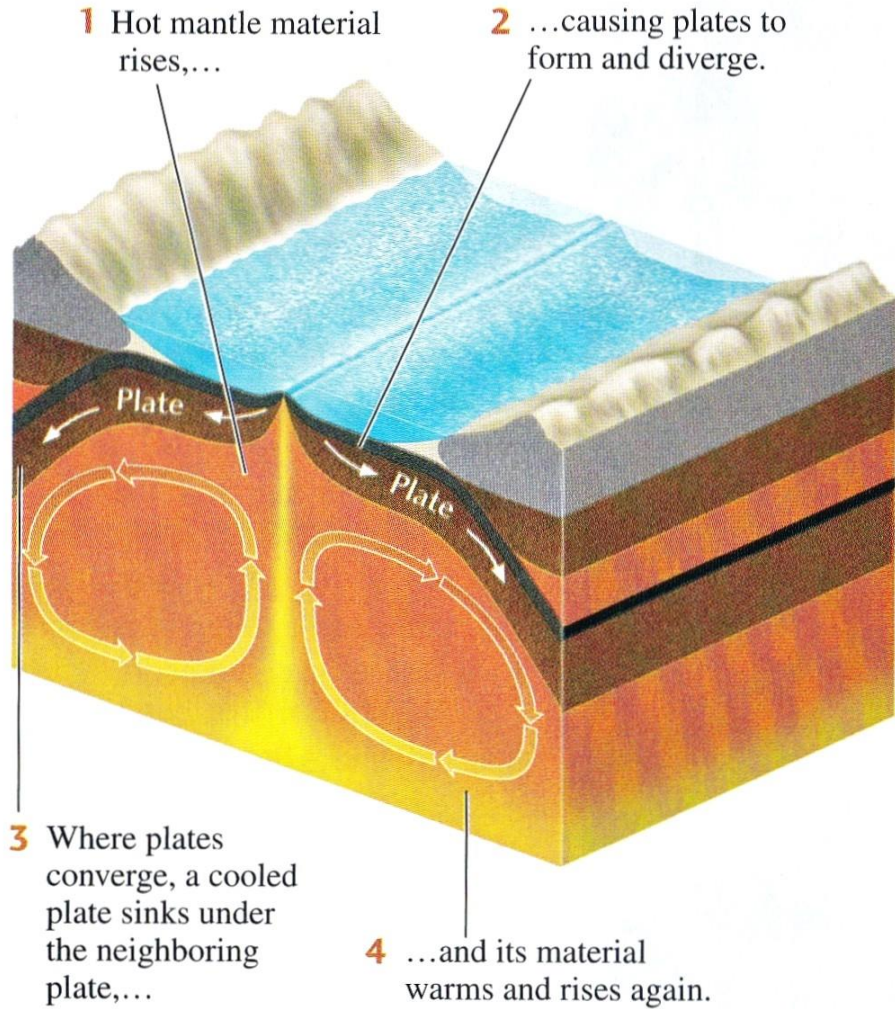
NASA (Magellan radar image)

- First flyby by Mariner 2 in 1962
- First images of the surface by Russian Venera spacecrafts in 70s and 80s
- Very dense atmosphere of carbon dioxide, clouds of sulfuric acid and water vapor
- Very young surface with few craters because of intense tectonic activity (“flake tectonics”), and many active volcanoes.
- No rainfall on the surface (too hot): very little erosion



NASA (Mariner 10, 1974)

(a) Plate tectonics on Earth



(b) Flake tectonics on Venus

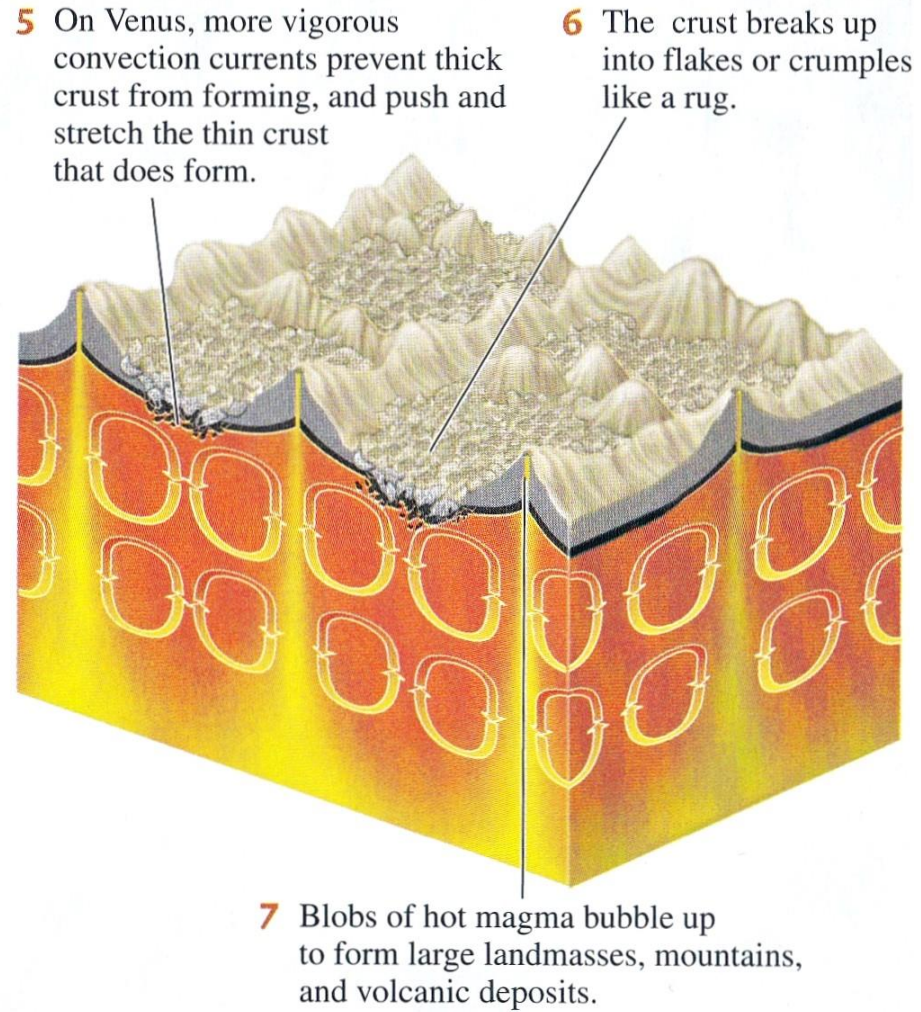
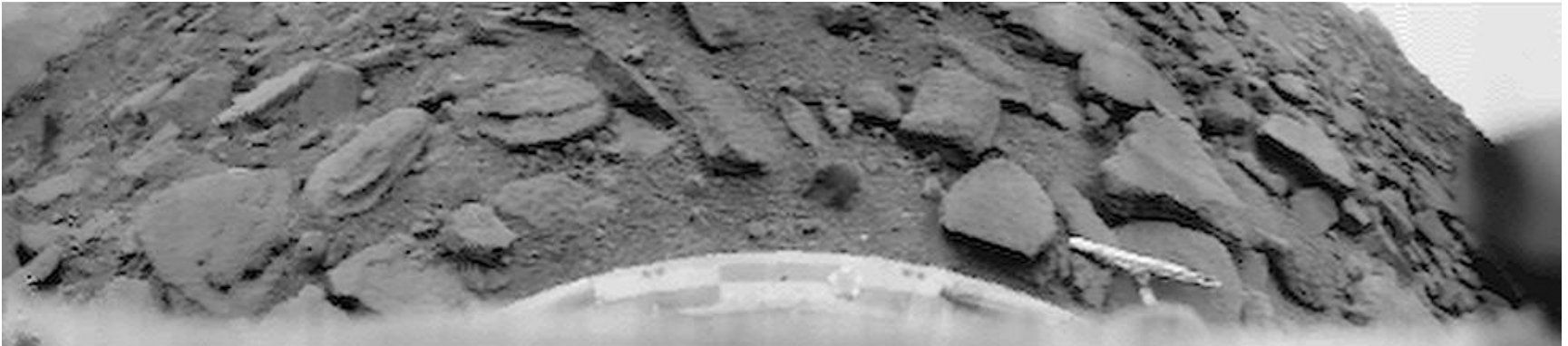


FIGURE 9.16 ■ Flake tectonics on Venus is very different from plate tectonics on Earth, but could be similar to tectonic processes on early Earth.

