# SOME PROPERTIES OF IONIZING RADIATIONS AND RADIOACTIVE IONS IN THE NATURAL ENVIRONMENT

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#### Abstract

This paper consists of two parts. Chapter 1 is concerned with the results of measurements of ionization intensities due to the beta- and gamma-radiations near the ground carried out from 1963 through 1969. The influences of fallout radioactivity deposited on the ground surface were studied from the view point of atmospheric ionization. The beta-radiation from the fallout radioactive nuclides deposited on the ground surface decreased remarkably during the last 6 years, and recently, a newly designed ionization chamber made of platic plates was used for meauring the beta-radiation. The ionization intensity due to the gamma-radiation didn't change so remarkable as that due to the beta-radiation during the same period. Recently, spectroscopic method with NaI scintllation counter was adopted for obtaining some detailed informations on the terrestrial gamma-radiation. In the beginning of January 1967, sudden increases of the ionization intensities due to the beta- and gamma-radiations from the fallout radioactivity caused by a nuclear explosion test were measured.

Chapter 2 is concerned with the physical properties of RaA ion. To clarify the interaction between RaA ions and condensation nuclei, simultaneous measurements of the concentrations of RaA ions  $(n_A)$ , radon-222, condensation nuclei (Z) and the diffusion coefficient of condensation nuclei (D) were carried out. It was found that the correlation among them may be well expressed by the simple formula:  $q_A = \beta_A n_A Z$ . The correlation between  $\beta_A$  and D (or radius of nuclei) was also obtained. The effective attachment coefficient of RaA ions was compared with that of small ordinary ions, and it was presumed that about 25 per cents of RaA atoms formed on the decay of radon-222 seem to have some charge, and remaining 75 per cents have no charge.

### Chapter 1. The ionization intensity due to the betaand gamma-radiations

#### 1. 1. Introduction

Although the fallout radioactivity caused by nuclear explosion tests was well studied from the view points of radiology, environmental sciences and technology, biology, and medical sciences, almost all of them were mainly concerned with gross counting, nuclide analysis, local anomaly, and their effects on human body, animals, and plants, and the influences of fallout radioactivity on the ionization

intensity near the ground in the natural radiation environment have not been studied so well.

In the previous paper<sup>1.1)</sup>, the present authors (M. K., Y. I., Y. N.) published the results of measurements on the ionization intensities due to the alpha-, beta-, and gamma-radiations in the atmosphere near the ground in 1963 and 1964. The time variations for several years and the influences of the fallout radioactivity on the ionization intensity due to the naturally occuring radiations were not clarified in the paper. In the present paper, the results of measurements of the ionizations due to the beta- and gamma-radiations from the fallout radioactive nuclides deposited on the ground surface during the last 6 years were described, and the abnormal increases of the beta- and gamma-radiations from the fallout radioactivity were discussed.

As is described below, the rate of ion pair production due to the beta-radiation was so small as considerably hard to measure it with the ionization chamber described in the previous paper. A newly designed ionization chamber used for measuring ionization intensity in low level due to the beia-radiation from the ground surface will be shown. The spectroscopic method with a scintillation counter (NaI  $3''\phi \times 3''$ ) was used for studying the detailed characteristics of the terrestrial gamma-radiations near the ground, and the response matrix method was used for calculating the real energy spectrum from the pulse height distribution detected with the scintillation counter.

#### 1.2. Instruments used for measurements

The ionizations due to the beta- and gamma-radiations were described in the previous paper<sup>1-1</sup>. The ionization due to the beta-radiation decreased, and the value of the rate of ion pair prodution is about 1 J recently. Therefore, an improved equipment is necessary to measure precisely the ionization intensity due to the beta-radiation from the ground surface.

A newly designed ionization chamber used for measuring the ionization intensity due to the beta-radiation from the ground surface was designed to minimize the wall effect and other secondary effect of the beta-radiation. The chamber consists of two ionization chambers. One of the two chambers is used for measuring the beta- and gamma-radiations, and another one for the gamma-radiation only. The structures and the dimensions of every parts of two chambers are quite the same each other except materials and thickness of their windows.

Figure 1.1 shows the structure of the ionization chamber used for measuring the beta- and gamma-radiations. The chambers are constructed by acrylite (one kind of polystyene) plates of 20 mm thick with dimensions 440 mm  $\times$  440 mm  $\times$  225 mm, having a volume of 43.5 litres. All walls except the lower surface are 20 mm thick (2,200 mg/cm²), and their inner surfaces are coated with carbon acqudac to get static conductivity. The lower surface (window area of 1,936 cm²) is made of Metalumy sheet (3.4 mg/cm²). Metalumy is a commercial name of vinyl sheet coated with aluminium. The inner electrode consists of an acrylite plate with dimensions 430 mm  $\times$  430 mm, and both surfaces of the plate are perfectly covered with Metalumy sheet.

This electrode is fixed to the wall of the top with four pieces of insulators (polystyrol). Each piece is divided into two parts, and a copper plate of which area is almost the same with that of the electrode is supported by these pieces,

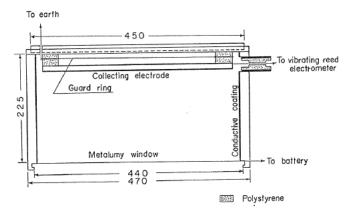


FIG. 1.1. Structure of the ionization chamber used for measuring the ionization intensity due to beta- and gamma-radiation.

and is grounded. Therefore, the copper plate plays a role of guard ring in this chamber. The collecting electrode protruded through an polystyrene plug, and was connected to a vibrating reed electrometer and an automatic recording microammeter. The window of the gamma-chamber is made of an acrylite plate of 10 mm thick which is perfectly covered with Metalumy sheet.

The acrylite plate of 10 mm thick is sufficient to absorb all beta-particles radiated from natural radioactive nuclides on the ground surface. The residual radioactivity, specially alpha-particle, contained carbon acquadac and acrylite is negligibly small as compared with those of any kinds of metals. Furthermore, the scattering of beta-particles by acrylite is very small as compared with that by metals because of its low atomic number.

Figure 1.2 shows the characteristic curves of two chambers; *i.e.*, solid line shows the characteristic curve of the beta- and gamm-chamber, and the dotted line shows that of the gamma-chamber. As is shown in these curves, 45 volts was found to be sufficient for the saturation current at the ionization level of back-ground. The saturated value in the ionization current of beta- and gamma-chamber was larger than that of gamma-chamber as shown in the figure.

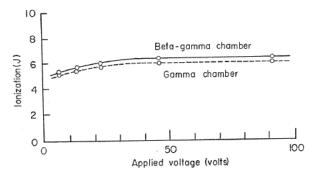


FIG. 1.2. Characteristic curves of ionization chambers used for beta- and gamma-radiations, and gamma-radiation, respectively.

