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## RESEARCH LETTER

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### Special Section:

Initial results of the ERG (Arase) project and multi-point observations in geospace

### Key Points:

- Unique postmidnight purple auroral rays and global Pc1/EMIC waves were observed during a CIR-driven solar wind density enhancement
- Pc1/EMIC waves were found over a wide longitudinal range extending from midnight through morning to the afternoon due to CIR arrival
- Entry of high-density solar wind plasma into the magnetotail may have created tall purple auroral rays in the sunlit ionosphere

### Supporting Information:

- Supporting Information S1
- Movie S1
- Movie S2

### Correspondence to:

K. Shiokawa,  
shiokawa@nagoya-u.jp

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## Purple Auroral Rays and Global Pc1 Pulsations Observed at the CIR-Associated Solar Wind Density Enhancement on 21 March 2017

K. Shiokawa<sup>1</sup> , M. Ozaki<sup>2</sup> , A. Kadokura<sup>3,4,5</sup> , Y. Endo<sup>4,6</sup> , T. Sakanoi<sup>4,6</sup> , S. Kurita<sup>1</sup> , Y. Miyoshi<sup>1</sup> , S.-I. Oyama<sup>1,3,7</sup> , M. Connors<sup>8</sup> , I. Schofield<sup>8</sup> , J. M. Ruohoniemi<sup>9</sup> , M. Noše<sup>1</sup> , T. Nagatsuma<sup>10</sup> , K. Sakaguchi<sup>10</sup> , D. G. Baishev<sup>11</sup> , A. Pashinin<sup>12</sup> , R. Rakhmatulin<sup>12</sup> , B. Shevtsov<sup>13</sup> , I. Poddelsky<sup>13</sup> , M. Engebretson<sup>14</sup> , Tero Raita<sup>15</sup> , Y.-M. Tanaka<sup>3,4,5</sup> , M. Shinohara<sup>16</sup> , M. Teramoto<sup>1</sup> , R. Nomura<sup>17</sup> , A. Fujimoto<sup>18</sup> , A. Matsuoka<sup>19</sup> , N. Higashio<sup>17</sup> , T. Takashima<sup>19</sup> , I. Shinohara<sup>19</sup> , and Jay M. Albert<sup>20</sup>

<sup>1</sup>Institute for Space-Earth Environmental Research, Nagoya University, Nagoya, Japan, <sup>2</sup>Kanazawa University, Kakuma, Japan, <sup>3</sup>National Institute of Polar Research, Tachikawa, Japan, <sup>4</sup>Polar Environment Data Science Center, Joint Support-Center for Data Science Research Research Organization of Information and Systems, Japan, <sup>5</sup>The Graduate University for Advanced Studies (SOKENDAI), Japan, <sup>6</sup>Graduate School of Science, Tohoku University, Aoba-ku, Japan, <sup>7</sup>University of Oulu, Oulu, Finland, <sup>8</sup>Observatories, Athabasca University, Athabasca, Alberta, Canada, <sup>9</sup>Bradley Department of Computer and Electrical Engineering, Space@VT, Virginia Tech, Blacksburg, VA, USA, <sup>10</sup>National Institute of Information and Communications Technology, Koganei, Japan, <sup>11</sup>Yu.G.Shafer Institute of Cosmophysical Research and Aeronomy, Siberian Branch of the Russian Academy of Sciences, Yakutsk, Russia, <sup>12</sup>Institute of Solar-Terrestrial Physics, Siberian Branch of the Russian Academy of Sciences, Irkutsk, Russia, <sup>13</sup>Institute of Cosmophysical Researches and Radio Wave Propagation (IKIR), Far Eastern Branch of the Russian Academy of Sciences, Paratunka, Russia, <sup>14</sup>Department of Physics, Augsburg University, Minneapolis, MN, USA, <sup>15</sup>Sodankylä Geophysical Observatory, University of Oulu, Sodankylä, Finland, <sup>16</sup>National Institute of Technology, Kagoshima College, Kirishima, Japan, <sup>17</sup>Japan Aerospace Exploration Agency, Tsukuba-shi, Japan, <sup>18</sup>Kyushu Institute of Technology, Iizuka, Japan, <sup>19</sup>Institute of Space and Astronautical Science, Japan Aerospace Exploration Agency, Sagami-hara, Japan, <sup>20</sup>Air Force Research Lab/RV/BXA, Kirtland, OH, USA

**Abstract** This paper reports two unique auroral features: postmidnight purple auroral rays and global Pc1 geomagnetic pulsations, observed before the onset of the corotating interaction region (CIR) storm of 21 March 2017, at the beginning of the first campaign of the new Particles and Waves in the Inner magnetosphere using Ground-based network observation (PWING) longitudinal ground network with the Arase satellite. The purple auroral rays were observed from ~0315 to 0430 UT (~03–04 magnetic local time) in the northeastern sky at Husafell, Iceland (magnetic latitude: 64.9°N). We newly propose that the entry of high-density CIR plasma into the magnetotail created purple auroral rays in the sunlit ionosphere. Pc1 geomagnetic pulsations at frequencies of 0–0.5 Hz were observed after ~00 UT over a wide local time range, of 13 hr, from midnight to afternoon sectors at subauroral latitudes associated with CIR arrival. These results indicate preconditioning of the magnetosphere due to crossing of a CIR.