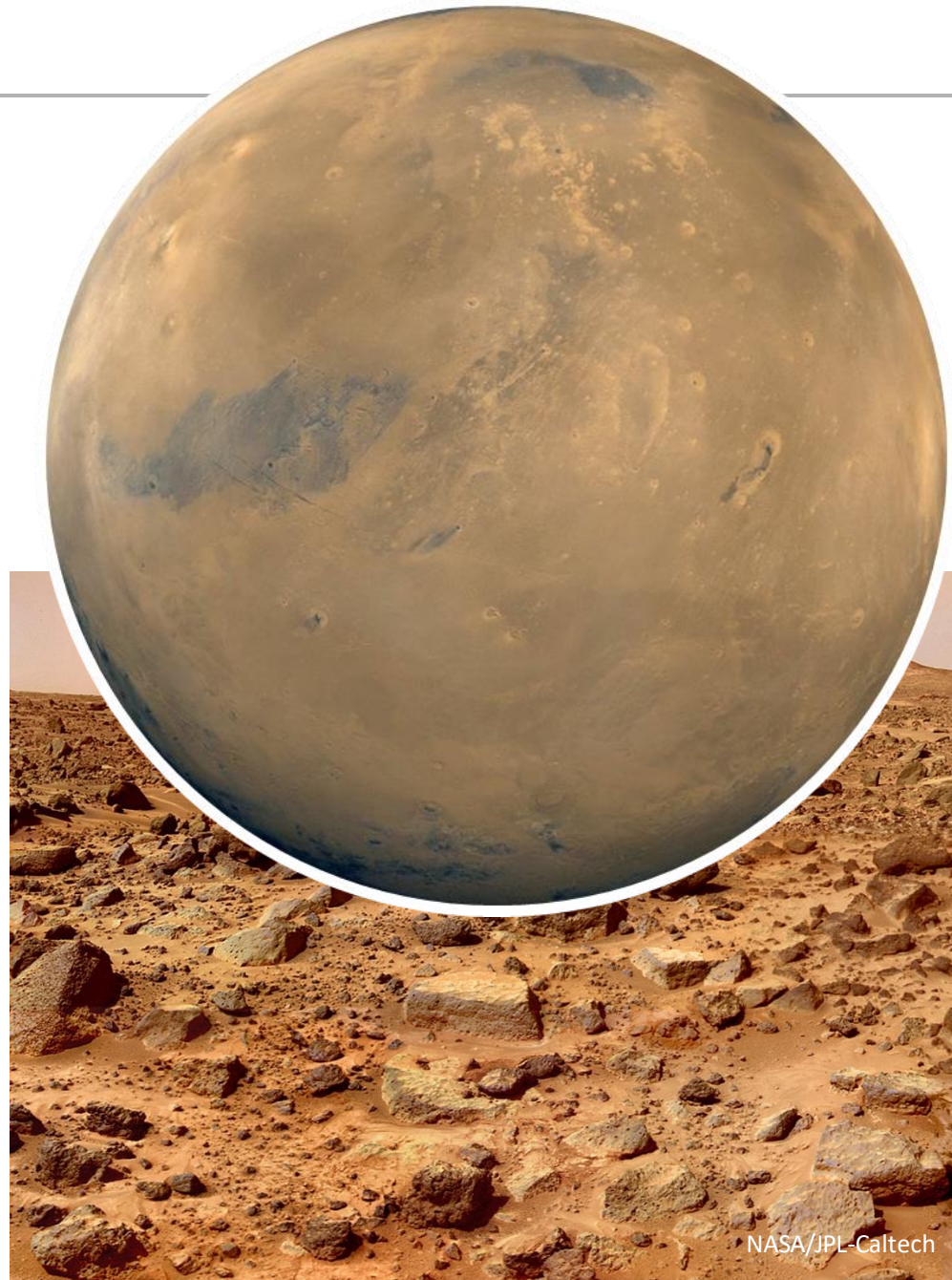
Venera missions to Venus launched by the Soviet Union (1961-1983)



*Venera 9 Lander image taken in 1979
Venera 13 Lander image taken in 1981*

- Mars

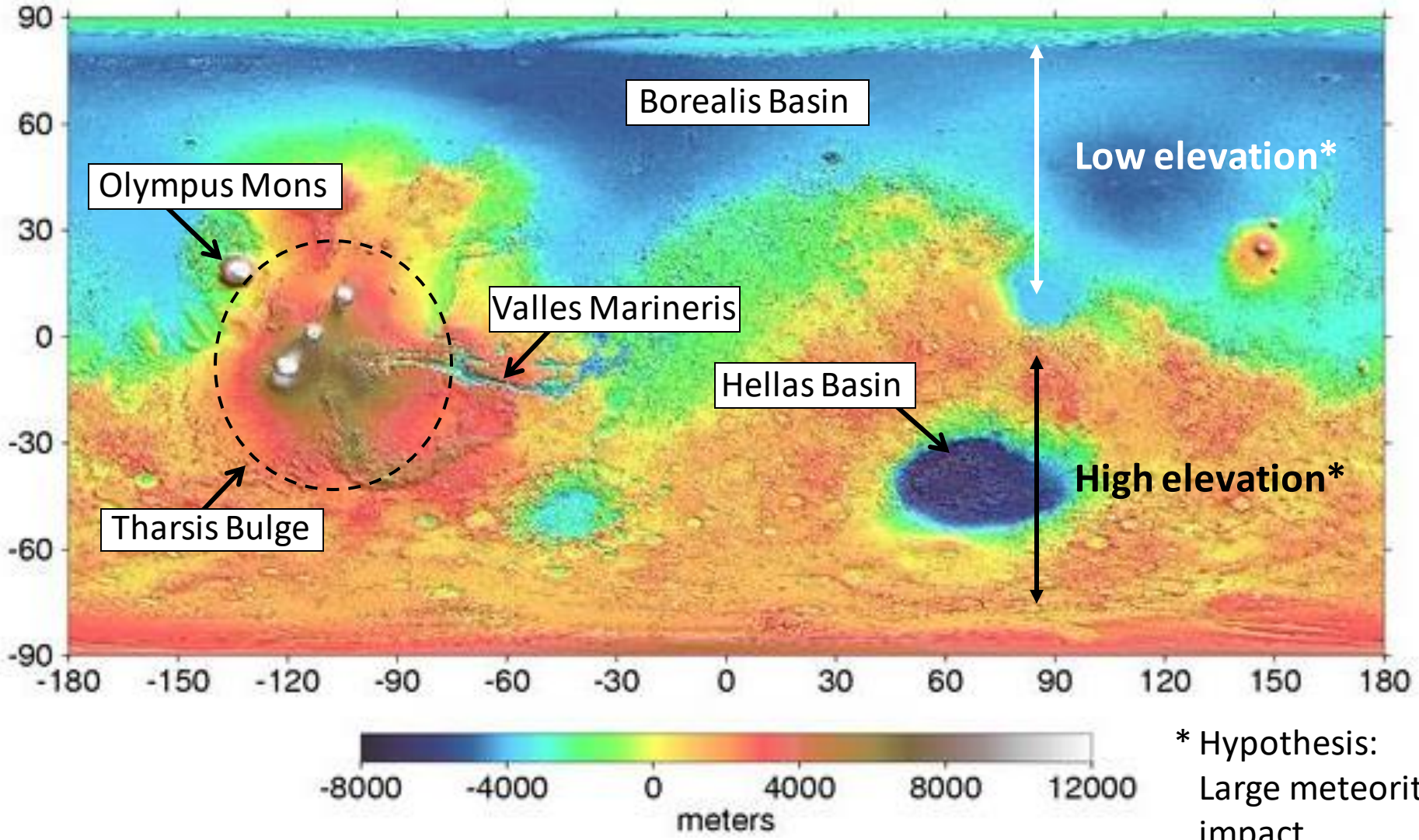
- Mars is red because its surface is covered with iron oxide/hydroxide dust (rust).
- Water occurs on Mars as ice on and below the ground and as gas in the atmosphere.
- The low temperature and low atmospheric pressure means that liquid water cannot occur today permanently at the surface of Mars.
- There is evidence of liquid water present on the surface of Mars in the distant past.