The ionization intensity q (measured in the unit of ion pair/cm<sup>3</sup>/sec, J) is usually given by the relation

$$qeV_e = i (1)$$

where, all quantities are measured in the proper units, and e is the electronic charge  $1.6 \times 10^{-19}$  coul., i is the current measured with a vibrating reed electrometer, and  $V_e$  is the effective ionizing volume of the chamber. The effective ionizing volume of the chamber is obtained by the experimental and theoretical methods, and described in Reference (1.2). The results of this study show that the effective ionizing volume of the chamber amounts to 0.60 V for the betaradiation from the ground surface (in the case of 100 cm high above the ground), and amounts to 0.48 V for the beta-radiation from the atmosphere. V is the volume of the chamber, and is  $3.11 \times 10^4$  cm<sup>3</sup>.

Substituting the effective volume into the relation (1), the ionization intensity measured with the chamber may be given as follows;

$$q = \frac{i}{1.6 \times 10^{-19} \times 0.6 \times 3.11 \times 10^4} = \frac{i}{2.99 \times 10^{-15}}$$
 (2)

where, q is expressed in the unit of ion pair/cm<sup>3</sup>/sec. The rate of ion pair production shown in Fig. 2 were calculated by using the relation (2).

The difference in the saturated ionization current between the two chambers shown in Fig. 1.2 gives the value of the ionization intensity to be 0.8 J. The difference seems to be caused by the beta-radiations from the floor inside a concrete building.

Figure 1.3 shows the arrangement of the two chambers to measure the ionization intensity due to the beta-radiation from the ground surface. The collecting electrodes of two chambers are joined with a brass rod which was provided with "T" fitting to allow connection with the pre-amplifier of a vibrating reed electrometer. The interconnecting "T" arrangement was enclosed in grounded metal shield making contact with both guard rings. The electrometer was operated in conjunction with an TOA-DEMPA recording microammeter. As described above, 45 volts was found to be sufficient for saturation at these levels of ionization.

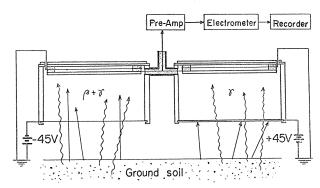


FIG. 1.3. Arrangement of the two chambers for measuring the ionization intensity due to beta-radiation from the ground.

Opposite polarity was applied to each chamber, so that the measured current was proportional to the difference in ioization current between the two chambers. To ascertain that the difference in the ionization current between the two chambers caused by the irradiation of the gamma-radiation is negligibly small, the joint chambers were irradiated by the gamma-radiation from Ra-226 of 5 mCi. The difference in ionization current between the two chambers was found to be always less than 1.5 per cents of the ionization current in each chamber. This result seems to show that the joint chambers can be used for measuring the difference in ionization current between the two chambers. Therefore, the current measured with the joint chambers is proprtional to the ionization intensity due to the beta-radiation penetrated into the chamber.

#### 1.3. Measurements and Results

#### i) Results of measurements for 6 years

The measurements of the ionization intensities due to the beta-, and gamma-radiations from the ground surface were continuously carried out inside and outside of the campus of Nagoya University from 1963 to 1969. The amplitudes of diurnal variation curves of both ionization intensities were small. As shown in

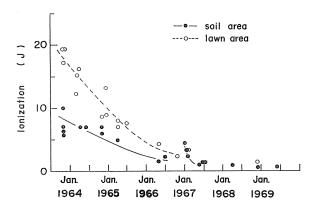


FIG. 1.4. The variation of the ionization intensity due to the beta-radiation.

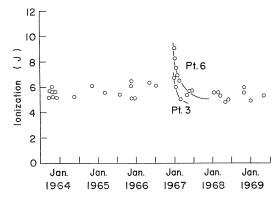


FIG. 1.5. The variation of the ionization intensity due to the gamma-radiation.

TABLE 1.1. The results of measurements of the ionization intensities after 1964

Places	Period	Surface Conditions	Beta- radiation	Gamma-radiation (including (Cosmic Ray)	
		Conditions	(J)	(J)	(µR/hr)
Kyoto Univ.	April. 1965	soil moss lawn	5.2 6.8 7.6	7.0 7.0 7.3	12.1 12.1 12.6
Nagoya Univ. Pt 6  2  6  7  5 3 2 6 2 6	June, 1964 Nov., 1964 Dec 1964 April, 1965 Aug., 1965 Dec., 1955	moss grass soil  " lawn grass soil grass soil	7.2 7.0 13 8.6 — — — — — — 1.7 4.2 2.5	5.1 6.0 — 5.5 5.4 6.0 6.4 5.1 5.1 6.3 —	8.8 10.3 — 9.5 9.3 10.3 11.0 8.8 8.8 10.9
Kyoto Univ.	Nov., 1966	lawn	2.3	7.2	12.4
Nagoya Univ. Pt. 6  " " " " " " " " " " " " " " " " " "	Jan. 6, 1967 Jan. 11, 1967 Jan. 14, 1967 Jan. 18, 1967 Jan. 24, 1967 Jan. 30, 1967 Feb. 14, 1967 Jan. 18, 1967 Jan. 18, 1967 Jan. 18, 1967 Jan. 19, 1967 Jan. 30, 1967 Feb. 14, 1967 Jan. 30, 1967 Feb. 14, 1967 May 17, 1967 June 18, 1967 June 23, 1967 June 26, 1967 Feb. 2, 1968 Feb. 6, 1968 Feb. 9, 1968 May 2, 1968 May 9, 1968 May 9, 1968 Mov. 20, 1968 May 9, 1968	soil  ""  lawn  soil  soil  lawn  grass  soil  lawn  lawn  soil soil  soil soil	4.5	9.0 8.3 8.3 7.5 7.1 6.9 6.7 6.0 5.9 5.0 5.4 5.8 5.6 5.6 5.6 5.6 5.6 5.6 5.1 6.0 5.9	15.5 14.3 14.3 12.9 12.2 11.9 11.2 11.5 10.3 10.2 8.6 9.3 10.0 10.0 9.6 9.6 9.8 8.5 8.8 10.3 9.6 9.6 9.6 9.3

3rd Chinese nuclear test explosion: May 9, 1966. 5th Chinese nuclear test explosion: Dec. 28, 1966. 6th Chinese nuclear test explosion: June 17, 1967.

Fig. 1.4, the mean value of the ionization intensity due to the beta-radiation from the ground surface decreased gradually from 1963 through 1969. It was about 20 J in the grass area and 10 J in the soil area in the fall of 1963, and decreased to 2 J in the grass area and 1 J in the soil area in the beginning of 1969.