**Plain Language Summary** We report auroral ray structures, which show a unique purple color, and global geomagnetic pulsations observed on 17 March 2017. The purple auroral rays were observed in the northeastern sky at Husafell, Iceland. The geomagnetic pulsations at frequencies of below 0.5 Hz were observed over a wide longitudinal range extending from midnight through morning to afternoon sectors at subauroral latitudes. These phenomena took place associated with the arrival of the so-called corotating interaction region (CIR) in the interplanetary space, which is characterized by high-density high-speed solar wind plasma. The CIR is one of the phenomena frequently occurring during the minimum phase of the 11-year solar cycle. We suggest that the CIR can cause these unique purple auroras and global geomagnetic pulsations. The present observations also suggest a possible mechanism for dropout of radiation belt electrons due to electromagnetic ion cyclotron waves (=geomagnetic pulsations) associated with the CIR arrival.

## 1. Introduction

Corotating interaction regions (CIRs) are caused by high-speed solar winds emanating from solar coronal holes that interact with upstream slow solar winds (e.g., Tsurutani et al., 2006). CIRs are among the major structures in the solar wind that can cause geomagnetic storms in the absence of solar eruptive events,

such as coronal mass ejections (CMEs), particularly during solar quiet times. There have been several studies of the differences between geomagnetic storms associated with CMEs and CIRs (e.g., Borovsky & Denton, 2006; Fraser et al., 2010; Kataoka & Miyoshi, 2006; Miyoshi & Kataoka, 2005; Sandanger et al., 2009; Tsurutani et al., 1995; Zhou & Smith, 2015). CIR storms tend to be weaker than CME storms in terms of ring current intensities as indicated by the Dst index. CIR storms may generate radiation belt particles more effectively due to their longer durations and fluctuating interplanetary magnetic field (IMF), resulting in hazardous conditions for spacecraft (e.g., Iles et al., 2002; Miyoshi et al., 2013; Miyoshi & Kataoka, 2008; Tsurutani, 2001).

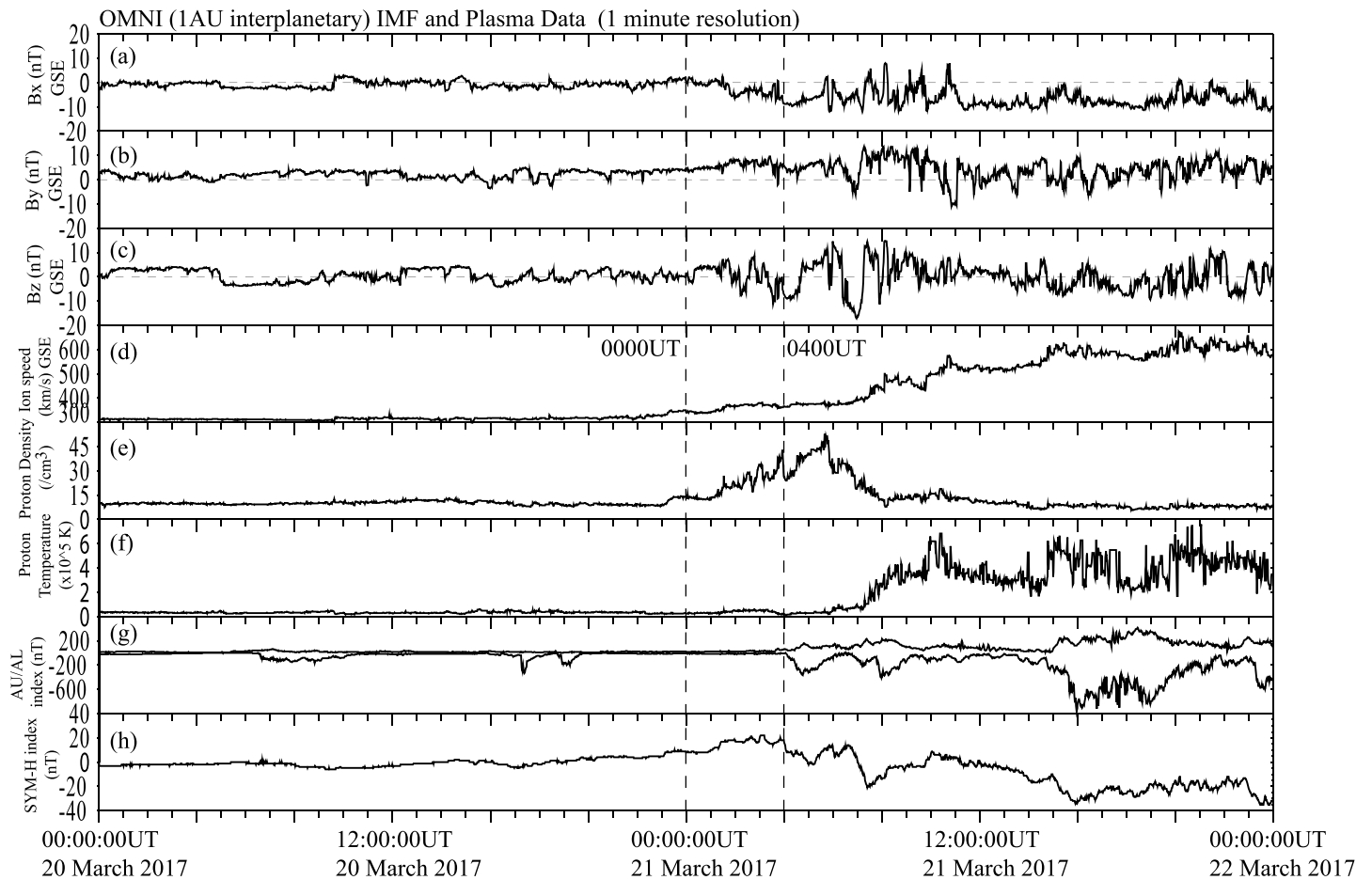
On the arrival of a CIR, solar wind monitoring satellites at 1 AU typically observe an enhancement in the solar wind plasma density, followed by an increase in the solar wind speed and development of a turbulent IMF containing a southward IMF-Bz, which initiates the main phase of a geomagnetic storm. Because the time period prior to the CIR storm is characterized by enhanced solar wind density, there might be unique effects in the Earth's magnetosphere associated with the injection of high-density plasma into the magnetotail region before the start of the main phase of the storm (i.e., preconditioning), as pointed out by Borovsky and Denton (2006). Auroral features at the beginning of the CIR storm may reflect the effect of this preconditioning, although these features have not been well studied to date. This preconditioning and subsequent increase in solar wind speed and IMF fluctuations may also contribute to the generation of various waves by setting up unstable plasma conditions in the magnetosphere.