Mars topography

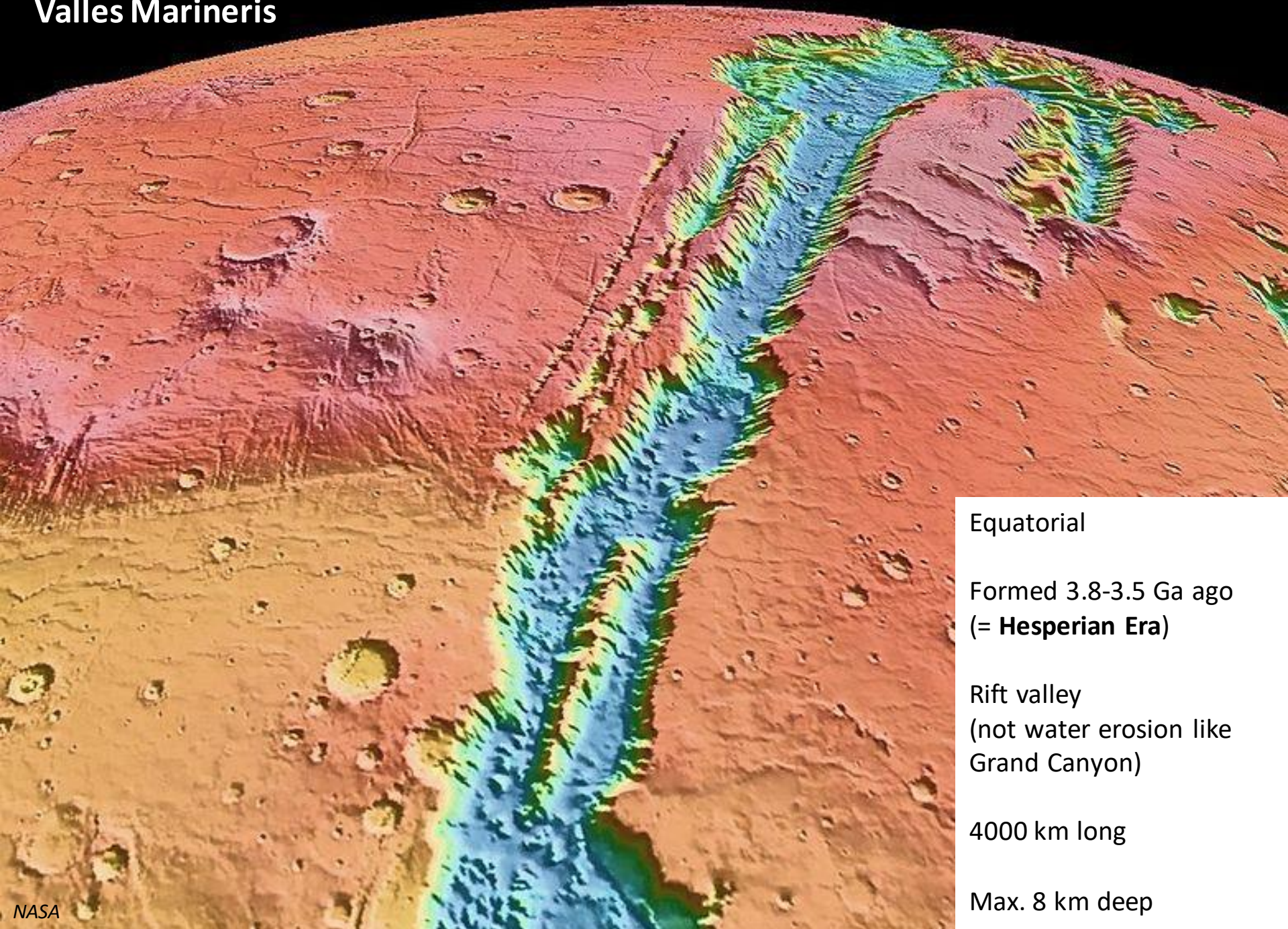
Olympus Mons height = 25 km

Valles Marineris depth = 8 km



* Hypothesis:
Large meteorite
impact

Valles Marineris



Equatorial

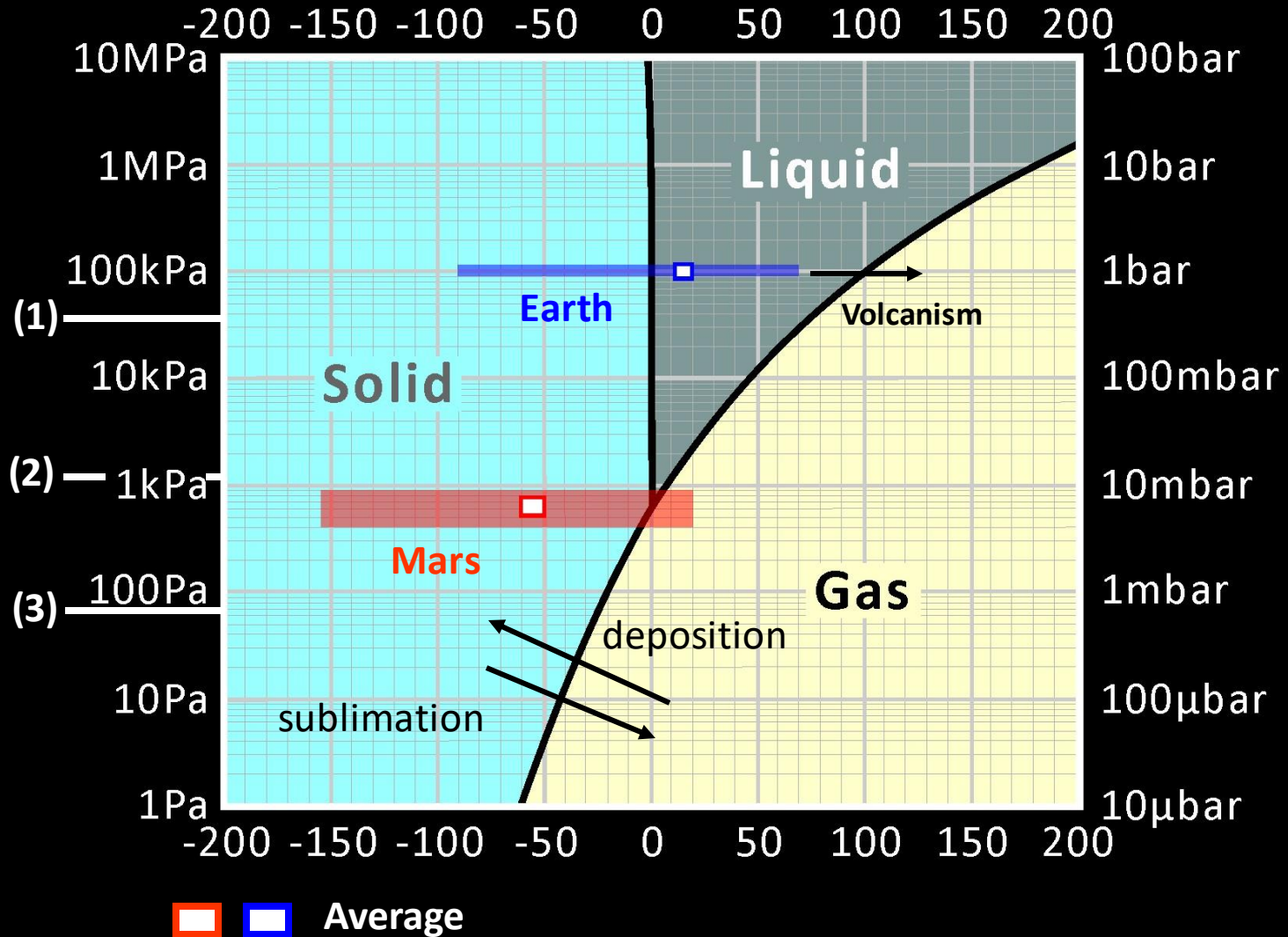
Formed 3.8-3.5 Ga ago
(= **Hesperian Era**)

Rift valley
(not water erosion like
Grand Canyon)

4000 km long

Max. 8 km deep

PHASE DIAGRAM OF WATER



(1) Atmospheric pressure at the top of Mount Everest, 8,850 m (0,3 bars)

(2) Atmospheric pressure at the bottom of Hellas Planitia, a 7-km deep depression on Mars (14 mbars)

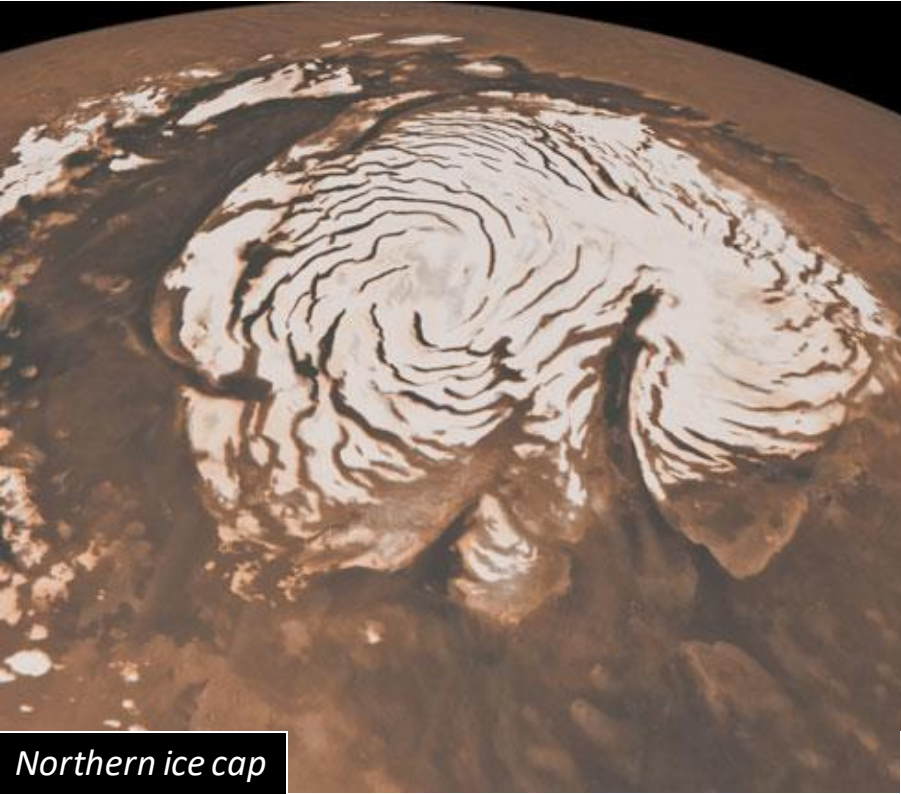
(3) Atmospheric pressure at the top of Olympus Mons, highest elevation on Mars, 25 km (3 mbars)

Water on Mars

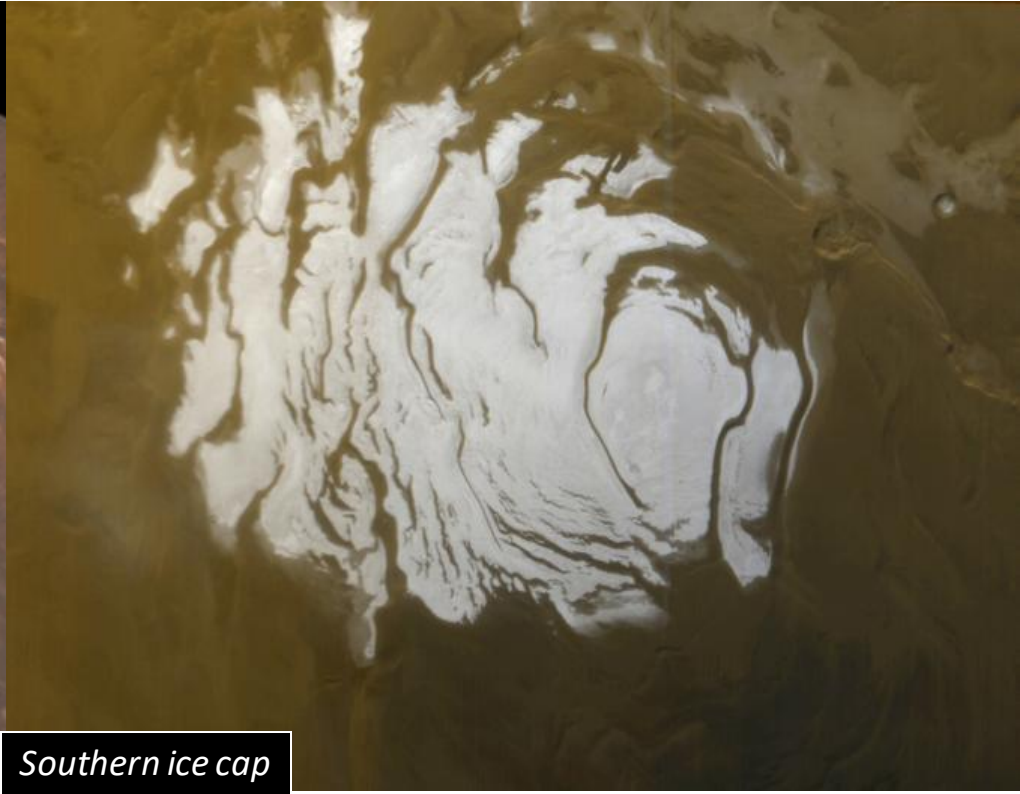
- Current conditions on Mars are not favorable for the formation of liquid water at or near the surface for an extensive period of time (very low temperatures and very low atmospheric pressures). There is however evidence for seasonal flows of liquid brines.
- Water is mainly present as **ICE**:
 - POLAR ICE CAPS** composed of water ice (H_2O) and dry ice (CO_2).
 - SUBSURFACE ICE** found from mid-latitudes to poles (soil prevents ice from sublimating).
 - FROST** at the surface of the ground (not permanent).
- Some water is also present as **GAS** in the atmosphere.