The decrease of the beta-radiation seems to be due to the decrease of fallout radioactivity dpeosited on the ground surface, because the additional nuclear ex-

plosion tests were not carried out often after 1663. The fact that the intensity of the beta-radiation in the grass area was higher than that in the soil area seems to be due to the accumulation of the fallout radioactivity caused by the nuclear explosion tests on the leaves and roots of the grasses.

Figure 1.5 shows the variation of the ionization intensity due to the gamma-radiation on the campus from 1963 through 1969. The variation is not so remarkable as that due to the beta-radiation. The results of measurements of the ionization intensity in several places after 1965 are shown in Table 1.1 (The results before 1964 were described in reference (1.1)).

# ii) Sudden increase of the ionization due to the beta-, and gamma-radiations from the fallout radioactivity caused by the nuclear explosion tests

As is evident from Figs. 1.4 and 1.5, sudden increases of the ionizations due to the beta-, and gamma-radiations were observed in the beginning of Jan., 1967. The variations of the ionization intensities in Jan., 1967 are shown in detail in Fig. 1.6. The 5th Chinese nuclear explosion test was carried out at Lop Nor (90° E, 40° N), Chinese People Republic on Dec. 28, 1966. Radioactive particles produced by the test were transported by the Jet Stream, and reached the central area of Japan on Dec. 30, 1966<sup>1.3)</sup>.

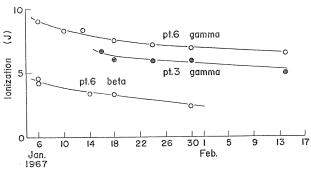


FIG. 1.6. Variations of the ionization intensities due to the beta- and gamma-radiations after the 5th Chinese nuclear explosion test.

It happened to rain on Dec. 30, and the radioactive particles were washed out by the rain and deposited on the ground surface. Unfortunately, no measurements were carried out on the campus before Jan. 6, 1967.

The ionization intensity which increased suddenly in the beginning of January decreased gradually, and reached ordinary level before March, 1967. The measured values of the ionizations due to the beta-, and gamma-radiations seems to be rather constant after March, 1967. The ionization due to the beta-radiation at one meter high above the ground surface amounts to about 1 J at present, while that due to the gamma-radiation (including the cosmic ray) amounts to about  $5\sim6$  J. According to Okano<sup>1.4)</sup>, similar sudden increases of the dose rate due to the gamma-radiation were detected in and around Tokyo on Dec. 30.

On the occation of the 3rd Chinese nuclear explosion test (May 9, 1966), dry fallout particles were detected with a G-M counter on the roof of the building of this department. Average number of dry fallout particles (higher than 10<sup>3</sup> cpm/

particle) on the roof was found to be about 18 particles/cm², and the maximum intensity of the radioactivity hold in one particle amounted to about  $5\times10^5$  cpm. However, the ionizations due to the beta-, and gamma-radiations near the ground surface didn't increase remarkably. On the occations of the 6th and 7th Chinese nuclear explosion test (June 17, 1967 and Dec. 27, 1968 respectively), neither the deposition of fallout particles caused by the tests nor the increase of the ionizations were found at all.

## iii) Nuclides of the fallout radioactivity

As described above, the ionization intensities due to the beta-, and gamma-

radiations decreased gradually from 1963 through 1966, and the ionization intensity due to the beta-radiation in the area coverd with grass or lawn exceeded that in the soil area. The intense beta-radiation seems to be caused by highly concentrating fallout radioactive debris on the grass or lawn. To study th radioactive nuclides concentrated on grass and lawn, the energy spectra of the beta-, and gamma-radiations from the grass samples on the campus were measured.

About 5 gr. of grass was sampled at Pt. 3 (one of the stations on the campus) on June 14, 1964. Figure 1.7 shows the absorption curve of the beta-rays from the grass sample. <sup>137</sup>Cs and <sup>90</sup>Sr+<sup>90</sup>Y were detected distinctly. A lump of grass with

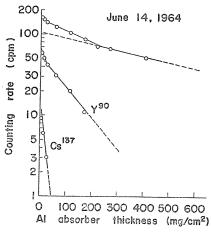


FIG. 1.7. The absorption curve of beta-rays from a lump of grass and soil sampled on the campus in 1964.

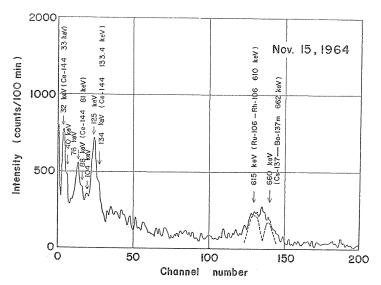


FIG. 1.8. The energy spectrum of gamma-rays from a lump of grass and soil sompled on the campus in 1964,

soil (30 cm $^2 \times 2$  cm in depth, 200 gr.) was also sampled at Pt. 3 on Nov. 15, 1964, and analyzed with a Nal(Tl) scintillation spectrometer. Figure 1.8 shows the gammaray spectrum of the sample. Photo-peaks due to  $^{144}$ Ce,  $^{106}$ Ru( $^{106}$ Rh), and  $^{137}$ Cs ( $^{137}$ mBa) were found. These nuclides were contained in the fallout radioactive debris. The obtained results coincide fairly well with those obtained by Fujita<sup>1.5)</sup>.

Shortly after the 3rd Chinese nuclear explosion test on May 9, 1966, dry fallout particles were collected on the roof of the building of this department, and radio-activity was analysed. Figure 1.9 shows the absorption curve of the beta-ray from the sample prepared by dissolving three particles in HCl solution. As is shown in the figure, <sup>144</sup>Ce and <sup>140</sup>La were detected. Figure 1.10 shows the gamma-ray spectrum of the fallout particles. Photo-peaks due to <sup>95</sup>Zr+<sup>95m</sup>Nb, <sup>132</sup>I, <sup>132</sup>Te, <sup>140</sup>La, and <sup>144</sup>Ce were detected. However, as described above, radiations from these highly radioactive debris had no influence on the ionization intensity of the air near the ground surface.

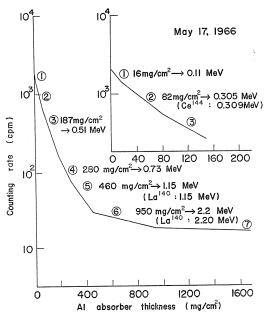


FIG. 1.9. The absorption curve of beta-rays from fallout particles caused by the 3rd Chinese nuclear explosion test.