The first observational campaign of the inner magnetosphere by the Arase satellite (and ground observation networks) was held from 20 March to 30 April 2017. A new longitudinal ground network covering all local times at subauroral latitudes became available through the "Study of dynamical variation of Particles and Waves in the Inner magnetosphere using Ground-based network observations (PWING)" project. In this paper, we used data from the first campaign observation and report two unique features: postmidnight tall purple auroral rays and global Pc1 geomagnetic pulsations, observed just before the beginning of the CIR storm of 21 March 2017. The tall purple auroral rays in the postmidnight sector suggest the entry of high-density solar wind plasma into the magnetotail flank. The unique PWING longitudinal network makes it possible to identify Pc1 pulsations occurring over a wide longitudinal range extending from midnight to the afternoon sector that are associated with the arrival of high-density solar wind plasma, even though the CIR storm had not yet begun. These Pc1 pulsations, specifically electromagnetic ion cyclotron (EMIC) waves in the magnetosphere (e.g., Fujita & Tamao, 1988; Greifinger, 1972), may contribute to the scattering loss of radiation belt particles in the inner magnetosphere to the ionosphere through wave-particle interactions.

## 2. Observation

Since April 2016, the PWING project has deployed all-sky imagers, induction magnetometers, riometers, and very low frequency receivers at eight stations at subauroral latitudes of  $\sim 60^\circ\text{N}$  magnetic latitude (MLAT) around the north geomagnetic pole, to identify the longitudinal distribution of waves and particle precipitation arising from the inner magnetosphere (Shiokawa et al., 2017). The Arase (ERG) satellite was launched on 20 December 2016 and has been in full operation since March 2017 (Miyoshi et al., 2018). Thus, the first Arase-ground campaign observation of the inner magnetosphere was held from 20 March to 30 April 2017. The combination of these ground and satellite instruments provides a unique opportunity to measure longitudinal distribution of various types of waves together with the in situ measurements of radiation belt electrons.

Figures 1a–1h show the three-component IMF, solar wind speed, density, temperature, AU and AL indices, and SYM-H index obtained from the OMNI database of National Aeronautics and Space Administration (NASA) for 20 and 21 March 2017 at the beginning of the campaign. The solar wind density started to increase from  $\sim 23$  UT on 20 March from 10 to  $50/\text{cm}^3$  at  $\sim 0530$  UT on 21 March (Figure 1e). The increase in solar wind speed was less significant during this interval. Solar wind density enhancement is a typical feature of the CIR prior to the arrival of high-speed solar wind, indicating the compression of plasmas in front of the high-speed stream (e.g., Tsurutani et al., 2006). Solar wind speed (Figure 1d) increased significantly from  $\sim 300$  km/s to more than 600 km/s after 15 UT, with increasing proton temperature in Figure 1f from 07 UT, indicating the arrival of the high-speed solar wind. The IMF Bz (Figure 1c) remained northward until  $\sim 0320$  UT, before turning southward from 0320 UT, and then again more significantly from 0630 UT; after 13 UT, the CIR storm developed, as shown by the SYM-H index (Figure 1h). At the beginning of the CIR storm, a small substorm



**Figure 1.** (a–c) Three components of the interplanetary magnetic field (IMF), (d) solar wind speed, (e) density, (f) temperature, (g) AU and AL indices, and (h) SYM-H index obtained from the OMNI database over 2 days (20–21 March 2017).