POLAR ICE CAPS are composed of an external layer of CO₂ ice (dry ice) underlain by water ice (in summer, water ice is exposed to the air in the NH but not in the SH)

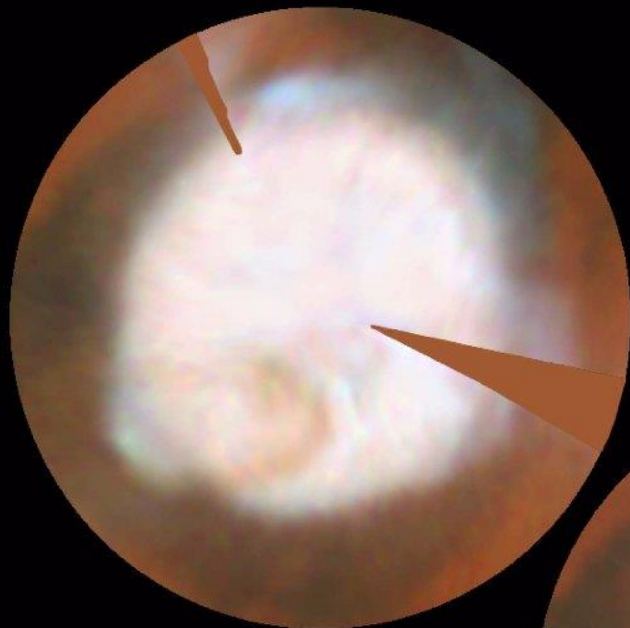
Northern Ice Cap



Southern Ice Cap

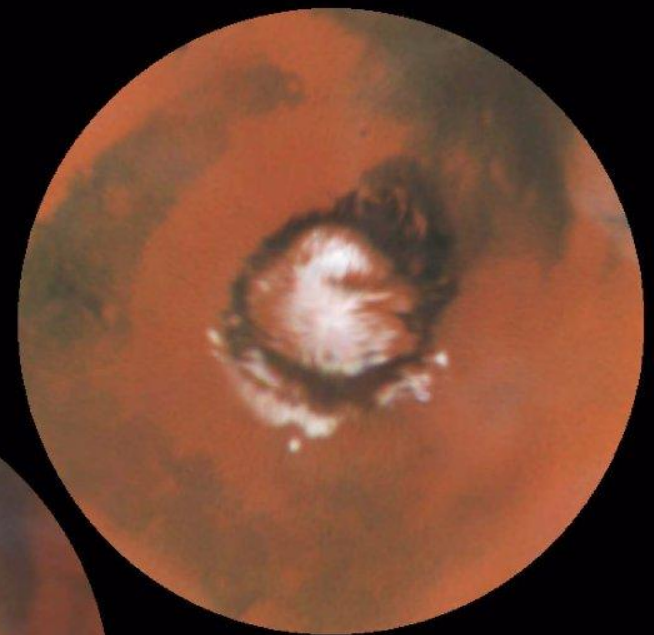
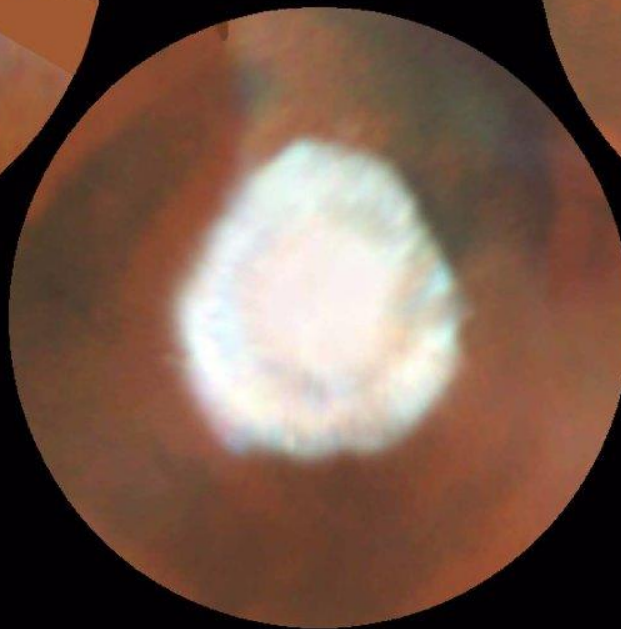


The dry ice sublimates in summer (seasonal ice). This results in large seasonal variations of CO₂ in the Martian atmosphere, which cause large variations of atmospheric pressure.