1.4. Spectrometric method used for estimating the ionization intensity due to terrestrial gamma-radiation

The results described above were obtained by measurements with the ionization chamber without nuclide analysis. The ionization intensity due to the gamma-radiation described above includes that due to the cosmic ray. The ionization intensity due to the cosmic ray on the geomagnetic latitude of Nagoya was measured as 1.9 J. Therefore, the ionization intensity due to only terrestrial gamma-radiation may be estimated by subtracting 1.9 J from the value of the gamma-radiation described above.

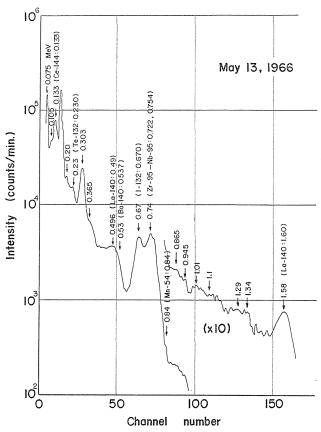


FIG. 1.10. The energy spectrum of gamma-rays from fallout particles caused by the 3rd Chinese nuclear explosion test.

To study the energy spectrum of terrestrial gamma-ray is important for detaied knowledges on terrestrial gamma-radiation. In general, it is considerably hard to estimate an energy spectrum in the air from a pulse height distribution measured with a scintillation spectrometer. The present authors used the method of transformation of the pulse height distribution obtained with a NaI scintillation spectrometer into real energy spectrum in the air by the response matrix method.

Correcting the response spectra calculated by Zerby  $et~al.^{1.6}$ , we obtained the response spectra available to analyze the experimental results. Applying the response spectra, we made a response matrix for the broad beam incident to the rectangular direction of the axis of a cylinder  $(3'' \varphi \times 3'')$  NaI scintillator), and analyzed the observed pulse height distribution by the response matrix. We denote the pulse height distribution of terrestrial gamma-radiation, response matrix, and real spectrum by  $P_0$ , R, and  $N_0$ , respectively. The following relation may be consisted among these three functions.

$$P_0 = R \times N_0 \tag{1}$$

If we adopt pulse height distribution for the first approximation of spectrum  $N_0$ , the following relation is obtained.

$$N_1 = P_0 \tag{2}$$

Multiplying  $N_1$  by the response matrix R we denote the product by  $P_1$  as follows:

$$P_1 = R \times N_1 \tag{3}$$

The *i*th element of the second approximation of spectrum,  $N_2$ , is as follow:

$$(N_2)_i = \frac{(N_1)_i}{(P_1)_i} \times (P_0)_i \tag{4}$$

Multiplying  $N_2$  by the response matrix R, we denote the product by  $P_2$  as follow:

$$P_2 = R \times N_2 \tag{5}$$

The ith element of the third approximation of spectrum,  $N_3$ , is as follow:

$$(N_3)_i = \frac{(N_2)_i}{(R_2)_i} \times (P_0)_i \tag{6}$$

The iterative cycle is repeated untill the series of trial vectors,  $N_m$ , achieves a satisfactory degree of convergence, and real spectrum of the terrestrial gammaray can be obtained.

Figure 1.11 shows the pulse height distribution obtained at 1 m high above the ground surface on the campus in Dec., 1968. Figure 1.12 shows the result of analyzing the distribution shown in Fig. 1.11. The tranformation of the result shown in Fig. 1.12 into the ionization intensity is given by the following relation.

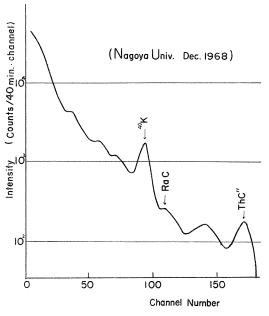


FIG. 1.11. Pulse height distribution of terrestrial gamma-ardiation,

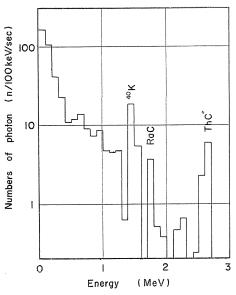


FIG. 1.12. The terrestrial gammaray spectrum analyzed by the responce matrix method,

$$q = \frac{\sum_{i=1}^{30} E_i \times N_i \times \mu_i \times L}{V \times W \times t}$$

where.

q: the ionization intensity (J)

E: the incident photon energy (MeV)

N: the number of incident photons (n/100 keV)

 $\mu$ : the energy absorption coefficient of the air (cm<sup>-1</sup>)

L: the effective length of NaI(T1) (cm)

V: the volume of a NaI(T1)

W: W-value of the air (MeV/ion pair)

t: the measuring time (sec)

The comparison between the ionization intensity measured with the ionization chamber of plastic wall, 30 litres in volume, and that measured with the scintillation spectrometer is made. The result is shown in Table 1.2. As shown in this table, the former coincides fairly well with the latter. Subtracting the primary components of Thorium series, Uranium series, and <sup>40</sup>K successively, we can obtain the scattered component of terrestrial gamma-radiation by soil. The contribution of each series to the total ionization intensity can be also obtained from the result. Their contributions are shown in Table 1.3 and Fig. 1.13. The dose build up factor of soil, shown in Table 1.3, coincide well with that obtained by Hulqvist<sup>1.5)</sup>.

TABLE 1.2. Comparison of observed values obtained by the chamber method with values obtained by the present method

Location	Present method	ion chamber		
		(including cosmic-ray)	( <sup>7</sup> -radiation) Only	
No. 1	2.53 (J)	4.6 (J)	2.7 (J)	
2	3.50	5.6	3.7	
3	2.45	4.8	2.9	
4	2.65	4.7	2.8	
5	2.85	5.0	3.1	

TABLE 1.3. Constitution of terrestrial gamma radiation obtained by the dresent method

The state of the s		Number of Photons* (sec-1)	$(MeV. sec^{-1})$	ionization (J)
Primary component Scattered con Build up fact	total of P.C.	19.6 30.5 21.5 71.6 371 6.2	28.6 26.0 29.5 84.1 85 2.0	0.439 0.450 0.413 1.302 1.41 2.1

<sup>\*</sup> Numbers and total energies of photons passing through NaI Scintillator (3"  $\varphi \times 3$ ").

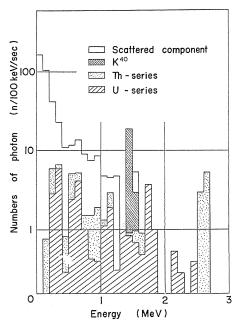


FIG. 1.13. Detailed profile of terrestrial gamma-radiation.