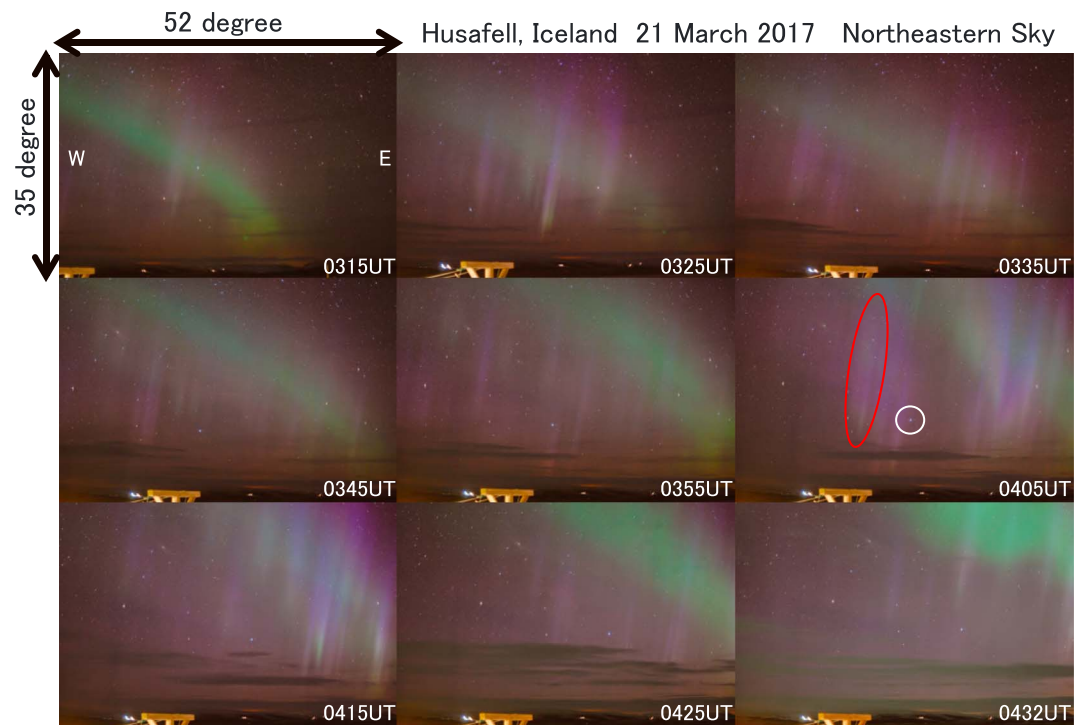
took place from 0400 UT, with a minimum AL index of  $\sim 300$  nT, associated with the initial southward turning of IMF Bz.

Figure 2 shows sequential auroral images taken by a Nikon D40 digital camera (ISO1600), with a 35 mm/F1.4 lens and an exposure time of 15 s, of the northeastern sky at Husafell (HUS, 64.67°N, 338.97°E, MLAT: 64.9°N), Iceland, at 0315–0432 UT on 21 March 2017. The complete set of images from 0308 to 0432 UT is shown in Movie S1. These images were taken manually through a window by an observer (K. S.) who stayed at HUS for the campaign observation. Notably, room light contaminated the images several times (e.g., at 0308, 0311, and 0340–0341 UT).

Clear purple auroral rays are visible in the images of Figure 2 and Movie S1 from 0308 to 0432 UT. Because no images were taken before 0308 UT, we do not know when these purple rays appeared initially. In Movie S1, the purple auroral rays showed repeated eastward (sunward) motion. The rays were brightest and most active from 0405 to 0420 UT. It is noticeable that the purple emission was observed mostly in the upper part of the auroral rays.

A stable green east-west auroral arc can also be seen with the purple auroral rays. The green arc gradually moved southward. After 0432 UT, the green aurora transformed into active diffuse auroras with pulsating patches covering the entire sky of HUS; this phenomenon was associated with the onset of a small substorm at 0400 UT (03:40 MLT at HUS).

Movie S2 shows all-sky auroral images at HUS taken by a panchromatic charge-coupled device camera (WAT-120 N+; Watec) with a fish-eye lens (YV2.2X1.4A; Fujinon) from 0000 to 0532 UT on 21 March 2017. The top and left of the images correspond to the north and east directions, respectively. We could not identify the