October 1996

January 1997



March 1997

Mars • North Polar Cap
Hubble Space Telescope • WFPC2

SUBSURFACE ICE

Images taken by Phoenix Mars Lander on June 15 and 19 2008

Sol 20

Sol 24



Ground ice sublimation

NASA/JPL-Caltech/University of Arizona/Texas A&M University

FROST

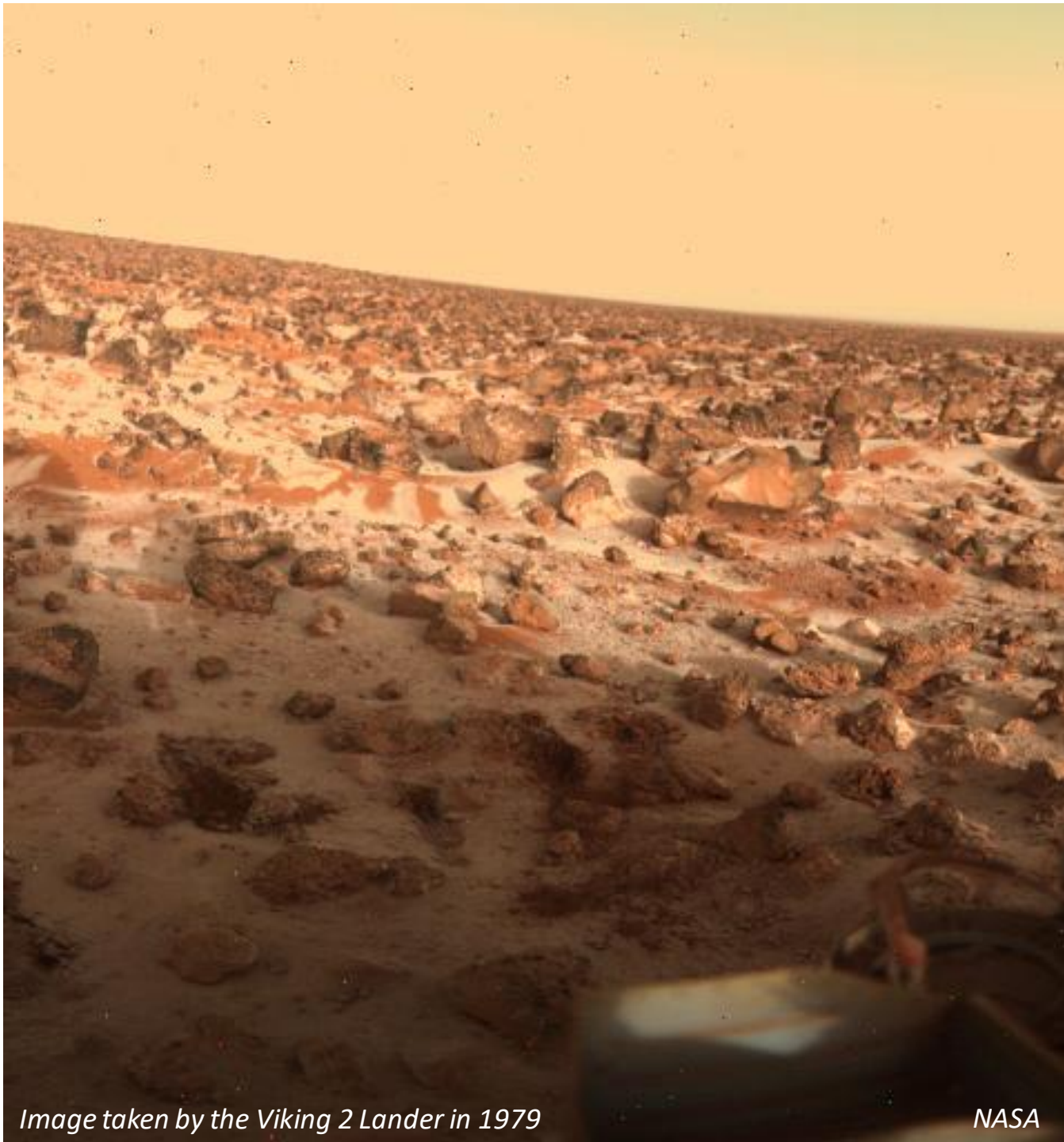


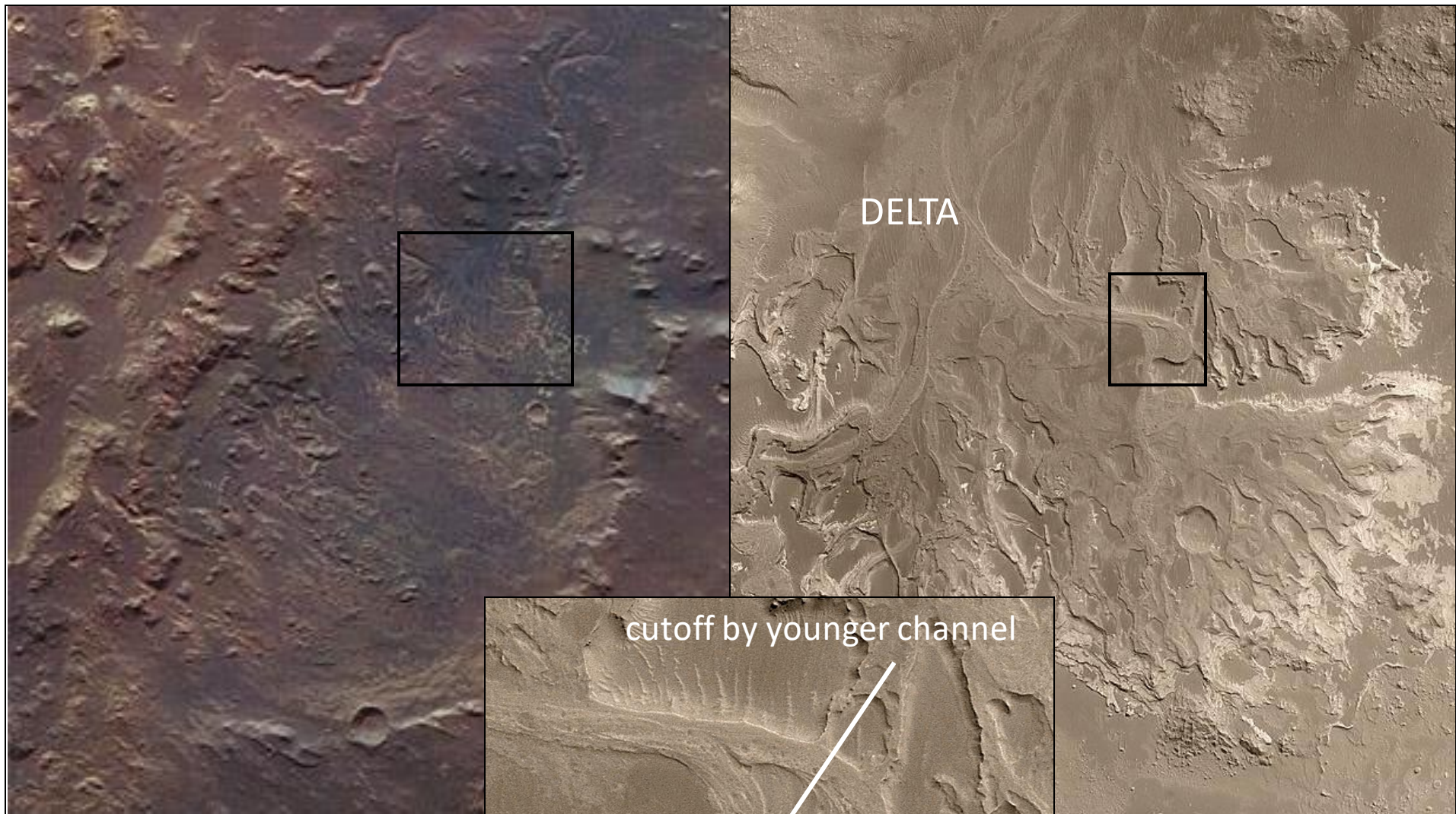
Image taken by the Viking 2 Lander in 1979

NASA

Evidence for the presence of liquid water in the past

Various lines of evidence (mostly 3-4 Ga ago)

1. Surface morphological features
 - MEANDERING STREAM BEDS, DELTAS**
 - OUTFLOW CHANNELS** (extremely large channels)
 - Frozen sea?**
2. Lithology and sedimentology
 - LAYERED SEDIMENTARY ROCKS**
 - CROSS-STRATIFICATIONS**
 - RIVER CONGLOMERATES**
3. Mineralogy
 - MINERALS** that form in the presence of liquid water



DELTA

cutoff by younger channel

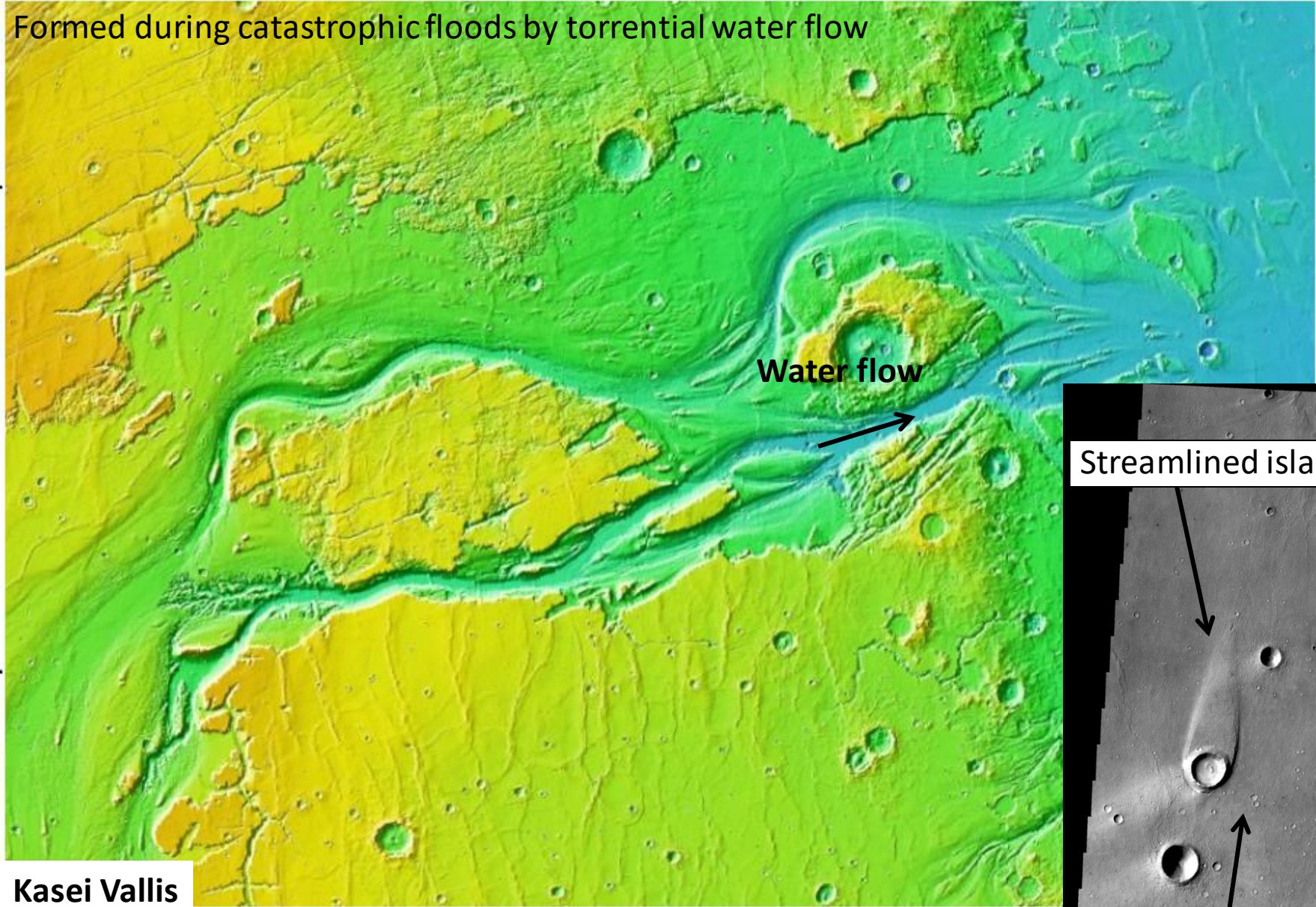
MEANDERING STREAM BED

Eberswalde crater
NASA/JPL-Caltech/MSSS

Large outflow channels (up to 100 km wide and 1000 km long)

Formed during catastrophic floods by torrential water flow

30°



Water flow

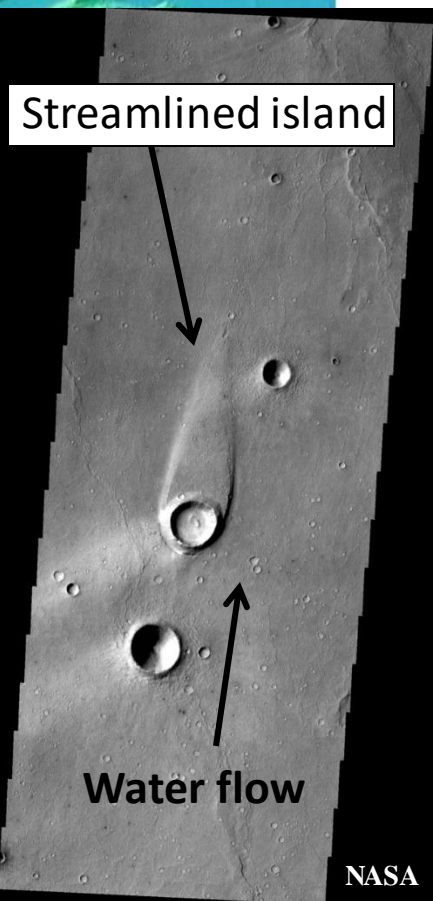
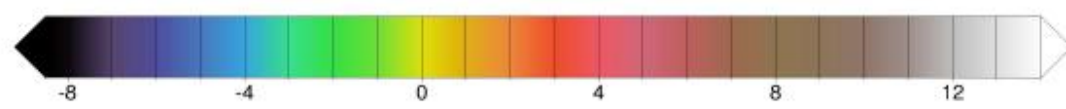
20°