#### Reference

- 1.1) Kawano, M., Y. Ikebe, Y. Nakashima, and K. Shimizu (1965): Some properties of naturally occurring radiations and radio-active ions in the atmosphere, Memoirs of the Faculty of Engineering, Nagoya Univ., 17, 125-159.
- 1.2) Iida, T. and M. Kawano (1970): Measurements of the ionization of the air by the betaradiation near the ground, to be published.
- 1.3) Yoshikawa, K., A. Fujita, N. Murayama, T. Matsunami, and T. Mamuro (1968): Fallout particles in the ground level air from the Chinese nuclear explosion of December 28, 1966, Journ. Geophys. Res., 73, 3637-3641.
- 1.4) Okano, S. (1967): Private comunication.
- 1.5) Fujita, A. (1965): Fallout in soil particles, Ann Rep. Radiation Centre, Osaka Pref., 6, 29-32.
- 1.6) Zerby, C. D. and H. S. Moran (1961): Calculation of the pulse-height responseof NaI(Tl) scintillation counters, ORNL-3169.
- 1.7) Minato, S. and M. Kawano (1970): to be published.
- 1.8) Hultqvist, B. (1956): Studes on naturally occurring ionization radiatiins. Kungl. Svenska Vetenskapskademiens Handlinger, 6(3), 7-32.

#### Chapter 2. Interaction between RaA ions and condensation nuclei

#### 2.1. Introduction

The RaA ions formed on the decay of radon-222 seem to be the most fundamental form of the naturally occuring radioactive aerosols. The mobility distribution of the naturally occuring radioactive ions were measured by Wilkening. Wilkening, Kawano, and Lane<sup>2,2)</sup>. Bricard<sup>2,3)</sup> also measured the mobility distribution of RaA ions formed on the decay of radon-222 which was emanated from Ra-solution. According to the result obtained by Bricard, the mobilities of primary ions were measured to be about 2.2 cm²/sec. volt, which coincide fairly well with the value assumed by Chamberlain and Dyson<sup>2,4)</sup>.

These radioactive ions attach to aerosol particles, and form "radioactive aerosols". The attachment theory of radioactive ions has been made by several authors<sup>2,5)-2,9</sup>, and the size distribution of natural radioactive aerosols have also been measured by many authors<sup>2,10-2,14</sup>). However, few measurements have been made concerning the attachment process of natural radioactive ions<sup>2,15</sup>).

In the previous paper<sup>2,16)</sup>, we reported the interaction between the natural radioactive ions and condensation nuclei. The correlation among the concentration of RaA ions  $(n_A)$ , the concentration of condensation nuclei (Z), and the production rate of RaA ions  $(q_A)$  were well expressed by the following simple formula:  $q_A = \beta_A n_A Z$ , where  $\beta_A$  is the effective attachment coefficient between RaA ions and condensation nuclei.

In the present paper, we report the results of simultaneous measurements of  $q_A$ ,  $n_A$ , Z, and D (the diffusion coefficient of condensation nuclei). The dependencies of the effective attachment coefficient between RaA ions and condensation nuclei ( $\beta_A$ ) on the diffusion coefficient of condensation nuclei (D) and the size of condensation nuclei (P) are also discussed. Finally, P0 are compared with the effective attachment coefficient between small ordinary ions and condensation nuclei, P0.

#### 2.2. Instruments used for measurements

The instruments used for the present work are as follows:

## i) Concentration of RaA ions (n<sub>A</sub>)

The apparatus used for collecting radioactive ions is the same with the one described in the previous paper<sup>2,17)</sup>. In the present work, the copper wire is maintained at negative potential of 225 volts. The critical mobility, denoted by  $k_c$ , of the radioactive ions trapped on the wire is calculated to be about 2.2 cm²/sec. volt. The radioactive ions are collected for one hour, and at the end of the collection period, the wire is wound upon a flat spool for activity counting. A ZnS scintillation counter is used for alpha-counting, and the counted values are automatically recorded for 90 minutes after the end of the collection. As is mentioned in the previous paper, it was clarified that the radioactive nuclides trapped on the wire are mainly RaA. Then the time variation of alpha-activity of the wire may be expressed as follows:

$$I_{\alpha} = \lambda_{A} N_{A} + \lambda_{C} N_{C}$$

$$= \eta Q n_{A} \left[ \frac{\lambda_{B} \lambda_{C} (1 - e^{-\lambda_{A} T})}{(\lambda_{B} - \lambda_{C}) (\lambda_{C} - \lambda_{A})} (e^{-\lambda_{A} t} - e^{-\lambda_{C} t}) \right]$$

$$-\frac{\lambda_{c}\lambda_{A}(1-e^{-\lambda_{B}T})}{(\lambda_{c}-\lambda_{B})(\lambda_{B}-\lambda_{A})}(e^{-\lambda_{B}t}-e^{-\lambda_{c}t})$$

$$+\left\{(1-e^{-\lambda_{c}T})+\frac{\lambda_{A}\lambda_{c}}{(\lambda_{R}-\lambda_{A})(\lambda_{c}-\lambda_{B})}(e^{-\lambda_{B}T}-e^{-\lambda_{c}T})\right.$$

$$-\frac{\lambda_{B}\lambda_{c}}{(\lambda_{B}-\lambda_{A})(\lambda_{c}-\lambda_{A})}(e^{-\lambda_{A}T}-e^{-\lambda_{B}T})\right\}e^{-\lambda_{c}t}$$

$$+\left.(1-e^{-\lambda_{A}T})e^{-\lambda_{A}t}\right]$$

$$(1)$$

where,

 $I_a$ : alpha-counting rate at the time t (dps)

η: collection efficiency for RaA ions

Q: flow-rate (cm3/sec)

n<sub>A</sub>: concentration of RaA ions (cm<sup>-3</sup>)

 $N_A$  and  $N_C$ : numbers of RaA and RaC atoms collected on a wire respectively

 $\lambda_A$ ,  $\lambda_B$ , and  $\lambda_C$ : decay constants of RaA, RaB, and RaC respectively (sec<sup>-1</sup>)

T: collection time (sec).

Substituting the values  $\eta=1$ ,  $Q=\left(\frac{25}{2}\right)^2\times 200$  cm³/sec,  $\lambda_A=3.79\times 10^{-3}$  sec<sup>-1</sup>,  $\lambda_B=4.31\times 10^{-4}$  sec<sup>-1</sup>,  $\lambda_C=5.86\times 10^{-4}$  sec<sup>-1</sup>, and T=7,200 sec into formula (1), and using the value 0.080 as a counting efficiency of the scintillation counter, we can estimate the concentration of RaA ions  $(n_A)$  in the air from the counting rate.

## ii) Concentration of condensation nuclei (Z)

A Pollak photoelectric nucleus counter was used for the present work. The calibration curve obtained by Nolan<sup>2,18)</sup> was used for estimating the concentration of condensation nuclei from the measured value of extinction.