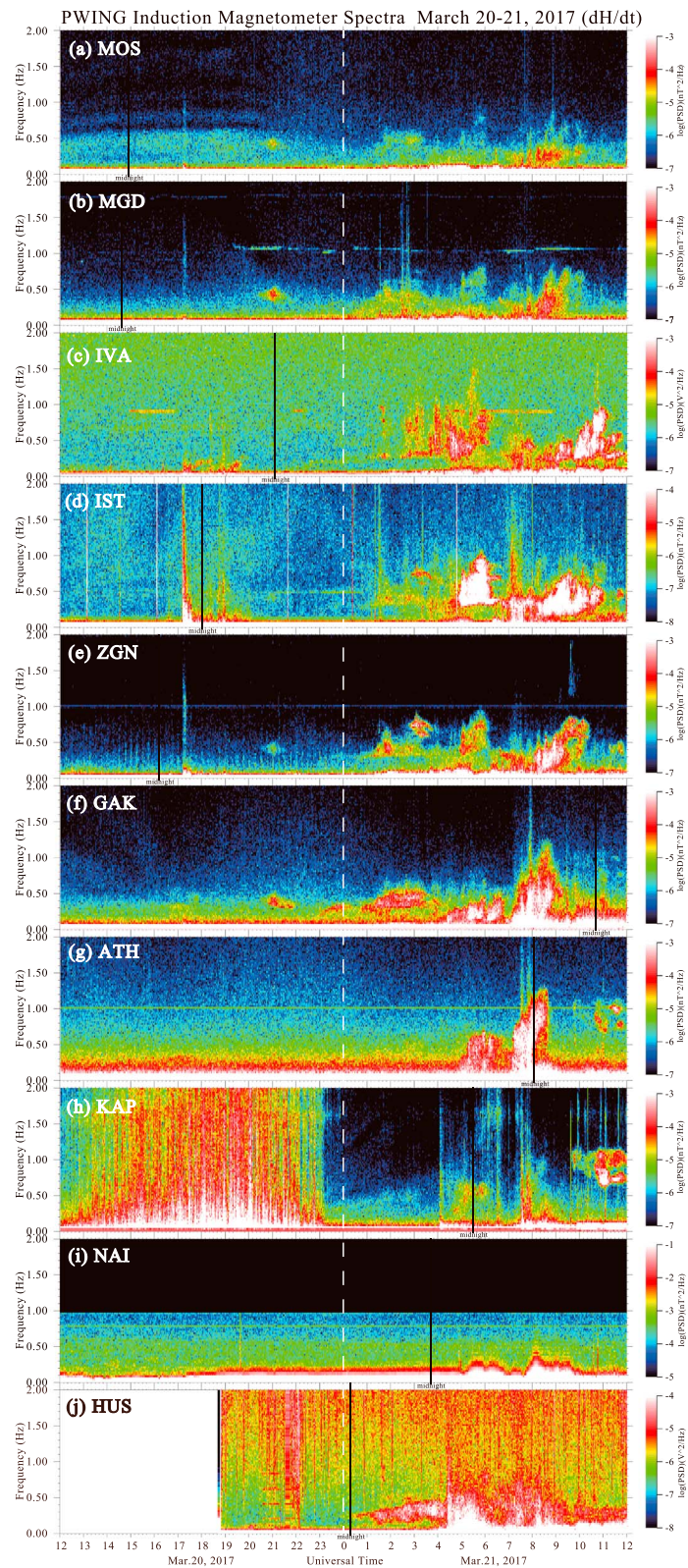
**Figure 2.** Sequential auroral images taken by a Nikon D40 digital camera with a 35 mm/F1.4 lens and an exposure time of 15 s in the northeastern sky at Husafell (HUS), Iceland, at 0315–0432 UT on 21 March 2017. The star marked by the white circle in the 0405 UT image is Alpheratz.

purple auroral rays in these images. This is probably because usually panchromatic charge-coupled device detectors have higher sensitivities at red to near-infrared wavelengths and less sensitive to the blue to ultraviolet wavelength. So it is natural that this auroral camera is less sensitive to the purple aurora. However, spatiotemporal development of the stable green arc was clearly evident in the images. The arc appeared gradually in the northern sky of HUS at ~01 UT and remained there until 0340 UT, after which it moved slowly toward the equator from 0340 to 0400 UT. After 0420 UT, the arc brightened several times from west to east, and finally converted into active diffuse auroras after 0430 UT, drifting eastward continuously.

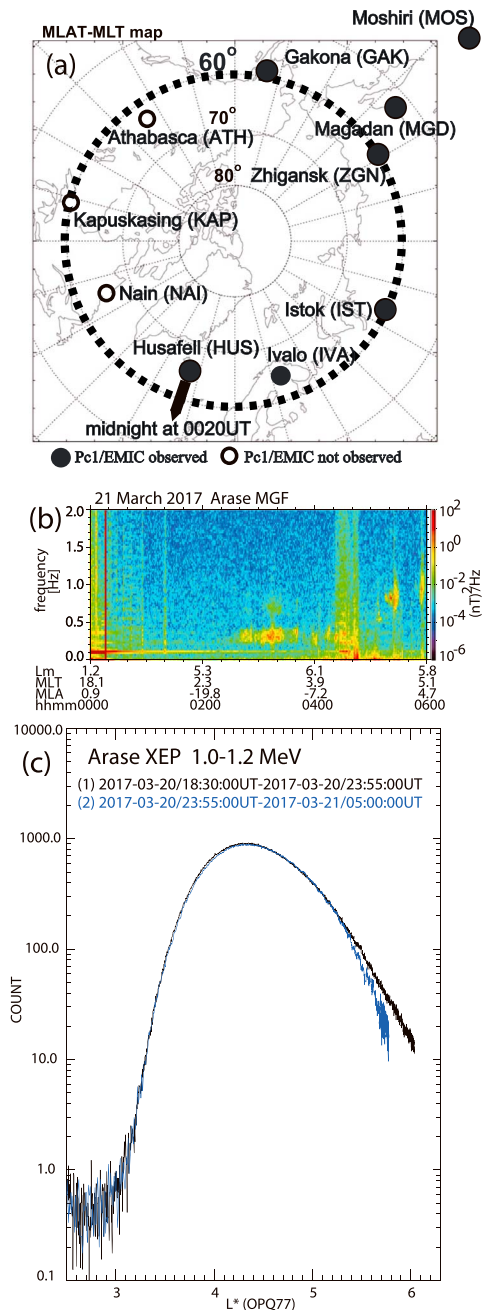
Figure 3 shows the dynamic spectra of  $H$ -component magnetic field variations observed by induction magnetometers at the PWING stations from 1200 UT on 20 March to 1200 UT on 21 March 2017. These spectra were calculated using an ordinary fast Fourier transforms every 128 s (8,192 data points). The magnetic local midnight is shown by a vertical black line in of each panel. The locations of these stations and their full names are shown in the MLT-MLAT map in Figure 4a. At 0020 UT, HUS is at magnetic midnight. The stations in Figures 3d–3f (IVA, IST, ZGN, GAK, ATH, KAP, NAI, and HUS) are longitudinally separated around the north geomagnetic pole at 60–65° MLAT, while those in Figures 3a–3c (MOS and MGD) are midlatitude stations in the longitude of ZGN. These spectra are shown in units of  $nT^2/Hz$ , with induction magnetometer sensitivities similar to those cited by Shiokawa et al. (2010, 2017), except for HUS, for which raw voltage data were used due to a lack of calibration data. The magnetometers at Nain are a part of the MACCS (Magnetometer Array for Cusp and Cleft Studies) array (Engebretson et al., 1995), with a sampling rate of 2 Hz.