Kasei Vallis

290°

300°

Wikipedia

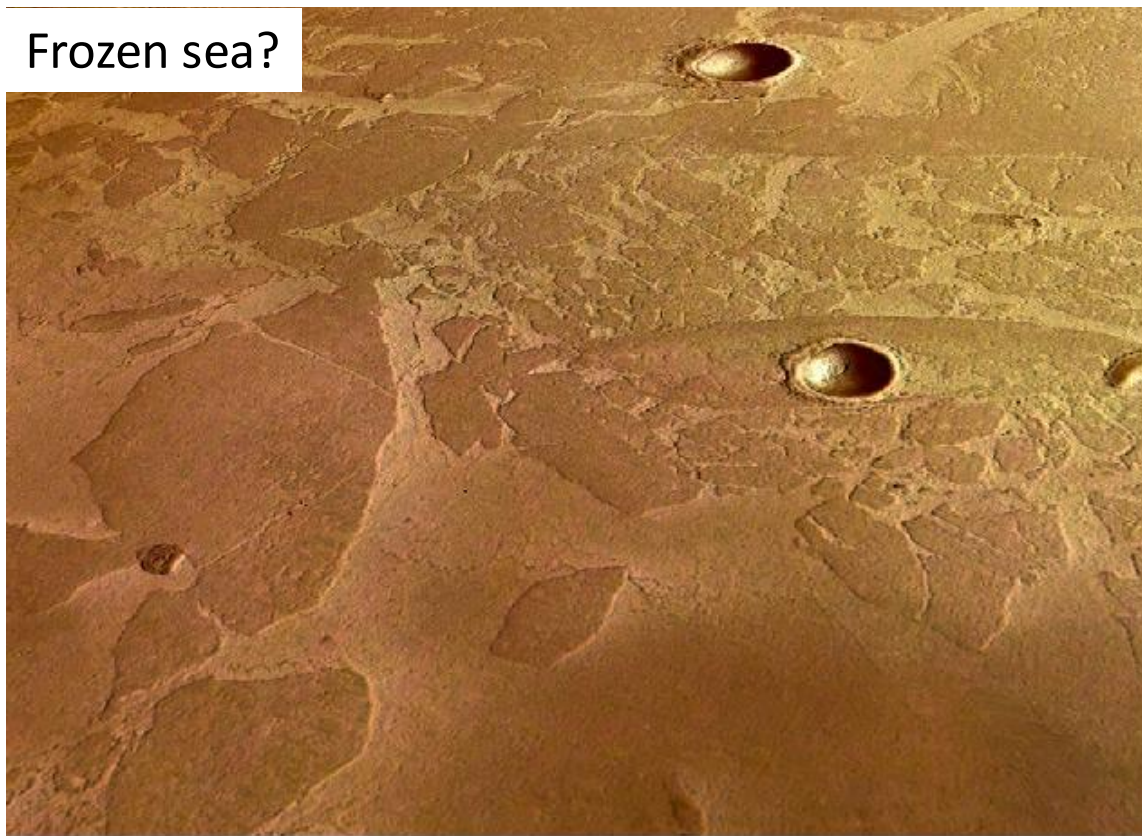


Streamlined island

Water flow

NASA

Frozen sea?

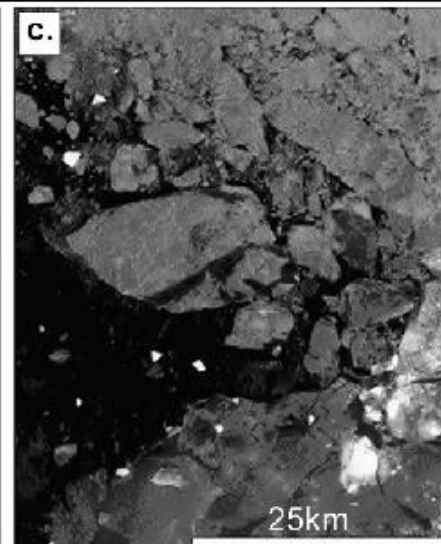
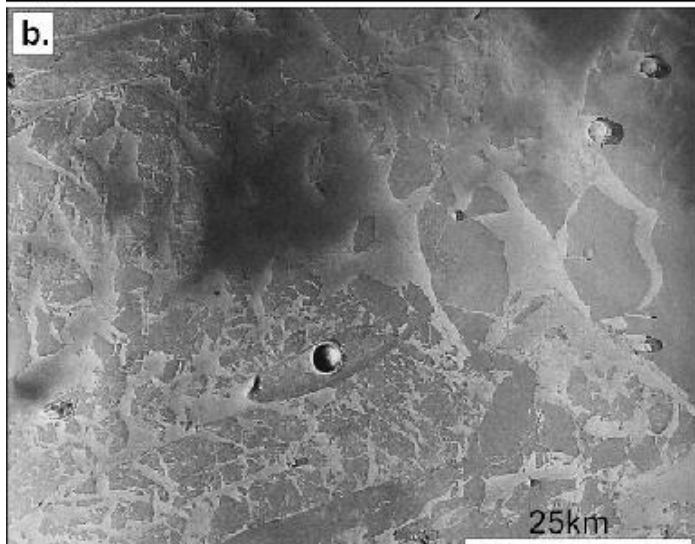


copyright: ESA/DLR/FU Berlin (G. Neukum)

Images of the 'frozen sea' on Mars (a,b) from the High Resolution Stereo Camera of the ESA Mars Express mission, and pack-ice (c) in the terrestrial Antarctic.

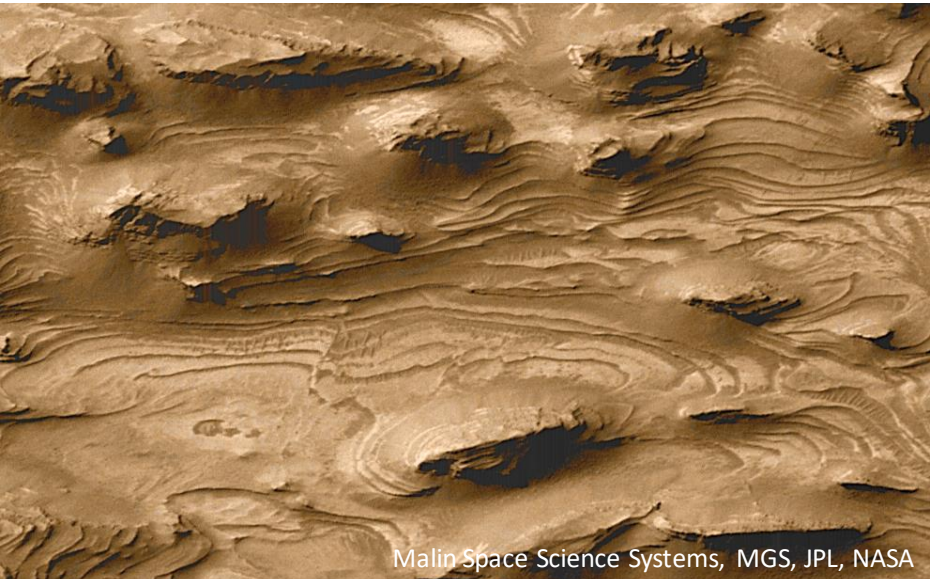
Source: Matt Balme (the Open University)

Water probably came from an underground source and came out through fractures. The water froze and formed pack ice before it became completely solid. The ice was covered by a layer of dust which prevented sublimation.

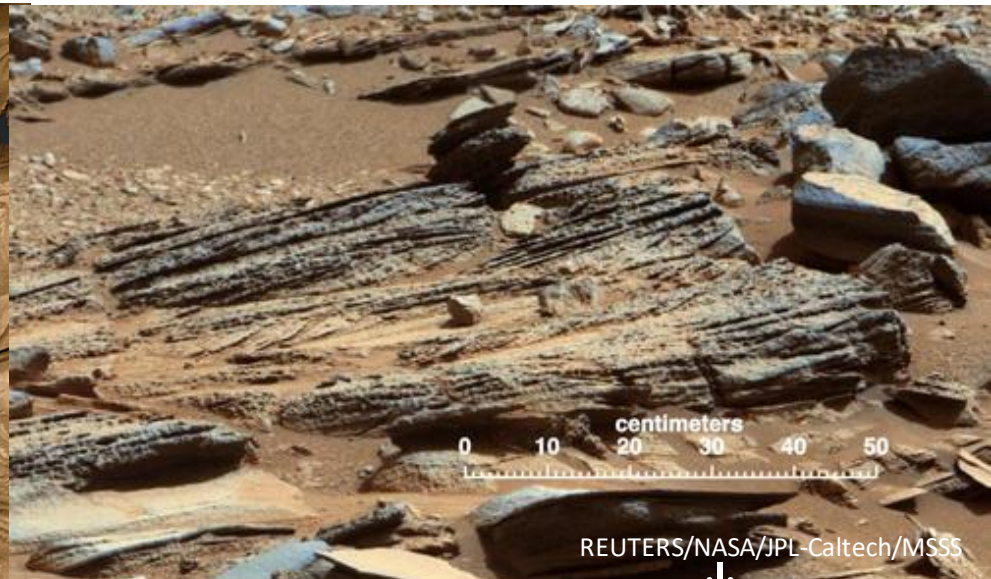


Lithology and sedimentology

Layered rocks in Candor Chasma



Cross-stratification (evidence of stream flow)



Ancient stream bed composed of sedimentary gravels and pebbles



Mars



Earth



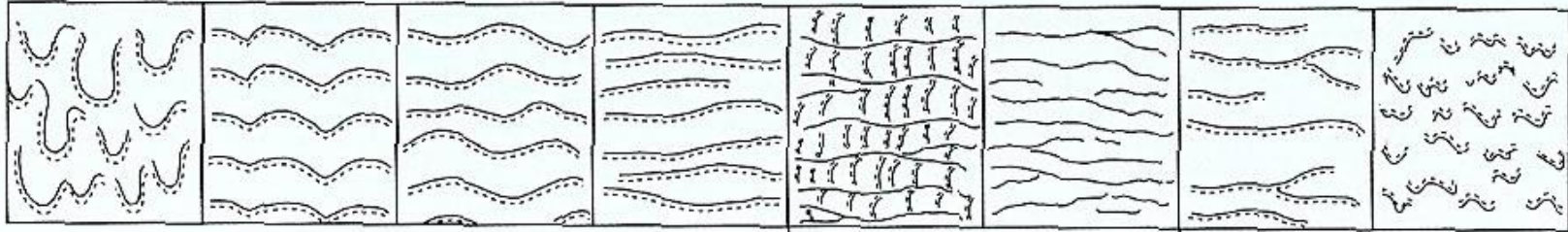
current ripples

wave ripples

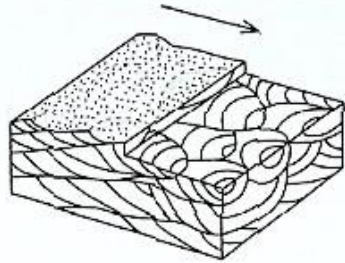
wind ripples

linguoid catenary undulatory straight

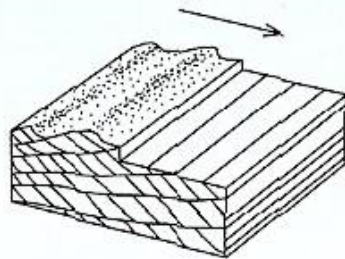
platform



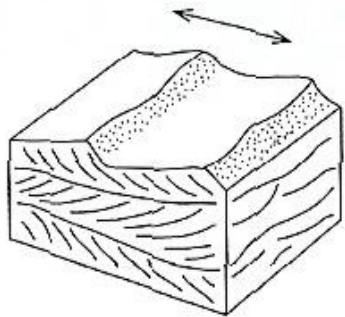
transverse section with preservation potential