## iii) Diffusion coefficient of condensation nuclei (D)

The diffusion battery methed<sup>2.19) 2.20)</sup> was used for measuring the diffusion coefficient of condensation nuclei. A metal diffusion battery with ten channels was used. The dimensions are:

Distance between two plates = 0.10 cm Width of plates = 20 cm Length of plates = 56 cm.

The total flow rate in the present work is about  $0.8 \sim 1.2 \ l/min$ . The diffusion coefficient of condensation nuclei was estimated by using Gormley's formula<sup>2.20</sup>. Corresponding particle diameter (r) was calculated by using Stokes-Cunningham formula.

## iv) Rate of RaA ion production (q<sub>A</sub>)

Although it is well known that positively charged RaA atoms are present in the atmosphere, it has not been perfectly clarified as to how many per cents of RaA atoms are positively charged and how many per cents are neutral. Therefore, we assume hereafter that " $100 \times f$  per cents of RaA atoms formed on the decay of radon-222 have a single positive charge and the remaining RaA atoms have no charge". Then we may estimate the rate of radioactive (RaA) ion pro-

duction using the concentration of radon gas in the air:

$$q_A = fq_{Rn} = f\lambda_{Rn}n_{Rn} \tag{2}$$

where,

 $q_{Rn}$ : rate of RaA atom production (cm<sup>-3</sup> sec<sup>-1</sup>)  $\lambda_{Rn}$ : decay constant of Rn-222 (sec<sup>-1</sup>)

 $n_{Rn}$ : concentration of Rn-222 (cm<sup>-3</sup>).

In the present work, the concentration of Rn-222 was measured by the charcoal trap method<sup>2,17)</sup> or the filter pack method<sup>2,21)</sup>.

### v) Concentration of small positive ions

A Gerdien type ion counter is used for measuring the concentration of small positive ions. The diameters of the outer cylinder and the inner electrode are 8 cm and 1 cm respectively, and the length of the inner electrode is 25 cm. The velocity of air flow is 200 cm/sec, and the potential difference between the cylinder and the electrode is 100 volts. For the values described above, the critical mobility of the ions is calculated to be about 1.3 cm²/sec. volt.

#### 2.3. Results of measurements

Simultaneous measurements of the concentrations of radon-222, RaA ions, condensation nuclei, and the diffusion coefficient of condensation nuclei were made in Feb., 1967 on the campus of this university. The concentration of small ordinary ions and the rate of small ion pair production were also measured.

In the case of RaA ion equilibrium, the following relation may be given:

$$\beta_A = \frac{q_A}{n_A Z} \tag{4}$$

where.

 $\beta_A$ : effective attachment coefficient between RaA ions and condensation nuclei (cm<sup>3</sup> sec<sup>-1</sup>).

The rate of RaA ion production  $(q_A)$  cannot be directly measured. We can only determine the values of  $q_{Rn}=q_A/f$  or  $\beta_A/f$ . However, as will be described in section 2.5, the value of f is estimated to be about 0.25.

According to the formula (4),  $n_A/q_A$  is expected to be inversely proportional to Z, if we assume  $\beta_A$  to be constant. In actual, the value of  $\beta_A$  is considered to be mainly controlled by the size of condensation nuclei (2r), and the time variation of r, hence  $\beta_A$  are remarkable. Figure 2.1 shows the correlation between  $n_A/q_A$  and Z. The data shown in the figure are classified into five groups\* accrding to the values of D measured simultaneously. In each group, the values of D, hence  $\beta_A$  are rather constant. As shown in this figure,  $n_A/q_A$  is almost inversely proportional to Z. Thus, if we take the variation of D into consideration, the correlation among  $n_A$ , Z, and  $q_A$  is expressed by the simple formula:  $q_A = \beta_A n_A Z$ .

The effective attachment coefficient between RaA ions and condensation nuclei

<sup>\*</sup>  $D = (5 \sim 10, 11 \sim 40, 41 \sim 60, 61 \sim 90, > 91) \times 10^{-6} \text{ cm}^2/\text{sec.}$ 

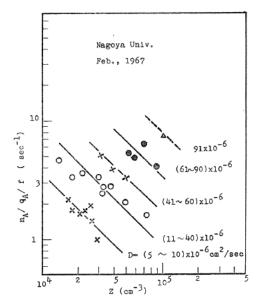


FIG. 2.1. The correlation between  $n_A/(q_A/f)$  and the concentration of condensation nuclei. Data are classified according to the value of D.

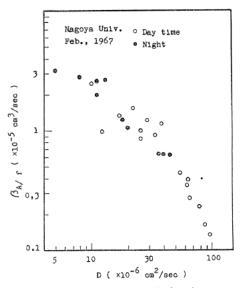


FIG. 2.2. The correlation between the attachment coefficient of RaA ions (divided by f) and the diffusion coefficient of condensation nuclei.

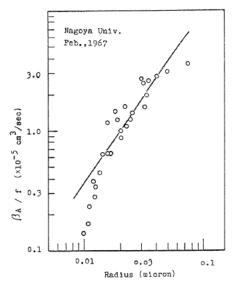


FIG. 2.3. The correlation between the attachment coefficient of RaA ions (divided by f) and the radius of condensation nuclei.

 $(\beta_A)$  may be estimated from the formula (4).  $\beta_A$  is considered to be a function of the diffusion coefficient of nuclei (D). Figure 2. 2 shows the correlation between  $\beta_A/f$  and D. As is shown in this figure,  $\beta_A/f$  correlates fairly well with D.

The average diameter of condensation nuclei (2r) may be calculated by using Stokes-Cunningham formula. In Fig. 2.3, the correlation between  $\beta_A/f$  and r is shown. Similar correlations were also obtained at Tadeshina, in Aug., 1966.

According to Keefe and Nolan<sup>2.22</sup>), the effective attachment coefficient between small ions and nuclei (b) is expressed theoretically as a function of that between ions and uncharged nuclei ( $\eta_0$ ), and  $\eta_0$  is given as follows:

$$\eta_0 = \frac{4\pi r D_0 (1 + \sqrt{\pi y})}{1 + \frac{4D_0}{rC}} \tag{5}$$

where,

r: radius of the nucleus (assumed spherical)

C: average speed of the small ions

 $D_0$ : diffusion coefficient of the small ions

$$y = \frac{\varepsilon^2}{2 \, rk \, T} \left( = \frac{2.88 \times 10^{-6}}{r} \text{ for a temperature of } 17^{\circ}\text{C} \right)$$

where  $\varepsilon$  is electronic charge, T is absolute temperature, and k is Boltamann's constant.

Our results of observations of small ions measured simultaneously with RaA ions suggested the excess of nucharged nuclei (or the lack of charged nuclei) in the atmosphere during the observation period<sup>2,23)</sup>. In this case, b closes  $\eta_0$ .