From 0100 to 0330 UT on 21 March, clear Pc1 geomagnetic pulsations at frequencies of 0–0.5 Hz were observed simultaneously at all of the stations shown in Figure 3, with the exception of NAI, ATH, and KAP. This time interval corresponds to the time when the solar wind density enhancement became significant in Figure 1, while the IMF was still northward and no substorm/storm had developed in the magnetosphere. The stations that observed this simultaneous Pc1 pulsation are denoted by black circles in Figure 4a. These stations are distributed over a wide longitudinal range extending from midnight (HUS), through morning (IVA, IST, and ZGN), to the afternoon (GAK) sector. Pc1 pulsations were not observed in the evening sector





**Figure 3.** Dynamic spectra of  $H$ -component magnetic field variations observed by induction magnetometers at the PWING stations from 1200 UT on 20 March to 1200 UT on 21 March 2017. The vertical white dashed line indicates 0000 UT on 21 March. Magnetic local midnight is shown by a vertical black line in each panel. Spectra from HUS and IVA are expressed in units of  $V^2/Hz$ , while the other stations used  $nT^2/Hz$ .



**Figure 4.** (a) PWING station locations (and their full names) where the induction magnetometer measurements were carried out to obtain the dynamic spectra shown in Figure 3. The black and white circles indicate the stations where Pc1/electromagnetic ion cyclotron (EMIC) waves were observed and not observed, respectively, from 0100 to 0330 UT on 21 March 2017. (b) Dynamic spectra of magnetic field variations obtained by the fluxgate magnetometer onboard the Arase satellite near the conjugate point of HUS. The spectra correspond to the 64-Hz-sampled z-component magnetic field in DSI coordinates. Pc1 pulsations/EMIC waves were observed at frequencies of  $\sim 0.2\text{--}0.4$  Hz at  $\sim 0220\text{--}0340$  UT and over a wider frequency range up to  $\sim 1$  Hz after 0400 UT. (c) Comparison of the high-energy electron fluxes in the radiation belts measured by the extremely high-energy electron experiments (XEP) onboard the Arase satellite for two sequential paths: (1) 1830–2355 UT on 20 March (before the EMIC waves/storm) and (2) from 2355 UT on 20 March to 0500 UT on 21 March (with EMIC waves).

(ATH, KAP, and NAI). The first signature of Pc1 pulsations was detected at GAK dayside at  $\sim 2240$  UT on 20 March. Another Pc1 pulsation burst occurred at 2050–2120 UT, localized at the dayside stations of GAK and ZGN and midlatitude stations (MGD and MOS) at  $\sim 0.3\text{--}0.5$  Hz. After 0400 UT 21 March, strong Pc1 pulsations with a signature of intervals of pulsations of diminishing periods were observed globally at most of the stations associated with substorm/storm activity.

During the Pc1 pulsation event, the footprint of the Arase satellite was close to HUS. The footprint moved from  $\sim 500$  km south of HUS at 01 UT (Arase at  $L \sim 3$ ; here,  $L$  is the Roeder  $L$ -value from the [International Geomagnetic Reference Field] IGRF and the Olson–Pfitzer Quiet model [OPQ77]; Olson & Pfitzer, 1977), then northeastward within  $\sim 700$  km east of HUS at 02–04 UT (Arase at  $L \sim 5\text{--}6$  and 2–4 MLT), and finally westward approaching within 100 km of HUS at 05 UT ( $L \sim 6$  and  $\sim 4.5$  MLT). The fluxgate magnetometer onboard the Arase satellite (Matsuoka et al., 2018) observed similar Pc1 pulsations (EMIC waves in the inner magnetosphere) at frequencies of  $\sim 0.2\text{--}0.4$  Hz at  $\sim 0220\text{--}0340$  UT, and over a wider frequency range up to  $\sim 1$  Hz after 0400 UT, as shown in Figure 4b in the spectra of the z-component magnetic field in Despun Sun sector Inertia coordinate (DSI). The DSI-z component is the direction of the satellite spin axis, which is basically toward the Sun (GSM-X direction), with a deviation of  $10 \pm 5^\circ$ . This Arase observation also confirmed the existence of Pc1 pulsations/EMIC waves in the inner magnetosphere at  $L \sim 5\text{--}6$ , at the beginning of the CIR storm.