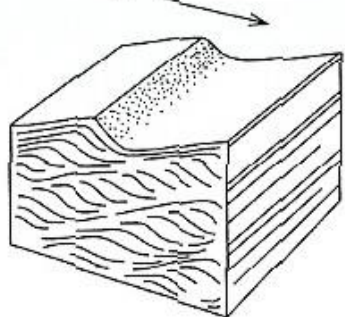
trough cross-lamination



tabular cross-lamination, low-angle of climb

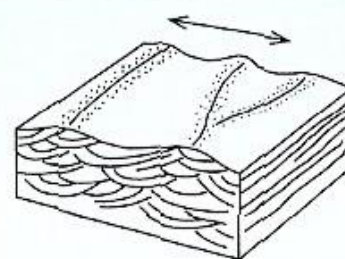


herringbone (bi-directional) cross-lamination

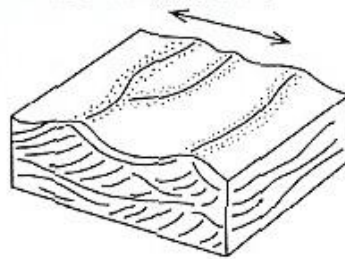


tabular cross-lamination high-angle of climb, stoss-side preservation climbing-ripples, high rate sediment fallout

moderate to good preservation potential

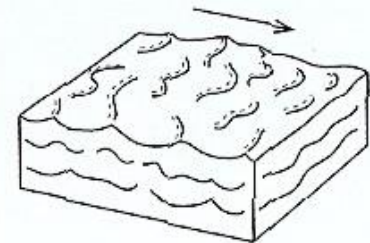


symmetrical wave ripples with upbuilding of bi-directional cross-lamination

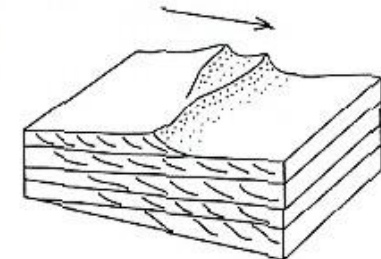


symmetrical-asymmetrical wave ripples with unidirectional cross-lamination

moderate preservation potential



adhesion ripples



aerodynamic ripples

low preservation potential

Deposition, exhumation, and paleoclimate of an ancient lake deposit, Gale crater, Mars

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The landforms of northern Gale crater on Mars expose thick sequences of sedimentary rocks. Based on images obtained by the Curiosity rover, we interpret these outcrops as evidence for past fluvial, deltaic, and lacustrine environments. Degradation of the crater wall and rim probably supplied these sediments, which advanced inward from the wall, infilling both the crater and an internal lake basin to a thickness of at least 75 meters. This intracrater lake system probably existed intermittently for thousands to millions of years, implying a relatively wet climate that supplied moisture to the crater rim and transported sediment via streams into the lake basin. The deposits in Gale crater were then exhumed, probably by wind-driven erosion, creating Aeolis Mons (Mount Sharp).

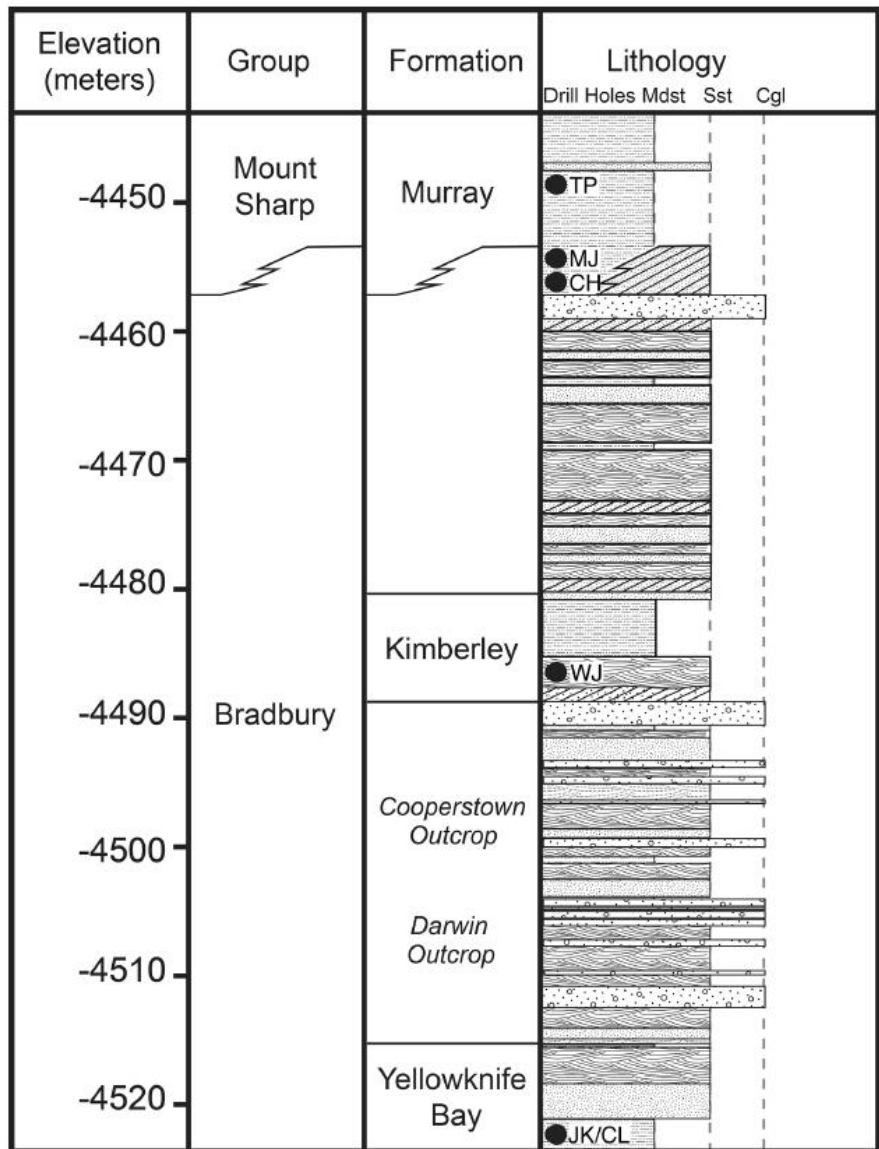


Fig. 2. Stratigraphic column for the sedimentary facies from Yellowknife Bay to Pahrump Hills.
 The contact between the Bradbury group (Aeolis Palus) and the Mount Sharp group (Aeolis Mons) is marked by interfingering of facies. Mdst, mudstone; Sst, sandstone; Cgl, conglomerate.

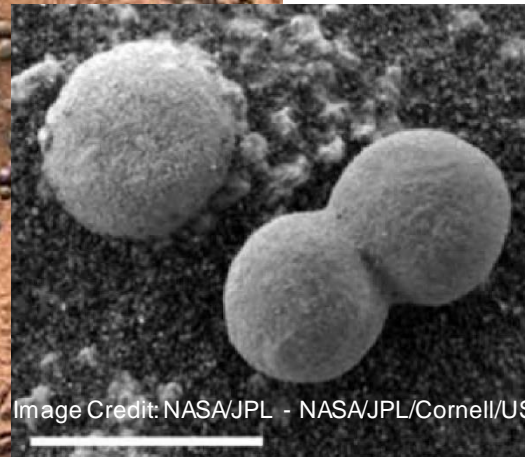
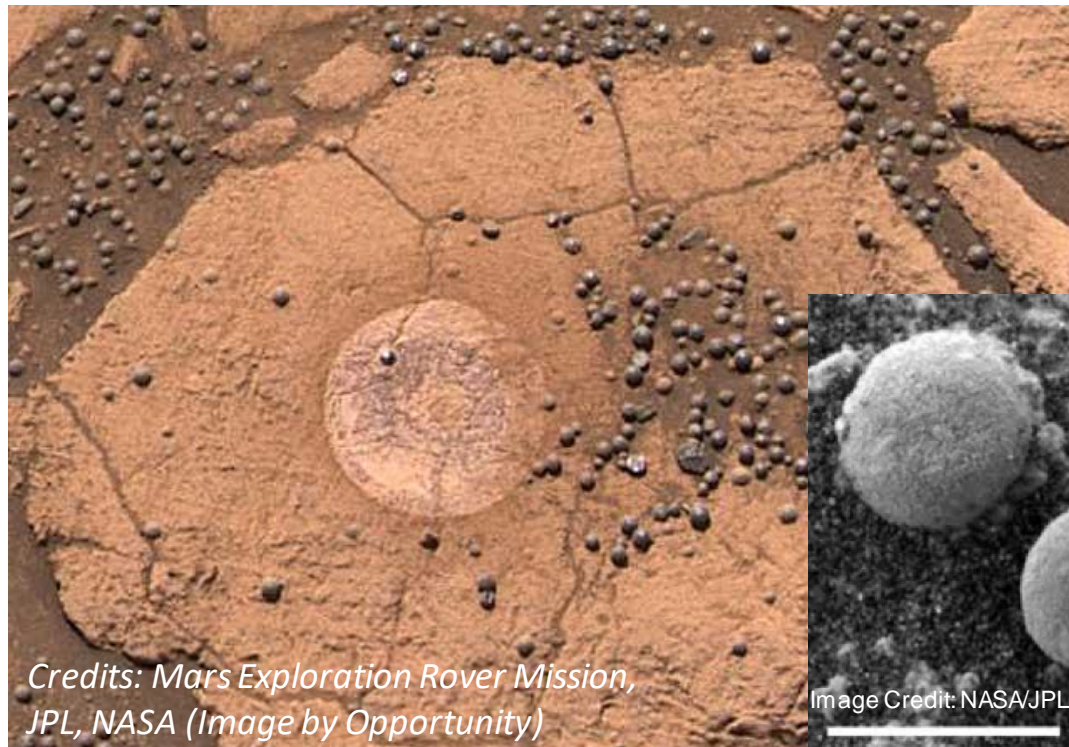
Minerals forming in the presence of liquid water and found on Mars:

1. Chloride salts (e.g. NaCl = halite)
2. Hydrated sulfates (e.g. $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$ = gypsum)
3. Hematite (Fe_2O_3 , “blueberries”)
4. Clay minerals
5. Hydrated silica ($\text{SiO}_2 \cdot n\text{H}_2\text{O}$ = opal)
6. Carbonates (e.g. CaCO_3 = calcite)

Note that the presence of these minerals does not necessarily imply the presence of surface water. These minerals could form underground in presence of groundwater.