In the case of RaA ion,  $D_0=0.054~{\rm cm^2\,sec^{-1}}$  and  $C=1.75\times10^4~{\rm cm.\,sec^{-1}}$ . Inserting these values into formula (5), we obtain  $\eta_0$  as a function of r. The solid line in Fig. 2.3 represents  $\eta_0$  (arbitrary unit). The measured values shown in the figure agree fairly well with the solid line in the range of r larger than 0.015  $\mu$ .

Lassen and Rau<sup>2.5)</sup> obtained a similar equation, but without the image term, to describe the attachment of ThB to aerosol particles. Our measured values seem to agree well with the theory of Lassen and Rau in the range of r smaller than 0.03  $\mu$ .

2.4. Comparison of the attachment coefficient of RaA ions ( $\beta_A$ ) with that of small ordinary ions ( $\beta_0$ )

As is mentioned above, the concentration of small ordinary ions  $(n_0)$  and the rate of small ion pair production  $(q_0)$  were also measured simultaneously. Therefore, we can calculate the effective attachment coefficient between small ions and condensation nuclei  $(\beta_0)$  using the following formula:  $\beta_0 = \frac{q_0}{n_0 Z}$ .

On the other hand, concerning RaA ions we can only calculate the value of  $\beta_A/f$ , and the values of f, therefore  $\beta_A$  remain unknown.

Keefe and Nolan calculated  $\eta_0$  for small ordinary ions of which mobilities are 1.34 cm²/sec. volt as a function of the radius of nuclei. We can also calculate  $\eta_0$  for RaA ions of which mobilities are 2.2 cm²/sec. volt using formula (5). We denote the former by  $\eta_{00}$  and the latter by  $\eta_{0A}$ . Calculations show that  $\eta_{0A}$  coincide almost with  $\eta_{00}$  in the range of radius from 0.91  $\mu$  to 0.6  $\mu$ :

$$\frac{\beta_A}{\beta_0} \equiv \frac{\eta_{0A}}{\eta_{00}} = 1.0.$$

Thus,  $\beta_A$  is considered to coincide well with  $\beta_0$ .

In actual,  $\beta_A/f$  is found to be almost proportional to  $\beta_0$ . Figure 2.4 shows the correlations between  $\beta_A/f$  and  $\beta_0$  obtained on Nagoya University. As is shown in the figure,  $\beta_A/f$  correlates fairly well with  $\beta_0$ . The value of the ratio  $(\beta_A/f)/\beta_0$  is about 4.0 (Fig. 2.4). In the figure, the solid line represents the correlation between 4.0  $\eta_{0A}$  and  $\eta_{00}$ .

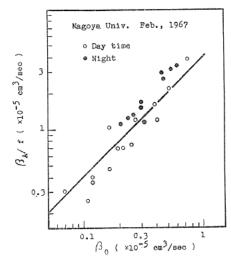


FIG. 2.4. The correlation between the attachment coefficient of RaA ions (divided by f) and that of small ordinary ions, obtained on the campus of this university.

Comparing the observed result with the theoretical one, we can obtain the following relation:  $(\beta_A/f)/\beta_0=1/f=4.0$ . Thus, the value of f should be estimated at about 0.25 on the average; therefore, in general, about 25 per cents of RaA atoms formed on the decay of radon-222 may have some charge, and the remaining 75 per cents may have no charge.

## 2.5. Correlations between each element obtained on Mt. Norikura

In the previous paper<sup>2.17</sup>, we reported the correlations among concentrations of RaA ions, small ions, and condensation nuclei. We can proceed with our studies concerning these correlations.

Figure 2.9 in Ref. (2.17) shows the correlation between the concentration of RaA ions and that of small ions. The concentration of RaA ions correlates fairly well with that of small positive ions. A similar correlation was reported previously by one of the authors<sup>2.2)</sup>. Now we assume an ionization equilibrium of small positive ions. Then the following relation may be given<sup>2.23)</sup>:

$$q_0 = \beta_0 n_0 Z \tag{6}$$

where,

 $q_0$ : production rate of small positive ions (cm<sup>-3</sup> sec<sup>-1</sup>)

 $\beta_0$ : effective attachment coefficient between small positive ions and condensation nuclei (cm³ sec<sup>-1</sup>)

no: concentration of small positive ions (cm-3).

Taking the formulae (4) and (6) into consideration, we get the following relation:

$$n_{A} = \left(\frac{\beta_{0}}{\beta_{A}} \frac{q_{A}}{q_{0}}\right). \tag{7}$$

The concentrations of radioactive ions and that of small ions are measured "simultaneously". Accordingly, the average size of the nuclei which are attached by small ions may be considered to be almost the same as that by radioactive ions. Therefore, in formula (7), we can consider the ratio  $\beta_A/\beta_0$  to be almost constant. The production rate of small ions, denoted by  $q_0$ , is almost independent of time, and we obtained:

Nagoya Univ.: 
$$q_0 = 14.9 \text{ J}$$
  
Mt. Norikura:  $q_0 = 14.5 \text{ J}$ . (8)

According to the formula (7), the concentration of radioactive ions is expected to be proportional to that of small positive ions, if we assume the production rate of radioactive ions  $(q_A)$  to be constant (see Fig. 2.9 in Ref. (2.17)). However, the time variation of  $q_A$  is considerably wide actually. Therefore, according to the formula (7), it is probable that the ratio  $n_A/q_A$  is proportional to the concentration of small positive ions  $(n_0)$ . Figure 2.5 shows the correlation between  $n_A/q_A$  and  $n_0$ . The correlation shown in Fig. 2.5 is better than that shown in Fig. 2.9 in Ref. (2.17).

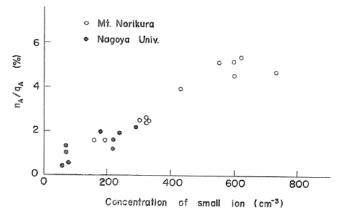


Fig. 2.5. The correlation between  $n_A/q_A$  and the concentration of small ordinary ions.

Fig. 2.10 in Ref (2.17) shows the correlation between  $n_A$  and Z. The correlation is not so evident as is expected by the formula (4).

According to the formula (4),  $n_A/q_A$  is expected to be inversely proportional to Z, if we assume  $\beta_A$  to be constant. However, as the time variation of the size of nuclei is considerably remarkable, that of  $\beta_A$  is also remarkable. Unfortunately, we did not measure the size of condensation nuclei on Mt. Norikura. But we can presume the value of  $\beta_A$  from the value of  $\beta_0$ , because  $\beta_A$  is considered to be almost proportional to  $\beta_0$ . Figure 2.6 shows the correlation between  $n_A/q_A$  and

Z. The data shown in the figure are classified into five groups\* according to the values of  $\beta_0$  estimated from the equation (6). In each group, the value of  $\beta_0$  are rather constant. As shown in this figure,  $n_A/q_A$  is almost inversely proportional to Z. Thus, if we take variation of  $\beta_0$  (hence  $\beta_A$ ) into consideration, the relation among  $n_A$ , Z, and  $q_A$  is expressed by the simple formula:  $q_A = \beta_A n_A Z$ .