### 3. Discussion

Purple auroras have been categorized as auroral color class *f* (blue- or purple-dominant) by the International Auroral Atlas (International Union of Geodesy and Geophysics [IUGG], 1963). The IUGG (1963) described the color class *f* aurora as follows: “Blue colour arises mainly from the ionized nitrogen bands and is commonly dominant during or after very active displays. Mixture with the red atomic oxygen lines gives the appearance of purple. Blue or purple is frequently more noticeable in the upper regions of the aurora and is enhanced when the upper regions are sunlit.”

The characteristics of the present purple auroral rays shown in Figure 2 are fairly consistent with the above description of color class *f* of IUGG (1963), because they were observed dominantly in the upper regions of the aurora. We also checked whether the auroral region was sunlit or not using the star Alpheratz, marked by the white circle in the 0405 UT image in Figure 2. Alpheratz was at an elevation of  $12.163^\circ$  and an azimuth of  $224.790^\circ$  at this time. The auroral ray marked by the red circle in Figure 2 shows green (557.7 nm) emission at the bottom at the same elevation angle as Alpheratz. If we assume that this green auroral bottom edge near this star was at an altitude of 110 km, the geographic location of the aurora becomes ( $67.28^\circ\text{N}$ ,  $13.19^\circ\text{W}$ ), giving the shadow height of the sunlight at this location at 0405 UT to be 350 km. If we change the assumption of the bottom edge height of the aurora from 110 km to 150 and 170 km, the shadow height changes to 300 and 280 km, respectively. So we can say that the ionosphere above  $\sim 280\text{--}350$  km was sunlit at this time. A simple geometric calculation shows the consistency of the present event with the statement of the IUGG (1963) that the blue or purple aurora is enhanced when the upper region is sunlit. Several previous studies after

that of the IUGG (1963) confirmed that the blue emission, mainly from  $N_2^+$ (1NG) at a wavelength of 427.8 nm, comes from the sunlit auroral region through resonant scattering of sunlight by  $N_2^+$  ions produced by auroral ionization (e.g., Hunten, 2003, and references therein). Because this auroral emission was observed at higher altitudes, the  $N_2^+$  ions may have been transported upward from their site of production at lower altitudes.

An interesting difference between the present event and that described by the IUGG (1963) is that the present event was observed during relatively quiet auroras, before the start of the first substorm (0400 UT) of the CIR storm. In particular, the present purple auroral rays at 0308–0432 UT on 21 March were observed during the period of solar wind density enhancement that began 23 UT on 20 March, with prolonged (~7 hr) northward IMF from ~20 UT on 20 March associated with CIR arrival, as shown in Figure 1. Terasawa et al. (1997) pointed out, based on statistical analysis of GEOTAIL satellite data, that during prolonged northward IMF periods (approximately several hours), solar wind plasma is transported diffusively to the near-Earth plasma sheet through the magnetotail flanks, forming the cold dense plasma sheet. The high correlation of plasma densities in the solar wind and plasma sheet was also shown by Borovsky et al. (1998). Thus, the high density solar wind with prolonged northward IMF seen for the present event most likely caused an increase in the cold plasma density in the near-Earth plasma sheet. Borovsky and Steinberg (2006) showed that “a calm before the storm,” typically seen just before CIR-storm onset, enhances the contribution to the inner plasma sheet and ring current by the convection of cool dense plasma into the dipole field. This may also contribute to the creation of purple auroral rays through low-energy large-flux electron precipitation in the postmidnight sunlit ionosphere. The repeated eastward (sunward) motion of the purple auroral rays also suggests that the source of this aurora exists in the sunward convecting near-Earth plasma sheet, rather than the anti-sunward convecting solar wind or low-latitude boundary layer.

The other interesting feature associated with solar wind density enhancement during CIR arrival is the Pc1 pulsations/EMIC waves observed over a wide longitudinal range from midnight to the afternoon sectors. The unique PWING stations enable Pc1/EMIC observation at subauroral latitudes over the entire longitudinal range. Engebretson et al. (2015) showed EMIC appearance over 12-hr local times, stimulated by a gradual 4-hr rise and subsequent sharp increase in solar wind pressure on 23 February 2014. The present Pc1/EMIC waves were observed from HUS to GAK over 13 hr of local time. The magnetosphere was probably compressed gradually due to the solar wind density enhancement from ~23 UT on 20 March ( $\sim 10/\text{cm}^3$ ) to ~0530 UT on 21 March ( $50/\text{cm}^3$ ); however, the compression was more gradual than the sudden compression events reported in previous studies (e.g., Engebretson et al., 2002; Engebretson et al., 2015; Kim et al., 2017; Olson & Lee, 1983; Usanova et al., 2008). Park et al. (2016) reported that Pc1/EMIC waves in a geosynchronous orbit can be generated even by small solar wind pressure enhancements under magnetically quiet conditions.

Before this CIR arrival, the magnetosphere was fairly quiet with only some small substorms (minimum AL index of ~300 nT) occurring from 18 to 20 March. The provisional Dst index was within  $\pm 5$  nT from 18 to 20 March, except for a reading of +9 nT at 23 UT on 20 March. Thus, we would not expect there to be a sufficient number of high-energy ions to create EMIC waves in the inner magnetosphere. The entry of high-density solar wind plasma into the near-Earth plasma sheet under the northward IMF condition, and the compression of the magnetosphere due to the solar wind density enhancement, may create another population of high-energy ions in the plasma sheet through betatron acceleration. Subsequent generation of EMIC waves may occur through temperature anisotropy due to perpendicular ion acceleration through the betatron process in the conservation of the first adiabatic invariant (e.g., Arnoldy et al., 2005). The Pc1/EMIC waves of the present event were observed from midnight to the morning and dayside sectors. This fact is opposite to the duskward drift of ions injected from the nightside plasma sheet and may suggest betatron acceleration of ions due to dayside magnetospheric compression. The perpendicular betatron acceleration of preexisting ring current ions may generate EMIC waves in the equatorial plane of the magnetosphere. Engebretson et al. (2015) also reported dayside-morningside enhancement of EMIC waves associated with an increase in solar wind pressure.