“Blueberries”: spherules of hematite (max 4 mm large) probably formed below ground surface when Fe^{2+} -rich fluids met with oxidizing groundwater

Equivalent on Earth:



Mars wet past: two competing hypotheses

4-3 billion years ago, Mars experienced...

1. A sustained warm, wet period with frequent rain (this implies an atmosphere thicker than today)
= Life-friendly environment
2. An icy and dry period with short-lived catastrophic floods resulting from ice melting events after meteorite impacts or periods of enhanced volcanism
= Life-unfriendly environment

Consensus on the fact that Mars has been icy and (mostly) dry for the past 3 billion years.

But the underground could be the next place to look for liquid water and perhaps forms of life...

Andromeda was once thought to be a proto-planetary disk. Now we know that it is a galaxy composed of billions of stars.

Since the 1990s, hundreds of exoplanets have been discovered and many more will be discovered in the future. Many of these planets could potentially be suitable places to shelter life.

1. Carbonaceous chondrite (Allende meteorite)

- Presence of Ca-Al-rich inclusions (white)
Amongst first solids to have condensed in the protoplanetary disk (oldest age measured in the solar system! → 4.567 Ga)
- Chondrules (rich in olivine and pyroxene)
- Rich in carbon, organic compounds

Undifferentiated, composition of C-type asteroids

2. Ordinary chondrite

- Chondrules (rich in olivine and pyroxene)
- Rich in iron, nickel

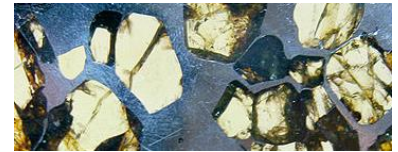
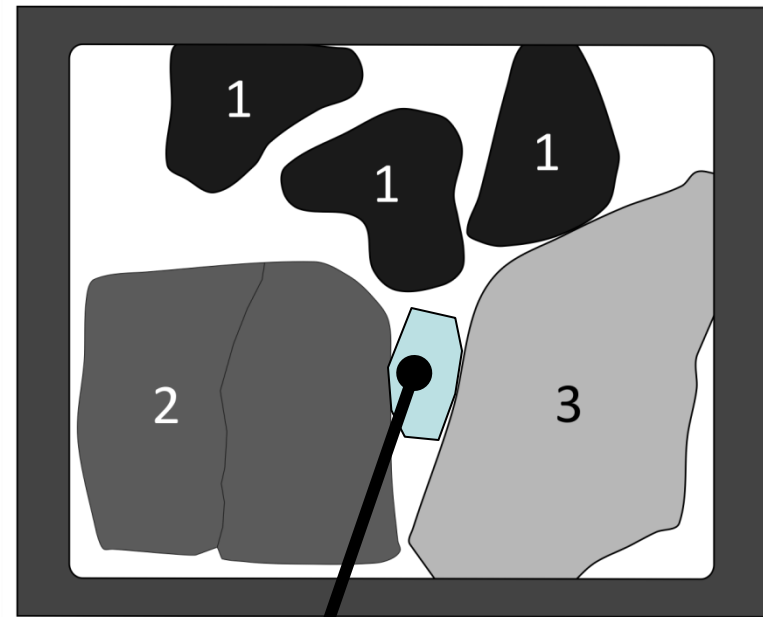
Undifferentiated, composition of S-type asteroids

3. Iron meteorite (Gibeon meteorite)

- Rich in iron, nickel
- Fragment of the core of a differentiated object



CHONDRULES: few μm to 1 cm spheres made of silicate minerals formed when droplets of molten rock solidified in space → amongst oldest solids in the solar system. <https://news.wisc.edu/comet-dust-reveals-unexpected-mixing-of-solar-system/2019/11/14>

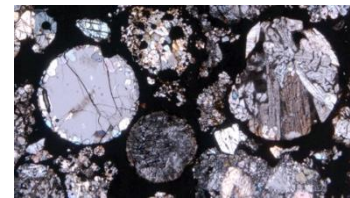


Wikipedia

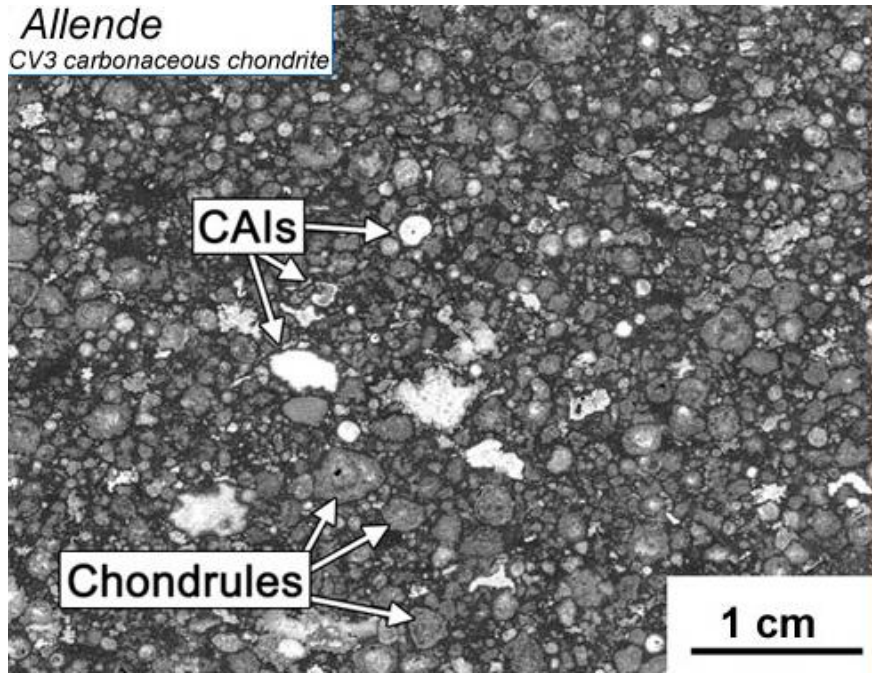
4. Pallasite

- Rich in iron, nickel
- Rich in olivine inclusions

Fragment of the core-mantle boundary of a differentiated object

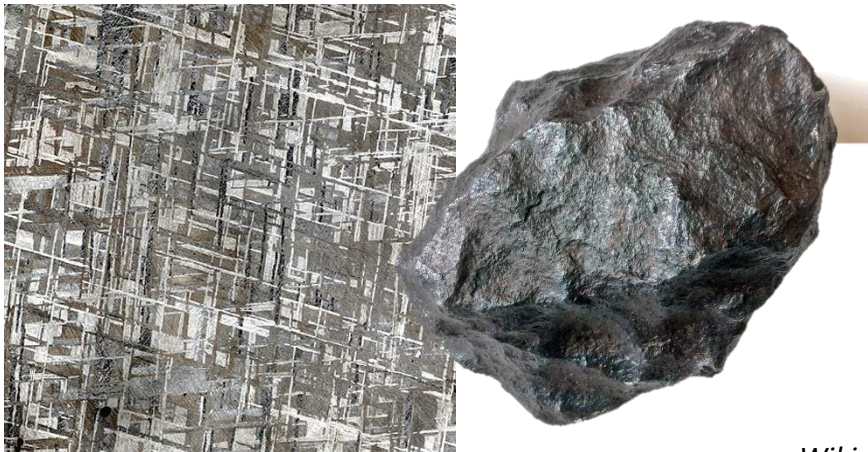


1. Carbonaceous chondrite (Allende meteorite)



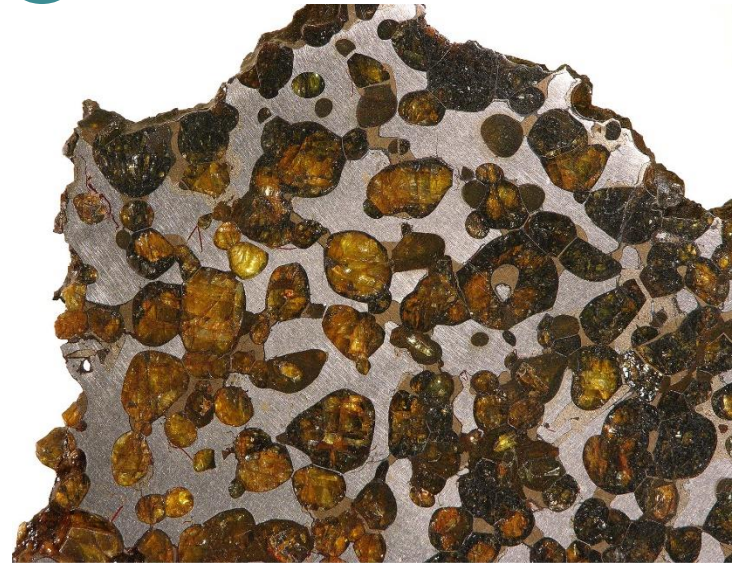
Taylor, G. J., 2012 (www.psr.d.hawaii.edu)

3. Iron meteorite (Gibeon meteorite)



Wiki

4. Pallasite



Scott, E., Goldstein, J., and Yang, J., 2010 (www.psr.d.hawaii.edu)