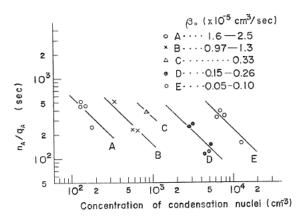


FIG. 2.6. The correlation between  $n_A/q_A$  and the concentration of condensation nuclei.

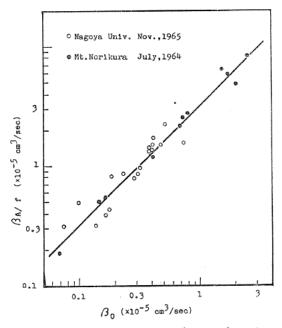


FIG. 2.7. The correlation between the attachment coefficient of RaA ions (divided by f) and that of small ordinary ions obtained on Mt. Norikura.

<sup>\*</sup>  $\beta_0 = (1.6 \sim 2.5, 0.97 \sim 1.3, 0.33, 0.15 \sim 0.26, 0.05 \sim 0.01) \times 10^{-5} \text{ cm}^3/\text{sec.}$ 

Figure 2.7 shows the correlation between  $\beta_A/f$  and  $\beta_0$  obtained on Mt. Norikura. The correlation is similar to that shown in Fig. 2.4.

#### 2.6. Conclusion

To study the interaction between RaA ions and condesation nuclei, the simultaneous measurements of the concentrations of RaA ions, condensation nuclei and radon-222 were carried out. The diffusion coefficients of nuclei were also measured. It was found that the correlation among them is well expressed by the simple formula:  $q_A = \beta_A n_A Z$ . The correlation between  $\beta_A$  and D or r was also obtained. The functional form of  $\beta_A$  coincide fairly well with those obtained theoretically by Keefe and Nolan in the range of radius larger than 0.015  $\mu$ .

The effective attachment coefficient of RaA ions was compared with that of small ordinary ions. Consquently, it was presumed that about 25 per cents of RaA atoms formed on the decay of radon-222 seemed to have some charge, and remaining 75 per cents have no charge.

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#### Reference

- 2.1) Wilkening, M. H. (1952): Natural radioactivity as a tracer in the sorting of aerosols according to mobility. Rev. Sci. Instr., 23, 13-19.
- 2.2) Wilkening, M. H., M. Kawano and C. Lane (1966): Radon-daughter ions and their relation to some electrical properties of the atmosphere. Tellus, 18, 679-683.
- Bricard, J., P. Girod et J. Pradel (1965): Radioactivitè—Spectre de mobilité des petits ions radioactifs de l'air. C. R. Ac. Sc., 260, 6587-6590.
- 2.4) Chamberlain, A. C. and E. D. Dyson (1955): The dose to the trachea and bronchi from the decay products of radon and thoron. Brit Journ. Radiol.. 29, 317-326.
- Lassen, L. and G. Rau (1960): Die Anlagerung radioaktiver Atome an Aerosole (Schwebstoffe). Zeit. Physik, 160, 504-519.
- 2.6) Lassen, L. (1961): Die Anlegerung von Zerfallsprodukten der natürlichen Emanationen an elektrisch geladene Aerosole (Schwebstoffe). Zeit. Physik, 163, 263-376.
- 2.7) Jacobi, W. (1961): Die Anlagerung von Natürlichen Radio-nucliden an Aerosolpartikel und Niedershlageselemente in der Atmoshäre. Geof. pura e appl., 50, 260-277.
- Bricard, J. (1962): La fixation des petits ions atmospherique sur les aerosols ultra-fins. Geof. pura e appl., 51, 237-242.
- Baust, E. (1967): Die Anlagerung von radioaktiven Atomen und Ionen an Aerosolteilchen. Zeit. Physik, 199, 187-206.
- 2.10) Kawano, M., and S. Nakatani (1961): Size distribution of naturally occurring radioactive dust measured by a cascade impactor and autoradiography. Geof. pura e appl., 50, 243-348.
- 2.11) Schuman, G. (1963): Investigation of radon daughters. Journ. Geophys. Res., 68, 3867.
- 2.12) Mohnen, V. und K. Stierstadt (1963): Die Verteilung der Natür lichen Radioaktivität auf das Größenspectrum des Natürlichen Aerosols. Zeit. Physik, Bd. 173, 276-293.
- 2.13) Lockart, Jr. L. B., R. L. Patterson, Jr., and A. W. Saunders, Jr. (1965): The size distribution of radioactive atmospheric aerosols. Journ. Geophys. Res., 70, 6033-6041.
- 2.14) Sekikawa, T. and S. Mochizuki (1968): On the simultaneous measurements of the particle size distribution of radioactive charged aerosols and concentration ratio of positive to negative (first report), J. Rad. Res., 9, 73-83.

- Burke, T., O. Carm and J. J. Nolan (1950): Observations on the radium A content of the atmosphere. Proc. Roy. Irish Acad., 53, Ser. A 145-161.
- 2.16) Kawano, M., Y. Ikebe, T. Nakayama and K. Shimizu (1970): The interaction between radioactive ions and condensation nuclei. J. Meteor. Soc. Japan, 48, 69-76.
- Kawano, M. et al. (1965): Some properties of naturally occuring radiations and radioactive ions in the atmosphere. Memoirs of the Faculty of Engineering, Nagoya Univ., 17, 125-259.
- Nolan, P. J. and L. W. Pollak (1946): The calibration of a photoelectric nucleus counter. Proc. Roy. Irish Acad., 51, 9-31.
- 2.19) Nolan, J. J. and V. H. Guerrini (1935): The diffusion coefficients and velocity of fall in air of atmospheric condensation nuclei. Proc. Roy. Irish Acad. 43, A 2, 5-24.
- Nolan, J. J. and P. J. Nolan (1938): Diffusion and fall of atmospheric condensation nuclei. Proc. Roy. Irish Acad., 45, 47-63.
- 2.21) Ikebe, Y., K. Shimizu and M. Kawano (1970): Measurement of the radon concentration by filter pack method. Oyobutsuri, 39, 114-118.
- Keefe, D. and P. J. Nolan (1962): Combination coefficients of ions and nuclei. Proc. Roy. Irish Acad., 62, 43-53.
- 2.23) Ikebe, Y. and K. Kawano (1970): Measurement of the effective attachment coefficient of small ions and condensation nucclei. Pure and Applied Geophys., in printing.