These Pc1/EMIC waves may contribute to the loss of radiation belt electrons in the inner magnetosphere through wave-particle interactions, as suggested in the literature (e.g., Hyun et al., 2014; Jordanova et al., 2008; Omura & Zhao, 2012; Summers & Thorne, 2003). We compared the high-energy (1.0–1.2 MeV) electron fluxes in the radiation belts measured by the extremely high-energy electron experiments (XEP; Higashio et al., 2018) onboard the Arase satellite for four sequential paths (1) at 1830–2355 UT on 20 March (before the EMIC



waves/storm) and (2) from 2355 UT on 20 March to 0500 UT on 21 March (during the presence of EMIC waves; Figure 4c). The fluxes in paths (1) and (2) did not differ much. Thus, the global Pc1/EMIC waves at 0100–0330 UT on 21 March may not appear to have contributed to the loss of radiation belt electrons, except for the decrease between (1) and (2) at  $L > 5.3$ , although the waves were observed over a wide longitudinal range.

We also estimated the loss lifetimes of high-energy electrons by the observed Pc1/EMIC waves based on the method and ion composition ratio described by Miyoshi et al. (2008). The background plasma density observed by Arase at 0100–0400 UT at  $L = 5$ –6 in the postmidnight sector near the conjugate point of HUS was relatively high at  $100$ – $200/\text{cm}^3$  as measured by the Plasma Wave Experiment/High Frequency Analyzer (PWE/HFA) instrument (Kumamoto et al., 2018). Arase was on an outbound path and was inside the plasmopause until  $\sim 04:40\text{UT}$  ( $L = 6.2$ ) when the substorm caused a sudden density decrease at the satellite location. During this event, the plasmasphere seems to have expanded significantly, probably because of prolonged quiet condition before this event. Assuming a  $100/\text{cm}^3$  plasma density at  $L = 6$  and a wave amplitude of  $0.5$  nT at  $0.2$ – $0.4$  Hz as observed by Arase at 0300–0330 UT, the estimated loss lifetimes are  $20,708$  s ( $5.8$  hr) at  $1$  MeV and  $1,385$  s ( $23$  min) at  $10$  MeV. The duration of EMIC waves at Arase (0300–0330 UT) can be due to the orbital motion of the satellite and the actual duration of EMIC waves can be much longer ( $\sim 0100$ – $0330$  UT as observed at ground stations in Figure 3). Thus, these Pc1/EMIC waves may contribute to the observed loss of radiation belt electrons between (1) and (2) at  $L > 5.3$  in Figure 4c.

#### 4. Conclusions

The first observational campaign of the inner magnetosphere by the Arase satellite and the PWING ground observation networks was held from 20 March to 30 April 2017. The new PWING longitudinal network covered all local times at subauroral latitudes. Based on this campaign, we report two unique features at the beginning of the CIR storm of 21 March 2017: postmidnight tall purple auroral rays and global Pc1 geomagnetic pulsations over 13-hr local times from midnight to the afternoon sectors. Purple auroras are known to appear during or after very active auroras (IUGG, 1963), whereas the present event occurred during the quiet time before substorm/storm activity. We suggest that the tall purple auroral rays in the postmidnight sector were caused by the entry of high-density solar wind plasma from the magnetotail flank into the near-Earth plasma sheet, creating low-energy large-flux electron precipitation in the postmidnight sunlit ionosphere during the period of *calm before the storm*. The resonant scattering of sunlight by the molecular nitrogen ions can cause purple auroras at high altitudes. The repeated eastward (sunward) motion of the purple auroral rays also suggests that the source of this aurora is in the sunward convecting near-Earth plasma sheet.

Pc1 pulsations over a 13-hr MLT interval were observed by the unique PWING longitudinal network. We suggest that the entry of high-density solar wind plasma into the near-Earth plasma sheet under northward IMF conditions and the compression of the magnetosphere due to the solar wind density enhancement may create another population of high-energy ions in the plasma sheet through betatron acceleration and subsequently generate EMIC waves through temperature anisotropy. The dawnside appearance of the Pc1/EMIC waves of the present event also indicates that they were not caused by duskward drifting ions injected from the nightside magnetotail. The Arase XEP high-energy electron fluxes and model calculations of electron scattering rate by these waves show that these Pc1/EMIC waves may contribute to the loss of radiation belt particles observed by Arase at  $L > 5.3$ .

These results indicate interesting preconditioning of the magnetosphere due to high-density solar wind at the beginning of the CIR storm. This preconditioning with high-density solar wind will contribute to the creation of unique features of CIR storms compared with CME storms. The present results also indicate that the conjugate observation of the PWING ground network with Arase is a powerful tool to investigate wave and particle processes in the inner magnetosphere on a global scale.

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